GROUND WATER DATA

LAKE MEREDITH SALINITY CONTROL PROJECT

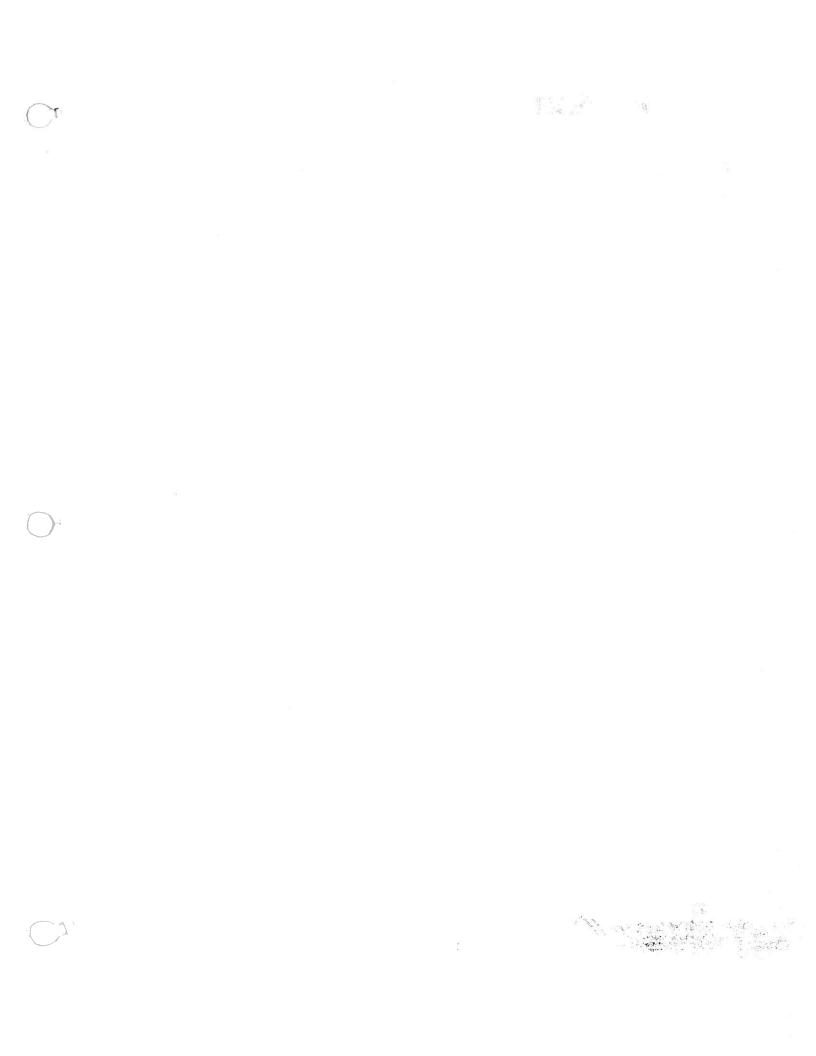
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CANADIAN RIVER MUNICIPAL WATER AUTHORITY

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EXECUTIVE SUMMARY

Ground water brine enters the Canadian River just below Ute Reservoir, near Logan, New Mexico. The brine inflow accounts for much of the excessive salt which reaches Lake Meredith, Texas, and thus is an important cause of degradation of the primary drinking water supply for 11 cities in the Texas Panhandle and South Plains which are members of the Canadian River Municipal Water Authority (CRMWA). CRMWA has a long-standing interest in developing a project to reduce or prevent the brine inflow, by intercepting and safely disposing of the brine before it reaches the river.

This notebook compiles available information on the brine distribution and flow. Important information sources include: studies of regional brine hydrogeology conducted by the Texas Bureau of Economic Geology, along with BEG's newly completed hydrogeologic cross-sections of the Logan area; studies related to Ute Reservoir conducted by the New Mexico State Engineer Office; and studies (including unpublished notes) specific to the brine inflow, conducted by or for the U.S. Bureau of Reclamation.

Selected information (e.g. stratigraphic columns; geologic and hydrologic maps; cross-sections) has been photo-copied from the source documents and organized by subject. We have built new data tables to present information on well characteristics, water levels and water quality. And, we have written short abstracts and an overall summary of the key documents.

No previous study effectively characterizes the local brine flow system. By piecing together information from many sources, and applying our own judgment, we believe that the evidence supports the following concepts.

- The Canadian River near Logan is the natural discharge zone for ground water which originates over a large part of northeastern New Mexico. Most ground water in the region flows laterally (rather than vertically) and with a gradient which includes a low but consistent net easterly component.
- Brine originates from dissolution of Permian evaporite deposits which are dominantly sodium and chloride in composition. Active dissolution occurs over a large area, including a broad front which parallels the Canadian River in the Logan area.
- Regionally, the dense brine can't flow through the relatively impermeable evaporites, so most moves downward and discharges to deep underflow beneath the Panhandle. Locally, brine moves upwards to the Santa Rosa Sandstone, the most permeable rock unit overlying the Permian evaporites in this area.
- Most of the brine found near Logan probably originates south and southwest of Logan, e.g. towards Tucumcari. It flows northeasterly and comingles with and is diluted by fresh water before discharging to the river.
- The Santa Rosa includes an upper and lower sandstone, the Trujillo and Tecovas respectively. Although no published study says so, it seems clear that the 150-foot thick Tecovas is the "brine aquifer" in the Logan area.

- The Tecovas brines are under artesian pressure. The total rate of brine inflow is 0.9 cfs, of which 0.6 cfs may originate from the evaporite dissolution and the rest is local fresh water which dilutes the brine. There are no structures which provide a known control over where the inflow occurs.
- The Tecovas slopes upwards from a depth of several hundred feet at Ute Reservoir to near the land surface at Revuelto Creek. Thus the brine aquifer is just beneath the Canadian River along the "gravel pit reach", which is the linear south-north reach 1 mile east of Logan.
- Upstream of the gravel pit reach, brine reaches the river by flowing upward across a 50-foot thick layer of shale and mudstone, then mixing with fresh water in the Trujillo and alluvium. But at the gravel pit reach, the clay rock is thin or absent, and flow is directly through the alluvium.
- Information on hydraulic properties of the brine aquifer is conflicting and in particular the permeability value utilized by USBR in their ground water model probably is in error. The model itself does not simulate a realistic flow condition.
- We interpret the water-level data as demonstrating that artesian pressure is greatest near Ute Reservoir, but the overall flow gradient is easterly toward a major discharge zone in the vicinity of the gravel pit reach.
- We interpret other data including previous geophysical studies and data on brine in the alluvium - as being consistent with the conclusion that lateral flow through the brine aquifer to the discharge zone along the gravel pit reach is a key part of the brine flow system.
- The geologic conditions similar to those which cause brine inflow near Logan also occur farther east, near the New Mexico state line. This is consistent with findings in the surface water notebook, which identified a probable "new" brine source between Logan and the state line.
- As established in the surface water notebook, the Logan brine aquifer is not as important to Lake Meredith as represented in previous reports, and a single control project will not adequately address salt inflows to the lake.

The literature does not provide a solid foundation upon which to design a brine control project. Fortunately, a stream survey recently conducted by CRMWA has generated the first set of data capable of quantifying key elements of the brine flow system. We recommend a detailed analysis of the data to confirm the importance of the gravel pit reach to the brine flow system, and to determine if this is the reach where a control project could be most effective. If CRMWA authorizes such an analysis, and it is successful, the next steps would be undertaken in the second half of 1992 and would include geophysical surveys and exploratory drilling.

1. INTRODUCTION

The Canadian River Municipal Water Authority (CRMWA) has contracted with Parkhill, Smith and Cooper Inc., and Lee Wilson and Associates Inc., to assist in implementation of the Lake Meredith Salinity Control Project. The Project objective is to reduce salinity in Lake Meredith, located NNE of Amarillo, Texas. The project method is to intercept and dispose of flow from a brine aquifer located near Logan, New Mexico.

This notebook contains readily available information on the brine aquifer, and more generally on geologic and ground water conditions along the Canadian River from Logan to Lake Meredith. In addition to a brief text summary of the information, selected data summaries and extracts are provided in the tabbed sections of the notebook. The Logan area is shown in TAB 1. This map includes the location of wells and other sites where data specific to the Project have been gathered; see subsequent discussions. Published documents from which information was extracted are listed in TAB 2.

The notebook was prepared for two different reasons: first, to compile and organize data and interpretations from many different sources, so that in the future it will be easier to find and use the information; and second, to allow the contractors to become more familiar with the Project. While the notebook summarizes what others have observed, it also contains our own interpretations of key aspects of the available data. A separate notebook provides information related to Canadian River surface water. Unlike the surface-water notebook, our interpretations are not provided in the context of specific references and informtion sources, but are consolidated together into a general discussion of the brine problem, which is presented in Chapter 3.

In general, the available information provides valuable insights regarding salinity problems at Lake Meredith and the brine aquifer near Logan. In particular, the hydrology of the brine flows to the river can be explained in a logical if somewhat generalized way. However, there is a lack of detailed information on the brine aquifer which makes it difficult to design a control project or predict project benefits. Interpretation of recently obtained stream survey data, and additional geophysical and subsurface exploration are all needed if CRMWA is to have a realistic prospect of implementing a cost-effective salinity control project.

2. LITERATURE SUMMARY

2.1 Basic data compilations and tabulations

The bibliography provided in **TAB 2** was developed through review of documents provided by CRMWA and the Texas Bureau of Economic Geology (BEG), and by cross-checking libraries such as those in our office and at the New Mexico Interstate Stream Commission; we have not done a detailed check of State libraries in Texas. Many studies were identified but not reviewed in detail because of their relatively early date or limited relevance; these are not listed on the bibliography.

For purposes of a reference notebook such as this one, it is useful to assemble basic data compilations and interpretations (such as tables which list well records, or cross-sections which interpret geologic logs) in some systematic way. Here, the choice has been to compile similar information from all germane reports into a particular tabbed section of the notebook. The organization of this information is as follows.

- TAB 3 compiles stratigraphic columns of the study area.
- TAB 4 compiles geologic, geomorphic and structure maps which include the study area.
- TAB 5 compiles geologic cross-sections of the study area.
- TAB 6 compiles hydrostratigraphic columns of the study area.
- TAB 7 compiles additional maps of the study area, especially those relating to ground-water conditions.
- TAB 8 compiles hydrogeologic cross-sections of the study area.
- TAB 9 compiles well and spring inventories from the study area. The first three tables in TAB 9 were prepared for this report, through careful integration of data from reports prepared by or for the U.S. Bureau of Reclamation (see subsequent discussion of Project Reports).
- TAB 10 compiles well log data for wells drilled in the study area.
- TAB 11 compiles water-level data from the study area. It includes an important graph showing that water levels in the brine aquifer appear to respond to to changes in storage in Ute Reservoir.
- TAB 12 compiles aquifer test data and analyses from the study area.

- TAB 13 compiles data on ground-water quality from the study area. The table in TAB 13 summarizing brine aquifer water quality was prepared for this report, through careful integration of data from reports prepared by or for the U.S. Bureau of Reclamation (see subsequent discussion of Project Reports). This tab also includes all records we have located which provide data on the quality of water in the Logan area brine aquifer.
- TAB 14 compiles maps, charts and graphs related to quality of ground water.

Note that additional tabs are provided in this notebook. For example, TAB 15 presents specific data on salt dissolution rates and river salt loads from Gustavson, et al. (1980a); and TAB 16 contains various USBR field data. Note also that all documents presented in the tabs come from reports cited in TAB 2, but only the most relevant reports are abstracted in the next sections.

2.2 Regional studies

Several reports provide useful regional perspectives on the hydrogeological conditions of the Canadian River Basin and/or the Logan area. These include: Berkstresser and Mourant (1966); Foster, et al. (1972); Spiegel (1972a, 1972b); Trauger and Bushman (1964); Kessler (1972); Gustavson, et al. (1980a); Gustavson and Finley (1985); Collins and Luneau (1986); Bassett and Bentley (1983); Orr, et al. (1985); Dutton and Orr (1986); and Dutton (1987). With limited exceptions, the following discussions present the major conclusions of these references, without criticism or comment.

Berkstresser and Mourant (1966) provide the earliest regional study that includes specific information on hydrogeologic conditions in the Logan area. The report includes a brief description of hydrogeologic units exposed in Quay County but the report otherwise lacks interpretation; the emphasis is on compilation and presentation of basic data. The data are largely adapted from earlier studies and all of it dates from the mid-1950's or earlier. Thus the information provides background on conditions that existed before the construction of Ute Reservoir. TAB 4 includes a portion of one of the geologic maps; TAB 5 includes relevant portions of two cross-sections; TAB 7 includes a portion of a water-level map; TAB 9 includes the spring inventory and parts of the well inventory; and TAB 13 includes relevant parts of the tabulation of chemical analyses of water samples.

The report identifies the Entrada Sandstone and various Cenozoic sedimentary units as the principal aquifers in the region. The Entrada is not exposed near the Canadian River, but Quaternary deposits are near or at the north side of the river from Logan to the state line. The Santa Rosa Sandstone and Chinle Formation, which are exposed along the river, were characterized as yielding 1 to 50 gpm to wells and springs.

Our review of the reference identified three noteable characteristics of ground water near Logan, as conditions existed before construction of Ute Dam.

- the Santa Rosa Sandstone southwest of Logan was under artesian pressure;
- ground water moved toward the river from both the north and south along its entire length from Logan to the state line;
- water levels north of the river near Logan were 50 to 100 feet higher than water levels a comparable distance south of the river.

Foster, et al. (1972) provide a fairly detailed summary of stratigraphy in surface exposures and subsurface rocks in east-central New Mexico. Although this study has a regional scope it includes details that are relevant to subsurface conditions along the Canadian River. Isopach maps of Triassic and older units and a contour map constructed on the top of the Precambrian surface are included in **TAB 4**.

Foster, et al. report salt deposits in the Artesia Group (Bernal Formation), San Andres-Glorieta interval and the Yeso Formation. In all cases, salt beds are absent north and west of Logan, and present to the southeast. The northern boundary of the salt-bearing area parallels the Canadian River from near Logan to the state line and lies under or just north of the river. Regional studies of the Palo Duro Basin indicate that the boundary parallel to the Canadian River is a dissolutional rather than depositional boundary.

The Precambrian surface is offset between Logan and Tucumcari by a west-northwest trending fault with displacement down to the southwest, as illustrated by the map in **TAB 4**. Near Logan, the displacement is approximately 1750 feet. Approximately 10 miles west of the state line the Precambrian surface is offset by a north-trending fault with displacement down to the east. The displacement is approximately 2500 feet where the Canadian River crosses the subsurface fault. Between these two faults the Canadian River crosses a buried horst.

<u>Spiegel (1972a, 1972b)</u> The New Mexico Geological Society Guidebook of East-Central New Mexico (Kelley and Trauger, 1972) contains several articles that describe geology in the Canadian River basin. Two articles by Zane Spiegel (who had done basic hydrogeology studies as part of the planning and design work for Ute Dam) are particularly important.

Spiegel (1972a) describes and attempts to resolve questions over the stratigraphy of Triassic rocks exposed along the Canadian River. He concludes that sandstones prominently exposed near Logan and along the river to the east are part of the Santa Rosa Sandstone and correlate eastward to the lower sandstone of the Trujillo Formation in Texas. Red mudstone and "friable white sandstone" underlie the prominent sandstone commonly exposed along the river

and are correlated on the basis of lithology and position to the Tecovas Formation in Texas. The red mudstone and white sandstone were also correlated to the Santa Rosa Sandstone. Later in the article, Spiegel confused his findings by correlating the Alibates of Texas to the San Andres Formation in New Mexico, which clearly is not correct. Well logs from Spiegel's report are included in **TAB 10**.

Spiegel (1972b) describes construction problems at dam and reservoir sites along the Pecos and Canadian Rivers, with particular attention on problems created by dissolution of gypsum or salt in the Permian section and collapse in the overlying Triassic section. The Dunes dam site located above the State Line (section 2, T13N, R35E) was described as the crest of a local anticline created by the collapse of adjacent areas. Collapse features in the reservoir site were specifically described.

Trauger and Bushman (1964) describe the hydrogeology of a portion of Quay County, New Mexico, extending generally southwest from Tucumcari. The report does not cover any area immediately along the Canadian River and most of the hydrogeologic units important in the area of the study have only limited occurrences elsewhere along the Canadian River. Only some aspects of the report will be abstracted and these mainly because we are very familiar with the report based on our work in the Tucumcari area.

The Santa Rosa Formation is present at depth within the study area and was not addressed by the authors because "the sandstone lies generally at depths of 1,500 feet or more beneath the land surface and has yielded only water of high salinity to oil-test wells...". (emphasis added)

Bedding in the Tucumcari area dips generally 1 to 3 degrees, but locally may dip as much as 30 degrees. The major structure is a large, northeast trending synclinal trough, locally modified by smaller east-trending anticlines and synclines. The convergence and divergence of the folds creates four distinct closed basins.

Trauger and Bushman attribute the folds to dissolution of bedded salt at depth and the subsidence of overlying units. Subsidence began at least by the Jurassic and is evidenced by the thickening of both the Entrada and Morrison Formations in the structural lows. Subsidence during the Quaternary is evidenced by the large thickness of alluvium (more than 500 feet) in a very local basin beneath Tucumcari. Quaternary alluvium beneath Tucumcari contains more clay than is typical in the area for other deposits of similar age. The clay beds probably were deposited in lakes or ponds analogous to existing Tucumcari Lake. A distinct sink hole, also attributed to salt dissolution in the underlying strata, was identified in the hills south of Tucumcari and described as evidence for recent collapse.

Trauger and Bushman describe only one small fault in the study area. Structural contours on the top of the Chinle Formation (TAB 4) illustrate 400

feet of structural relief across a one-mile wide, north-northeast trending zone on the west side of Tucumcari Mountain. The relief can be explained by a westward continuation of the relatively steep dip of beds on the west side of Tucumcari Mountain. A fault is not required to explain the offset.

Kessler (1972) described variations in channel geometry, gradient and sediment along the Canadian River in eastern New Mexico and the Texas panhandle. In New Mexico and western Oldham County, Texas, the river is sinuous, with relatively coarse sediments. In western Oldham County the river gradient is 7.2 feet per mile. The gradient drops to 5.5 feet per mile in the eastern panhandle, where the river is braided and less sinuous.

The Canadian River in New Mexico and western Oldham County is actively downcutting through earlier channel deposits. The river valley downstream from a point in central Oldham County preserves a fairly complete record of catastrophic flood and appears to be aggrading.

Gustavson, et al. (1980a) describe the rate, distribution and structural effects of salt dissolution in the Texas panhandle. Salt dissolution has affected parts of the Salado, Seven Rivers, San Andres, Glorieta and Clear Fork Formations (all Permian) in a zone along the Canadian River from New Mexico to near Amarillo, Texas, and along the caprock escarpment south of Amarillo. Dissolution has removed as much as 1,100 feet of bedded salt deposits.

Dissolution creates several types of structural displacements in overlying beds, including extensive breccias, breccia-filled collapse chimneys, faults, sinkholes and folded terrains. Collapse chimneys are common near Lake Meredith. Twenty seven collapse chimneys were discovered during the construction of Sanford Dam and others occur along the lake above the dam. Cross-sections through the Bonita and Alamosa faults in New Mexico show a marked thinning of the salt-bearing Artesia Group on the downthrown sides of the faults (TAB 5). Dissolution may cause or enhance the offset on the faults.

Dissolution began in the Dalhart and Anadarko Basins after deposition of the Kiowa Formation and Dakota Group (Cretaceous) and probably continued through the Tertiary. Dissolution during and after deposition of the Ogallala Formation created numerous large solution basins up to 120 square miles in area and troughs up to 50 miles long. Contours on the base of the Ogallala in the Anadarko and Dalhart Basins show broad depressions with as much as 400 feet of relief attributed to the dissolution of salt in underlying strata (TAB 4). The Canadian River flows through a series of basins and troughs that locally are more than 250 feet deep. Regionally, dissolution lowered the Great Plains surface north of the Canadian River by more than 250 feet.

Active dissolution in the Palo Duro Basin delivers a salt load of 2,800,000 tons per year to streams draining the basin. Almost half of the total load is carried by the Prairie Dog Town Fork of the Red River. The Canadian River at the Amarillo station carries nearly 700,000 cubic feet of dissolved salt a year, with about 450,000 of that originating above Tascosa, Texas and the remainder originating in the reach between Tascosa and the Amarillo station (TAB 15).

A preliminary conceptual model of salt dissolution in the Palo Duro Basin requires downward percolation of freshwater from the Ogallala Formation into the Permian section through vertical fractures cross-cutting the Dockum Group. The formation of the fractures and the extent of fracture development are influenced by prior dissolution and collapse. Dissolution is most active in the highest salt-bearing unit preserved in an area. Dissolution occurs along a steep front that gradually advances toward the center of the basin.

Gustavson and Finley (1985) describe the development of the existing drainage pattern in and around the Palo Duro Basin. Major streams that existed at the time of Ogallala deposition flowed southeast from highlands in New Mexico, directly down the land surface gradient. The present Canadian and Pecos rivers flow at high angles to the regional gradient.

Subsidence due to salt dissolution on the western margin of the Palo Duro Basin formed a series of lake depressions aligned along the north-south structural margin of the basin. The ancient Portales River was diverted into the lakes and subsequent integration of the drainage formed the present Pecos River.

Similarly, subsidence along the northern margin of the Palo Duro Basin formed a series of lakes aligned along the east-trending structural margin of the basin. Lacustrine deposits preserved at the top of the Ogallala section in the Rita Blanca drainage north of the Canadian River may be a remnant of those lakes. The southeasterly flow of a proto-Canadian River was diverted into the lake basins and subsequent integration of the system formed the existing Canadian River.

The report includes a roughly north-south geologic cross-section through Harding, Quay, Curry and Roosevelt Counties (New Mexico) that crosses Ute Reservoir west of Logan (TAB 5). The sections shows salt dissolution south of the reservoir near Tucumcari in the Artesia Group and subjacent upper San Andres Formation. It shows salt dissolution immediately below the reservoir in the lower San Andres Formation.

Collins and Luneau (1986) describe regional fracture patterns in the Palo Duro Basin and present detailed studies of joints in exposures and fractures and veins in cores. Fractures in exposures of Triassic and Permian rocks are commonly oriented parallel to faults and folds. The principal directions are east-west (275 to 295 degrees azimuth) and northwest-southeast (305 to 320

degrees azimuth). Northeast-striking fractures and faults (30 to 60 degrees azimuth) are found in Triassic units in the western part of the study area (including Quay County, New Mexico) but are rare elsewhere (TAB 4).

The in-situ principal compressive stress was measured by hydraulic fracturing of Permian beds at the SWEC Holtzclaw No. 1 well (southern Randall County, Texas). The principal compression was oriented on a northeast azimuth. Open fractures parallel to the principal stress are the most likely conduits for ground water flow.

Gypsum veins in Permian strata are normally found in units affected by salt dissolution. The veins were formed in response to collapse accompanying late-stage dissolution of underlying salts. Halite veins within the salt-bearing sequence formed soon after deposition. The orientation of some veins also indicates an influence from the regional stress system.

<u>Bassett and Bentley (1983)</u> describe the general hydrogeology and geochemistry of brines in the Palo Duro Basin. They divide the stratigraphic section into three hydrostratigraphic units on the basis of known or estimated hydraulic properties. The three units are: 1) a deep basin brine aquifer; 2) an evaporite aquitard, and 3) an upper freshwater aquifer (**TAB 6**).

The deep basin brine aquifer contains all pre-Leonardian (early Permian Major sedimentation began in the and older) strata in the basin. Pennsylvanian, synchronous with tectonism that formed the basin. Most of the aquifer consists of open-marine platform carbonates and fluvial-deltaic arkosic sandstones interbedded with mudstone. The distribution of the lithologies (TAB 4) was controlled by the structural development of the basin. Fluvial-deltaic sandstones were deposited on the basin margins from sediments derived on the adjacent highlands and from deltas extending into the basin from the northeast. Shelf and shelf-margin carbonates occur basinward from the deltaic deposits and intertongue with them. The basin center was filled with mud. The rate of subsidence diminished over time and by the late Wolfcamp the shelf-margin carbonates had prograded into the basin center and lapped over the basin-bounding uplifts. The average permeability of the deep brine aquifer was estimated at 0.02 gpd/ft².

The evaporite aquitard consists of Middle and Upper Permian (mostly Leonardian and Guadalupian) strata containing halite, anhydrite, dolomite and fine-grained siliciclastic red beds. The average permeability of the evaporite aquitard was estimated at 2×10^{-6} gpd/ft².

The fresh water aquifer consists of fluvial, deltaic and lacustrine deposits in the Triassic Dockum Group and alluvial deposits in the Tertiary Ogallala Formation. The properties of sandstones within the Dockum Group are not well known because of limited development. The approximate permeability (apparently of the sandstones alone) is 2 to 20 gpd/ft 2 . The Ogallala is well characterized and its average permeability was estimated at 200 gpd/ft 2 .

Freshwater flows in the upper aquifer generally eastward and some of it discharges at springs and seeps along the eastern escarpment of the high plains. Beds in the evaporite aquitard are exposed east of the escarpment, where dissolution of the evaporites and attendant collapse of the interlayered red beds has enhanced their permeability.

Data from drill stem tests show that the piezometric surface in the deep brine aquifer slopes generally eastward at about 6.3 feet per mile with little effect from topographic features, including the caprock escarpment. Under the existing gradient, with a permeability of 2 millidarcy and a porosity of 0.5 percent, the deep brine aquifer would flush approximately every million years.

The Wolfcampian units in the Panhandle Gas Field had an initial pressure of 435 psi, about 265 psi lower than adjacent beds toward the central basin. The hydraulic separation is created by faults that effectively isolate the reservoir area.

Brines in the deep aquifer are consistently dominated by sodium and chloride and typically have TDS concentrations between 150,000 and 170,000 mg/l (map in **TAB 14**). Brines from the carbonate and arkosic sections of the aquifer differ in their Ca/Mg ratios (higher in the clastics) and their sulfate concentrations, which are generally 2 orders of magnitude lower in the clastics than in the carbonates. The difference is attributed to sulfate reduction in the clastic deposits.

Equilibrium chemistry computer programs were used to determine the effects controlling the brine chemistry. The salinity was attributed to dissolution early in the ground water flow path of salts from the overlying evaporite sequence. Brines from both carbonates and arkose contained a Na/Cl ratio slightly lower than that expected from the dissolution of pure halite. The modification was attributed to ion exchange.

The available analyses were found to be affected by CO₂ outgassing and iron oxidation. Reconstruction of in situ conditions indicated that water from carbonate host rocks would be in equilibrium with calcite. Brines from the carbonates generally fall on the anhydrite phase boundary, but brines from the arkose are not in equilibrium with anhydrite because of sulfate reduction.

Orr. et al. (1985) provide information on variations of hydraulic conditions within the deep brine aquifer in the Palo Duro Basin.

The deep brine aquifer is generally underpressured; the brine hydraulic head is lower than indicated by a hydrostatic gradient and substantially below the land surface. The hydraulic gradient between the deep aquifer and the shallow freshwater aquifer is downward in all but the southeastern corner

of their study area. Head differences between the two aquifers (TAB 7) range from near 0 (southeast) to more that 1,800 feet (northwest, in central Oldham County).

The horizontal hydraulic head gradient within the brine aquifer (TAB 7) is generally northeastward. The gradient turns nearly due east along the south margin of the Palo Duro Basin.

Underpressuring of the brine aquifer was attributed in part to downward flow within the aquifer. Data from drill stem test, a simple hydrostatic model and a previously published flow model of the brine aquifer were combined to predict flow conditions within the aquifer. Upward flow within the aquifer was predicted in the northern part of the basin, while downward flow was expected in the central basin and east of the caprock escarpment. Flow parallel to bedding was predicted in a band around the basin center separating the areas of upward or downward flow along the basin margins from the area of downward flow in the basin center.

The variations in vertical flow components within the aquifer were attributed to lithologic changes.

<u>Dutton and Orr (1986)</u> describe hydraulic conditions and geochemical characteristics of brines in the San Andres Formation, a part of the Palo Duro Basin's evaporite aquitard.

Hydraulic heads within the San Andres are affected by topographic features and by a history of hydrocarbon production. Two maps (TAB 7) show that water levels slope from the northwest to southeast from central and eastern New Mexico to the southeast corner of the Palo Duro Basin. The heads are intermediate between heads in the Ogallala aquifer and heads in the deep brine aquifer.

Vertical flow through the San Andres occurs principally through fractured zones. Areally averaged vertical flow rates were estimated in previous model studies at 2×10^{-7} ft³/day per square foot. Average lateral flow rates were estimated at 7×10^{-8} ft³/day per square foot.

Geochemical study indicated that brines in the San Andres probably originated as evaporatively concentrated sea water (connate water). Ion exchange lowered the sodium concentration in the connate water and caused additional halite dissolution. Halite dissolution causes the Cl/Br ratio in the San Andres brine to be higher than in concentrated sea water. Sulfate reduction, which lowers sulfate/chloride ratios, and dolomitization which raises Ca/Mg ratios also affect brine composition.

A meteoric-water evolution path for the brine composition also was considered. Evolution from meteoric water to San Andres brine was considered unlikely because the path required unsubstantiated or speculative reactions.

<u>Dutton (1987)</u> reports the results of testing for hydraulic properties and chemical characteristics of water in salt dissolution zones in the Palo Duro Basin. The tests were conducted at three wells: 1) SWEC Mansfield No. 2, in the Canadian Breaks in eastern Oldham County; 2) SWEC Sawyer No. 2, on the Rolling Plains in Donley County and; 3) SWEC Holtzclaw No. 1, on the High Plains surface in southern Randall County. The SWEC Holtzclaw No. 1 well was used to test two different intervals. Water quality and water levels were reported from several additional wells.

SWEC Mansfield No. 2 was used in single-well tests to measure the transmissivity of an interval in the Seven Rivers Formation. The well was tested in a series of drawdown and recovery cycles (see data plots and curves in **TAB 12**) and results were analysed by type-curve matching or by Theis or Jacobs approximations. Transmissivity ranged from 1.5 to $5.0~\text{m}^2/\text{day}$ and permeability was calculated to range from 0.16 to 0.4 m/day. Water from the well contained 67,500 mg/l TDS and had a specific weight of 0.453 psi/ft.

Specific storage was calculated at 1×10^{-6} ft⁻¹, based on barometric efficiency (36%), neutron log porosity (30%) and the modulus of elasticity of water. The storage coefficient was calculated at $10^{-4.2}$.

SWEC Sawyer No. 2 was used in a single well test to measure the transmissivity of an interval at the top of the San Andres Unit 4 carbonate. The well was tested in a single drawdown and recovery cycle (see data plots and curves in TAB 12) and results were analysed by type-curve matching. Transmissivity varied from 2.3 to 3.7 m²/day and permeability was calculated at 0.4 to 0.6 m/day. Water from the well contained 94,900 mg/l TDS and had a specific weight of 0.461 psi/ft.

Specific storage was calculated, based on barometric efficiency (58%) using the same porosity and bulk modulus values estimated for the Mansfield well. Uncertainty over possible contribution to the well from the collapse zone overlying the production interval lead to a range of estimated storage coefficients from $10^{-4} \cdot 8$ to $10^{-3} \cdot 6$.

SWEC Holtzclaw No. 1 was used to complete drill stem tests in two intervals; one in the Seven Rivers Formation and the other including both the Salado and Tansill Formations (**TAB 12**). A type curve approximation was used to estimate the transmissivity of the Seven Rivers interval at 0.1 $\rm m^2/day$ and the permeability at 0.007 m/d. The same method was used to estimate the transmissivity of the Salado-Tansill interval at 0.0008 $\rm m^2/day$ and the permeability at $\rm 6x10^{-5}$ m/d.

The hydraulic head in each stratigraphic unit is higher than the head in underlying units, indicating downward flow throughout the section. Chemical analyses of major and minor constituents and hydrogen, oxygen and sulfur isotopic studies indicate that the water in salt dissolution zones originated

as meteoric water that entered the ground water system 16,200 to 23,500 years ago. Subsequent chemical evolution involved the dissolution of halite and gypsum, hydration of anhydrite and ion exchange.

2.3 Project Reports

A key element in any literature summary for the Project is the reports conducted by or for USBR in the 1970's and the first half of the 1980's: USBR (1979), HGC (1984a), HGC (1984b), USBR (1984) and USBR (1985). General aspects of this reports, along with specifics related to the Canadian River, are provided in the surface water notebook. For the most part, our comments on these reports are provided in Chapter 3, where we use the reports (or critique them) to help identify key findings regarding the role of ground water in discharging brine to the Canadian River near Logan.

<u>USBR (1979)</u> is an appraisal-level study that summarizes work performed from 1972 to 1978. The studies were intended to identify the sources contributing high concentrations of salts to the Canadian River between Ute Reservoir and Lake Meredith and to identify methods of alleviating salinity problems in the river.

The report includes a summary of geologic conditions along the Canadian River. TAB 3 includes the geologic column presented in the report. USBR describes the river as being entrenched as much as 100 feet into resistant sandstones in the Dockum Group, which are well exposed along the river. Upland areas north of the river are blanketed by "gravelly and sandy terrace and windblown deposits." Alluvial deposits in the valley consist of fine sand, silty sand and silt. The Canadian River downstream from Revuelto Creek is reported to be entrenched into the Permian Bernal Formation; note that this assertion is not supported by most other studies.

Beds within the Dockum Group downstream from Ute Dam dip at low angles to the west or northwest from the crest of an anticlinal flexure hinged about 1 mile downstream from the confluence of the Canadian River and Revuelto Creek. USBR indicates that the flexure is a depositional feature developed over a topographic high in the Precambrian basement, rather than a compressional or tectonic feature. USBR reports that it observed no faulting or surface evidence of solution collapse features in the Logan area.

Ground water along the Canadian River between Ute Dam and Lake Meredith is reported as generally highly mineralized, probably because of its proximity to soluble deposits within the Permian section. The highly mineralized water occurs at shallow depths in both Triassic and Permian rocks. The total thickness of the brine aquifers range from about 1,500 to 4,000 feet (p. 8). Records from a number of oil test wells indicate that the salt water is under sufficient artesian pressure to flow to the ground surface.

Water in the Canadian River, except for direct runoff, was attributed to baseflow contributions. Streams originating from above the Canadian Breaks carry relatively fresh water from the Ogallala Formation. Baseflow along tributaries in the inner valley or the main stem of the Canadian River originates from the Triassic and Permian beds. Water from the Triassic rocks is typically of fair to poor quality. Water from the Permian rocks is typically highly mineralized.

The report notes saline ground water contributions to the river below Ute Dam and attributes the salt pollution to the upward movement of mineralized water from the Permian through fractures and collapse features. They also note that despite the occurrence of saline seeps and pools along the river, there are no identifiable structural conduits which localize saline inflow. Ute Dam is reported to have caused an increase in seepage due to reservoir leakage and "possibly to hydraulic loading on the Triassic aquifers" (p. 9). Refer to the surface water notebook for information on the USBR stream surveys which identified the saline seeps.

USBR (1979) describes the results of several specific field studies. earliest investigations were in the period 1972-74, when the agency undertook investigations to locate the source(s) of saline water which reaches the Canadian River above Lake Meredith. A 1972 study sampled riverbed sands and evaluated evapotranspiration at various locations along the river. data have been requested from the Bureau of Reclamation and, if obtained, will **16**.) TAB The studies generally indicated in included salt storage and salt flushing do not contribute evapotranspiration, significantly to the salinity problem.

In 1974 the Bureau drilled four groups of five test holes into alluvium along the Canadian River from Logan to the U.S. Highway 87 bridge above Lake Meredith. (These data have been requested from the Bureau of Reclamation and, if obtained, will be included in **TAB 16**.) These showed saline conditions in New Mexico, above the State Line. Later, six groups of test holes were drilled along the channel between Ute Dam and the State Line. (These data have been requested from the Bureau of Reclamation and, if obtained, will be included in **TAB 16**.) Data from the more detailed program suggested the presence of saline inflow concentrated in an area along the channel from 2 to 5 miles below Ute Dam. This finding caused subsequent studies to concentrate exclusively on the identified inflow area.

Two relatively deep wells were drilled in 1975 near Logan, New Mexico and sampled for water quality. Drill hole 1 (DH-1) was drilled near the U.S. Highway 54 bridge at Logan, to a depth of 356 feet. The second hole (DH-2) was located 2,000 feet downstream from the Revuelto Creek confluence, to a depth of 536 feet. Drilling records are included in **TAB 10**.

DH-1 encountered water containing 4,000-5,000 mg/l sodium chloride between 80 to 155 feet. Artesian flows (30 gpm) of saline water (22,000 to >30,000

mg/l sodium chloride) at depths between 261 and 356 feet. The producing unit was a white or light tan, friable sandstone which was described as "Glorieta-like"). DH-2 encountered a minor artesian flow below a depth of 516 feet from fine sandstone and siltstone. It did not encounter a section of friable sandstone comparable to that identified in DH-1. Water from the deep zone in DH-2 contained about 8,000 mg/l sodium chloride, a concentration comparable in quality to water in the river. Neither well encountered any salt beds.

In 1978 the Bureau of Reclamation drilled a test hole (TW-1) and four observation wells (OW-1 through OW-4) near the U.S. 54 bridge south of Logan - the same area in which DH-1 had previously identified the brine aquifer. All wells were reported to be drilled in "Dockum" materials. The top of the brine aquifer was penetrated in each of the wells and correlated between them. Drilling records for the wells are included in TAB 10; TAB 5 includes two cross-sections, using lithologic logs and geophysical well logs. The aquifer was identified as a sequence of friable to well-cemented, generally fine grained or silty sandstones interlayered with shale within the Triassic section. The aquifer occupied an interval from about 200-220 feet to 340-360 feet.

USBR (1979) also summarized results from seismic refraction and surface electrical resistivity studies completed in 1975. The surveys were intended to determine the depth, thickness and extent of the brine aquifer. They reported that the seismic survey identified layering in the subsurface, with the sequence being similar to that in DH-1: alluvium, sandstone, shale, sandstone.

The resistivity survey consisted of 6 soundings in the area along the Canadian River from near Ute Dam to Revuelto Creek. (Location, contours and resistivity profiles are included in **TAB 16**.) A low-resistivity layer, tentatively considered to be the brine aquifer, was identified in four of the six soundings and was reliably characterized in three of the soundings. The layer was absent in two soundings, but only one of those two represented reliable data. The layer thickness was found to vary from 150 to more than 300 feet and to dip generally westward at about 4 degrees.

The depth to the top of the layer varied from about 450 feet at its eastern identified extent to about 800 feet near Ute Dam (elevation from about 3,400 feet to 2,900 feet); see map in **TAB 16**. The layer was absent approximately 1 mile west of Revuelto Creek, but apparently present at an unknown depth east of Revuelto Creek. One sounding produced unreliable results, but appeared to indicate a discontinuity or boundary within the brine aquifer.

The resistivity study may have identified the eastern extent of the brine aquifer found by drilling in DH-1, TW-1 and OW-1 through OW-4. Specifically, well DH-1 may be located near the eastern, upper limit of the unit. The study did not indicate the northern, southern or western extent of the aquifer. In

our judgment, the exactly location shown for the eastern boundary of the aquifer may not be accurate.

Hydraulic parameters of the brine aquifer were measured by pumping well TW-1 and measuring drawdown at observation wells OW-3 and OW-4. The well was pumped at 475 gpm for 97 hours, followed by 68 hours of recovery. Drawdown at OW-3 and OW-4 were interpreted (apparently by Theis curve matching, although this is not specified in the report) to arrive at a storage coefficient of 0.00015, a transmissivity of 2500 square feet per day and a hydraulic conductivity of 36 ft/day. If only 425 gpm of the total pumping was derived storage coefficient of 0.00013, transmissivity of 2250 square feet per day and hydraulic conductivity of 32 ft/day. The report does not include any of the measured data, hence it has not been included in the tabbed sections of this notebook. As discussed in Chapter 3, the conductivity value is suspect.

Heads in well TW-1 are artesian and typically above the level of water in the river. The well was sampled during the pumping test. The report includes only two partial analyses (sodium, chloride and sulfate) from that set (**TAB 13**). The analyses average 14,300 mg/l sodium, 22,000 mg/l chloride and 9,150 mg/l sulfate. Apparent problems with the sulfate concentration will be discussed in Section 3.

Conductance was measured at intervals during drilling of DH-1 and reported in a letter memorandum supplement to the 1979 report. The drillers found water from 0 to 71 feet similar to river water at 10,000 to 11,500 micromhos, water from 71 to 136 feet near 6,000 micromhos and water below 261 feet at 50,000 micromhos. They also reported a dry interval from 96 to 116 feet.

Conductance was measured for the artesian flow encountered during drilling DH-2. Water from the interval 466 to 516 feet had a conductivity of 16,000-17,000 micromhos. The river at that site carried water with a conductance of 17,500 micromhos.

A Bureau file memorandum dated August 8, 1975 and written by Jimmy K. Morrison provides the basis for the USBR (1979) analysis of river salt loading from the brine aquifer. The memorandum estimated the brine contribution between Ute Dam and DH-2. Prior to a stormwater runoff event, the river discharge was estimated at 2.5 cfs, with 8,000 mg/l of TDS; about 75% was from dam seepage at 400 mg/l and the remainder (about 0.6 cfs) was from ground water. These values were used to estimate that ground water contributed 31,000 mg/l of TDS, close to the 35,000 actually found at DH-1. After the storm, the river discharged an estimated 7 cfs. The author then used the original values and assumed that the additional 4.5 cfs of runoff carried 490 mg/l. That resulted in an estimated TDS in the river after runoff of 3,200 mg/l. The river actually carried about 2,800 mg/l.

USBR (1979) combined this brine inflow estimate of 0.6 cfs with the average concentrations for sodium, chloride and sulfate measured in TW-1 to

estimate the tonnage of each constituent contributed annually by the brine aquifer. In units of tons per year (tpy), the loads are 8,500 tpy sodium, 18,000 tpy chloride and 5,400 tpy sulfate, for a total of 26,900 tpy. In units of kilograms/second, the loads are 0.24 kg/s sodium, 0.37 kg/s chloride and 0.16 kg/s sulfate. The loads were compared to total loads to Lake Meredith; the details of the latter were not provided. The 26,900 tons/year is 32 percent of a total loading of sodium, chloride and sulfate of 84,095 tons/year. Individual percentages are: sodium - 31%; chloride - 44%; sulfate - 20%.

These latter percentages have been used in some subsequent studies to indicate the potential improvement in water quality that might occur if brine inflows in the Logan area could be eliminated. At p. 20, USBR indicates that any improvement would be gradual, because large amounts of salt and salty water are stored in the alluvium and will be flushed out over time. As discussed in Chapter 3, the assessment of sulfate loading is suspect.

HGC (1984a). This is the most extensive interpretation of subsurface conditions related to the Project, including both a regional summary of geologic conditions along the Canadian River from Ute Dam to Lake Meredith and a detailed study of the geology in the Logan area. The stratigraphic column presented on page 14 of the report is provided here in TAB 3. The report contains information such as tables with information related to well logs, and interpretative maps (structure contours, isopachs). The most useful of these, a structure contour map for the San Andres Formation, is included in TAB 4. TAB 5 presents three cross-sections along with a location map (these are Figures 12 through 15 in the report). Data from the report also are included in the tables which were compiled for this study: see TAB 9 and TAB 13.

The report indicates that the course of the Canadian River follows fractures and a dissolution zone along the updip edge of Paleozoic age salt deposits. The subparallel (NE) nature of the dissolution front, regional fracture patterns and flexures suggests that the dissolution of salt has been the major control upon the geologic structure in the region.

More specifically, detailed mapping of the Logan area defined two major structural trends (TAB 4). Numerous, broad folds trend northeast. South of the river these folds plunge 5 to 10 degrees southwest. North of the river the folds appear to plunge to the northeast. A large anticlinal flexure is mapped roughly parallel to the river and defined by the reversal in the plunge of the northeast-trending flexures. The courses of Revuelto Creek and Tuscocoillo Creek run roughly parallel to the northeast-trending flexures. The east-trending flexure is roughly parallel to the salt dissolution front in underlying Permian rocks.

Aerial photographs and low-level overflights were used to identify a group of depressions aligned on a roughly east-west trend through Logan and along

the Canadian River, parallel to the salt dissolution front. Field investigations confirmed the presence of collapse structures along the river.

Fracture studies indicate four dominant trends: N50-70W, N30-10W, N30-50E and N70-90E. The best development is seen in fractures that run parallel and perpendicular to the northeast-trending flexures (N30-50E and N50-70W). These fracture sets - particularly the northeast trending set - appear to be of tectonic origin and may influence the orientation of the salt dissolution front. The northeast trending set also seems to influence the courses of Revuelto Creek and the Canadian River, as both streams flow parallel to this trend through some of the area.

HGC indicates that relatively little is known about ground water conditions in the <u>Permian formations</u> of the area, except on a regional scale. Flows originate as recharge in the Sangre de Cristo Mountains and are eastward at a fairly even gradient between 15 and 20 feet per mile. Ground water in the Yeso and San Andres formations in New Mexico flows stratigraphically downward to the Wolfcamp aquifer in Texas and discharges at salt springs along and east of the caprock escarpment in Texas. Water levels in the Permian near and west of Logan are above ground level, while water levels in the Wolfcamp of Texas are generally below ground level because the water level gradient is steeper than the slope of the land surface. Permeabilities are generally low, but locally can be high due to fractures and dissolution.

Stiff diagrams showing the quality of water in the Permian units are included in TAB 14. These show that water in recharge areas typically is dominated by calcium, bicarbonate and sulfate. Sodium sometimes is a major component but chloride generally is minor. West of Logan water from the Permian section is very high in sodium and chloride as the result of halite dissolution. Sulfate is present on an equivalent basis at roughly 1/12th the chloride concentration and bicarbonate is relatively minor. Water from the Wolfcamp section in Potter County generally is similar to water from the Permian in Quay County except that sulfate is found at proportionately lower concentrations. In addition to halite dissolution, sulfate reduction and cation exchange are believed to be responsible for the compositional variations.

Information on ground water in the <u>Triassic units</u> of the area is taken mostly from Berkstresser and Mourant (1966), which is summarized above. Figure 22 in the HGC report is a water-level map for the Triassic; it is reproduced here in **TAB 7**. The flow is markedly toward the Canadian River. Hydraulic conductivity has been estimated to range from 0.25 to 2.5 feet per day. HGC estimates the Triassic transmissivity near the Canadian River at about 300 $\rm ft^2/day$.

Based on this value and on the water-level map, HGC calculates a discharge from the Triassic into the Canadian River in New Mexico of about 0.15 cfs/stream mile, or a total of 5 cfs between Ute Reservoir and State Line.

There is little additional gain from this unit in Texas (p. 51). A localized gain of about 1 cfs near the Highway 54 bridge was observed on several occasions. Stiff diagrams showing the quality of water in the Triassic units are included in **TAB 14**. They show the water to be dominated by sodium, bicarbonate and sulfate. Chloride is a relatively minor constituent. Dissolution of limestone and gypsum and cation exchange probably are responsible for the water composition. Water from the Triassic section typically has higher concentrations than the local surface water. Some occurrences of relatively high chloride concentration may be caused by mixing with deep ground water.

Based on drilling conducted by USBR, the <u>shallow brine aquifer</u> of the Logan area is located beneath a lower Triassic shale, and above a shale which is near the top of the Permian section. The saturated thickness of the aquifer is about 100 to 150 feet and the hydraulic head appears to be slightly above river level. An aquifer test indicates a transmissivity of roughly 2500 ft²/day. HGC states that to date there is no known increase in brine inflow caused by the construction of Ute Reservoir. <u>All</u> water-quality data known for the brine aquifer are included in **TAB 13**.

Water from the brine aquifer at USBR well OW-3 is a sodium-chloride brine similar to (though more dilute than) water in the Permian section. Isotopic and minor element analyses also show the water at OW-3 to be similar to saline brines in the Permian section, but with a component of water from the Triassic aquifer. The total discharge to the Canadian River from the shallow brine aquifer near Logan was estimated at 0.90 cfs (p. 94). The chloride load from the shallow brine aquifer was estimated at 0.7 kg/s. Out of the total discharge, 0.57 cfs and virtually all of the chloride load was believed to flow from the Permian section. The remaining 0.33 cfs originates as relatively fresh water from the Triassic aquifer. Refer to Chapter 3 for our modifications of these interpretations.

The channel alluvium of the Canadian River is described as 50 to 75 feet thick (100 feet thick by the State Line) and 400 to 600 feet wide. The stream gradient and water-level gradient are presumably similar, at about 0.0001. Water levels are generally within a foot or two of the land surface and there are no noticeable water-level variations between piezometers open at different depths at the same sites. Throughflow in the alluvium is considered to be small, due to the low gradient. Stiff diagrams showing the quality of water in the channel deposits are included in TAB 14, as are charts showing ranges of chloride concentrations in water from wells in channel deposits.

A graph (Figure 36) plots surface water and ground water chloride levels between the Logan and Amarillo gages; see **TAB 14**. Based on this graph, salinity clearly drops by the State line, parallel with a drop in ground-water salinity and stratification. Salinity remains low to Tascosa, indicating no brine inflow. The report does not comment on the fact that there is a marked

increase in ground water chlorides at Amarillo, which is not paralleled by an increase in surface water chlorides.

HGC interprets stream survey data from the Logan area as indicating that there are no brine springs as such, but rather upflow through the alluvium ... probably in three locations (presumably miles 0.9, 3.5, 4.2). Construction at the railroad bridge may have disturbed channel sediments such that brine is forced to the surface in the area. At this time, our interpretations of the flow system differ somewhat from that given by HGC; see Chapter 3 of this notebook and Chapter 4 of the Surface Water Notebook.

The movement of salt down the Canadian River channel is a dynamic process, mostly occurring during high flows when channel sediments are flushed. At low flows that channel sediments store and retain the salt inflow because the rate of ground-water salt transport is extremely slow in comparison to transport by the river water. A water budget constructed for Lake Meredith shows that flow in the Canadian River is strongly affected by ground-water losses and bank storage.

A mixing cell model was constructed to predict the timing of salinity reduction on the salt load delivered to Lake Meredith. The model represented the river flow, alluvial ground water flow and the connection between the river and alluvial ground water along the channel from Ute Dam to Lake Meredith. The model simulated base flow conditions and accounted for the river discharge and salt load, tributary inflows and salt loads, alluvial ground water and brine concentrations, storage, flow rates and underflow. The transfer of water and salts from the alluvium to the river was simulated with an empirical transfer coefficient. The transfer coefficient was estimated through model calibration.

The model indicated that the salinity reduction resulting from 100% capture of the brine inflow would occur over a period of time. Salinity in the river above Lake Meredith was reduced by 24% over a ten year period. The maximum salinity reduction (70% of the existing salt load) occurred over a period of about 42 years. Simulations were also completed assuming a 50% brine capture. The results indicated a salinity reduction of about 12% over a ten year period. The model also was used to simulate the effect of pumping the alluvium. Pumping a rate equal to the brine inflow rate (0.9 cfs) produced a 16% reduction in the river salinity over a 10 year period.

The mixing cell model was considered to be conservative because it did not account for flushing of the alluvium during periods of high river flow. As discussed in the Surface Water Notebook, we do not regard the mixing model as being reliable.

HGC (1984b) summarizes the result of geophysical exploration for an injection site southeast of Logan. The first stage of the study consisted of examination and interpretation of electric logs from the National Oil

Company's Ute Anticline No. 1 test well (T12N, R32E, section 11). **TAB 10** shows these logs. They concluded from the study that sands in the Abo Formation and possibly in the underlying Sangre de Cristo Formation provide reasonable injection zones, with the best zones formed by sands in the middle of the Abo Formation. The total porosity of sands within the larger Abo-Sangre de Cristo section was measured at 10 to 17 percent, with shale contents ranging to 30 percent. Data from three additional wells were used to construct correlated sections in the area (**TAB 5**).

In the second stage of the study contractors completed and interpreted two deep seismic reflection profiles south of the river and east of Logan (TAB 4). Profile A was approximately 6 miles long and aligned east-west. Profile B was approximately 5 miles long and aligned north-south. Several large, possibly northwest-trending, faults with down-to-the southwest displacement were identified in the section. The Sangre de Cristo Formation north and east of the major faults was about 1500 feet thick. South and west of the fault it was about 4000 feet thick.

The most favorable location for exploration drilling was selected as an area in sections 29, 30 and 32, T13N, R34E (**TAB 4**). The most favorable target (middle Abo sands) are at a depth of 3800 to 4500 feet. They estimate that about 200 feet of sandstone will be required for injection of 450 gallons per minute. That section is available within the middle Abo but it may be necessary to complete the well over a larger section in order to obtain the largest possible capacity.

<u>USBR (1984)</u> summarizes and interprets information presented in HGC (1984a and 1984b) and presents additional data from agency field studies, related particularly to test drilling and collection of data from a piezometer network.

Drill hole DH-3 was drilled north of the Canadian River between Logan and Ute Dam. The well was intended to resolve the question over the stratigraphic position of the brine aquifer and was cored to total depth of 569.5 feet. A 147 foot section of grayish-white to bluish-gray sandstone was cored from about 350 feet to 497 feet and was correlated to the brine aquifer identified by earlier drilling. The top of the Permian section was found at 514 feet. The brine aquifer was separated from overlying sandstones by 54 feet of shale and mudstone and was underlain by about 21 feet of shale. TAB 10 includes the lithologic and natural gamma ray logs for DH-3.

DH-3 was completed as an observation well. Screen was set in the brine aquifer from 361 to 418 feet. The initial water level measured at DH-3 was about 90 feet below the land surface (i.e. above the top of the aquifer, indicating artesian conditions). Subsequent pumping apparently developed the well and the water level dropped about 5 feet (TAB 11). Water from DH-3 was sampled, though not without problems (see discussion on page IV-56). The

water is slightly less saline than, but otherwise chemically similar to, the brine sampled at wells on the canyon floor ($\mathsf{TAB}\ 13$).

USBR's experience with DH-3 led them to conclude that the lithologic log from DH-2 was probably incorrect because of caving during drilling. DH-2 was also found to be plugged at 160 feet and to be hydraulically connected to the river, that is, water level changes corresponded to fluctuations in river flow. It was not possible to know for sure where a water sample from DH-2 came from, but it was considered to represent a mixture of Triassic water and Permian brine.

Brine from well OW-3 was analysed for tritium activity (**TAB 13**). No activity was found and USBR concluded that the brine aquifer at OW-3 does not contain modern water from Ute Lake. Samples for OW-3 and OW-4 confirmed water typical of a sodium-chloride brine (e.g. see Stiff diagrams in **TAB 14**).

The USBR report provides a brief interpretative summary of hydrogeologic conditions near Logan. A natural sodium chloride brine is formed in the salt-bearing Permian section. The brine moves upward (probably through fractures) to the overlying Triassic section at a rate of 0.6 cfs. The water is mixed with relatively fresh water in the Triassic section. Water from the Triassic section discharges (also probably through fractures) to alluvial deposits along the Canadian River at a rate of 0.9 cfs. The brine is further mixed with fresher water as it moves upward and downstream through the alluvium. It discharges to surface water at numerous points along the river, particularly within the first 10 miles downstream from Ute Dam.

Limited subsurface data did not allow conclusions about the flow system within the Permian, Triassic and alluvial sections. Comparison of records from Ute Reservoir and from a recorder at well TW-1 did not indicate that the brine aquifer was hydraulically connected to the lake; rather, small water level fluctuations reflected atmospheric pressure changes and earth tides.

USBR established 5 stations on the Canadian River and Revuelto Creek where the chemistry of the river water and alluvial ground water were monitored over a period of 17 months from May, 1983 to September, 1984. Ground water levels were measured in August of 1983. Each site included a nest of two or three piezometers completed at depths from 15 to 55 feet. The wells are all included in the inventory presented in TAB 9. Figures showing the locations and construction of the piezometers are included in TAB 10; these figures include water-level information. Water quality data, including a statistical analysis, are in TAB 13; Stiff diagrams and scattergrams of the data are in TAB 14.

Site 1 (at mile 1.6 below Ute Dam, i.e. upstream from the Highway 54 bridge near Logan) included two piezometers completed to 16 and 22 feet, each with 4 feet of screen at the bottom of the hole. USBR questioned whether the deeper piezometer reached the bottom of the alluvium. Ground water quality

showed some variation during a spill of Ute Reservoir in June and July, 1983. Composition otherwise showed little variation. The shallow piezometer averaged 15,585 mg/l TDS while the deeper one averaged 13,670 mg/l; as with the other sites, the water was dominated by sodium and chloride, with relatively low sulfate concentrations. At the time the piezometers were completed, there was a strong upward gradient in the alluvium, with the head in the deep piezometer being slightly higher than the river elevation.

Site 2 (2.2 miles below Ute, at the Highway 54 bridge) included three piezometers completed to 22, 40 and 55 feet, each with 4 feet of screen at the bottom of the hole. The alluvium/bedrock contact was at 59 feet. Steel bits were left in two of the holes. As at Site 1, ground water quality varied with river flow. The piezometers show very little vertical variation in water quality, with total dissolved solids averaging slightly over 15,000 mg/l. At the time the piezometers were completed, there was essentially no vertical gradient in head, but water levels were slightly above the river elevation.

Site 3 (5.4 miles below Ute, upstream from the Revuelto Creek confluence) included two piezometers completed to 20 and 34 feet, each with 4 feet of screen at the bottom of the hole. Although the deeper piezometer was completed near bedrock, it was not located at the point of thickest alluvium. The water at the shallow level is only slightly more saline than water in the river at the site; the total dissolved solids averaged 13,229 mg/l. In contrast, water in the deep piezometer was much more saline, averaging 24,846 mg/l. In fact, the deep piezometer at Site 3 shows the most impact of brine of any of the alluvium data points.

Site 4 was on Revuelto Creek, about 0.2 miles above its confluence with the Canadian River. The site included two piezometers completed to 15 and 20.5 feet, each with 4 feet of screen at the bottom of the hole. The deeper piezometer penetrated bedrock. The ground water salinity was much lower than observed at other sites and showed greater variation over time. The shallow piezometer averaged 3688 mg/l TDS while the deeper one averaged 5168 mg/l; unlike other sites, the chloride: sulfate ratio was relatively low, in the range 2:1 to 3:1. All constituent concentrations were relatively high at the outset and fell over a period of three months. They remained low through most of program, but peaked again briefly in June, 1984. At the outset of the study concentrations generally were higher in the deep piezometer than in the The difference disappeared as concentrations fell and shallow piezometer. there is little difference between the piezometers over most of the sampling period. At the time the piezometers were completed, there was no significant head gradient.

Site 5 was not completed.

Site 6 (9.9 miles below Ute, on the Canadian below the Revuelto Creek confluence) includes three piezometers drilled to 21, 31 and 50 feet, each with 4 feet of screen at the bottom of the hole. The alluvium/bedrock contact

was at 52 feet. Drill bits were left in the holes. Concentrations in the river are lower than at upstream sites because of dilution from Revuelto Creek. Concentrations in water from the shallow piezometer averaged 8,816 mg/l total dissolved solids, while the intermediate piezometer averaged 13,651 mg/l and the deep piezometer averaged 20,319 mg/l. No significant hydraulic gradient was observed at the site.

Salinity in the river increases from Sites 1 and 2 (which are essentially similar) to Site 3, then decreases to Site 6 because of dilution. Concentrations in shallow piezometers decrease steadily from Site 1 through Site 6. Concentrations in deep piezometers increase from Site 1 through Site 3 then decrease to Site 6. There is a good correlation between chlorides, TDS, and field-specific conductance for the piezometer and surface water data.

At p. IV-51, USBR summarized the data from the piezometer network thusly: "the data points and sampling frequency have not been great enough to answer the questions about the distribution and movement of brine in the alluvium." Among the specific possible problems mentioned in use of the data are: Problems may include: inadequate depth of piezometers; incomplete mixing of waters, freshwater springs, brine pools etc. USBR suggested installing continuously recording conductance meters in a shallow piezometer and adjacent stream reach (e.g. below Revuelto Creek) in order to observe fluctuations in water quality with flow.

USBR also presents information on water quality in the Triassic geologic units and the Permian Brines; for example, see stiffs diagram in **TAB 14**. Relatively high sulfate levels are reported for Triassic units near the mouth of Revuelto Creek; one speculation is that the brine source in this area is gypsum rather than halite. The brine aquifer samples were discussed at the beginning of the discussion of USBR (1984).

To evaluate hydrologic and geologic systems in the Logan area, USBR developed a quasi-three dimension (two-layer) ground water model. The model also was used to test the feasibility of controlling brine seepage to the alluvium. Details of the USBR model in the report were limited, so notes on the model construction and results were obtained by Lee Wilson and Associates from USBR; the discussion below incorporates those notes, as well as the published text. Our critique of the model is provided in Chapter 3.

The model simulated flow in a 4.5 mile (north-south direction) by 5 mile (east-west direction) area near Logan. The area included the Canadian River from Ute Dam to about 9 miles downstream from the dam and the lower 5 miles of Revuelto Creek. The finite-element mesh contained 249 nodes and 314 elements in each layer and was designed to provide a relatively high degree of resolution along the Canadian River.

The system was simulated in two layers. The top layer represented the Logan sandstone. The lower layer represented the brine aquifer. The layers

were connected vertically across a 50-foot confining bed. The top layer was considered to be unconfined, but spatial variations in the thickness of the layer were specified by the modeller, rather than computed by the model. The thickness varied from 150 feet to 600 feet. The lower layer was 100 feet thick and was everywhere confined.

The top layer was simulated with a hydraulic conductivity of 0.5 feet per day. Under transient conditions, the top layer was assigned a specific yield of 10% and a specific storage coefficient of 3.0×10^{-5} per foot. The brine aquifer was simulated with a hydraulic conductivity of 24 feet per day. Under transient conditions, the brine aquifer was assigned a specific storage coefficient of 1.4×10^{-6} per foot. The intervening confining bed was assigned a vertical hydraulic conductivity of 0.0011 feet per day.

The modeler specified constant head boundaries in both layers at nodes along the northern and southern borders of the area. All heads in the lower layer (brine aquifer) were equivalent fresh water heads; water density was not a model variable. The Canadian River throughout the model area was simulated as a leaky boundary. No special conditions were used to simulate Revuelto Creek.

In a steady-state simulation the head in the top model layer conformed to topographic changes. Water movement was toward the Canadian River. The potentiometric surface in roughly the western 2/3 of the model area sloped to the east with some tendency to converge toward the river, particularly just downstream of Ute Dam. In the eastern 1/3 of the model, the potentiometric surface sloped toward the river. Flow between the two layers was everywhere upward.

Constant head boundary flows to the brine aquifer supplied the vast majority of inflow to the model area. Discharge to the Canadian River provided most of the discharge. The mass balance for the steady state simulation is summarized below.

Net flow from constant heads to brine aquifer	7.74 cfs
Upward flow to Logan sandstone	7.74 cfs
Net flow from constant heads to Logan sandstone	0.17 cfs
Ground water discharge to the Canadian River	7.92 cfs

These values fail to sum to zero because of roundoff during conversion from cubic feet/day to cubic feet/second.

Of the total upflow from the brine aquifer to the Logan sandstone, only 0.89 cfs occurs directly beneath the river. The remaining 6.85 cfs of upward flow reaches the river by lateral flow through the Logan sandstone.

The model was run under transient conditions to simulate a salinity control program consisting of a single well at the site of TW-1 pumping 450

gpm for periods up to 10 years. USBR reported that drawdown at the pumping well stabilized within one month at 23 feet; the drawdown was negligible about 1 mile to the west of the well and 2 miles to the northeast.

The modeller's principal conclusion was that the results were highly sensitive to variations in the input and that without further studies to constrain the physical geometry of the system and the aquifer coefficients the system could not be reliably simulated. The transient results also imply that more than one well will be required to adequately control saline seepage to the river and that the pumping rates will probably need to be higher than simulated in order to account for the induced downward flow of fresh water.

<u>USBR (1985)</u> is the main project report, and with respect to hydrology it mostly summarizes information from the 1984 studies discussed above. The following are among the important conclusions or observations made in the report.

- . "The hydrogeologic investigations ... determined that a sodium-chloride brine of natural origin produced by dissolution of Permian halite beds flows into the Canadian River near Logan, New Mexico. The brine flows upward from the Permian deposits into a geologic unit in the upper Permian or lower Triassic Formations then upward into the river alluvium. The exact route of movement is not known but is probably through a complex fracture system. The movement of this brine through the alluvial system is not very well understood. Brine appears to discharge into the river at several discrete points; but because of influences from freshwater springs and floodflows, these sites have not been adequately defined. It is possible that brine seepage may be relatively continuous downriver from Ute Dam." (Pages b-c.)
- "The ... analysis of the regional and site geology (New Mexico and Texas) relating to the sources of brine contamination in the Canadian River concludes that about 70 percent of the sodium chloride entering Lake Meredith comes from New Mexico and that most of this contamination enters the river channel near Logan, New Mexico. The report also states that an additional 10 to 15 percent of the total salt load enters the river channel between the Tascosa and Amarillo gauges. Reclamation investigations indicate that this brine appears to flow continuously to the river system. Floodflows do not appear to affect concentration levels within the alluvium." (Page c.)
- "Results of a ... model ... provide estimates of the effect after 10 years of 100-percent reduction in brine inflow near Logan. The reduction is calculated to be about 24 percent (in milligrams per liter) of total dissolved solids (TDS) in the river water reaching Lake Meredith. If the brine inflow was only reduced by 50 percent, the time for the system to respond was nearly the same; but the amount of salinity reduction was about half of that calculated for the 100-percent reduction

in brine inflow. The response to the inflow salinity reduction in Lake Meredith would be direct but at a slightly reduced rate." (Pages d-e.)

- . "Based on existing information on deep formations in the Logan, New Mexico, area, a suitable disposal zone probably exists for deep-well injection." (Page e.)
- . "Additional fieldwork to include exploratory drilling and long-term pump testing is needed to verify the findings presented in this report and the effectiveness of the plan." (Page e, emphasis added. Elsewhere, the need for geophysical work also is noted.)
- The hydraulic conductivity of the alluvial aquifer is estimated at about 30 feet per day and the velocity of ground water is roughly 0.1 foot/day. (Page II-5.)
- . The data to support either large storage of salt in the channel alluvium or the flushing of salts from the alluvium by high flows is not available (page IV-3).
- . The most important result of the computer model of the brine aquifer is to demonstrate how little is really known about the hydrogeology of the area. (Page IV-9.)
- . 35 locations of possible seepage have been identified by EPA based on remote sensing (page V-36; details are not provided).

The report includes a 17-step program to implement the salinity control project, to include geophysical studies, test drilling, installation of additional piezometers, aquifer testing and sampling, and modeling. Report materials dealing with project alternatives are outside the scope of this notebook.

2.4 Other reports

Spiegel (1957) reports a preliminary evaluation of the geology and hydrology of the Dunes damsite (roughly, section 2, T13N, R35E) on the Canadian River, about 84 linear miles upstream of the state line. The available copy contains two geologic maps covering exposures along the river from near Logan to the eastern extent of the dam site area, a descriptive table of local stratigraphic units and two cross sections. TAB 3 includes the table of local stratigraphic units.

The report describes a large collapse feature, possibly centered in section 34, T14N, R35E, north of the river. The feature is not readily evident from the topographic map.

Geologic maps show that the Canadian River from Logan to the Dunes damsite flows alternately over the Trujillo and Tecovas Formations. The Tecovas Formation is exposed in the lower canyon walls from about 8.5 miles to 9.5 miles downstream from Ute Dam and again from about 24 to 28 miles downstream from Ute Dam. The second reach underlies the Dunes damsite.

 \underline{SEO} (1961) is a preliminary report on the geology of the Ute damsite. It provides an overview of area geology as well as detailed geologic maps and cross-sections of the damsite area, including the east end of the reservoir area and a small reach downstream from the damsite (TAB 4 and TAB 5).

The principal regional structure is a southwest-plunging anticline with its axis passing just west of the Olean No. 1 Woods well (IAB 4). Strata west of the axis dip gently westward into an irregularly shaped structural depression locally interrupted by several east-trending anticlines. Strata south of the axis dip south-southeast into the Palo Duro basin. Structural relief on the top of the San Andres Formation in this fold is as much as 1,427 feet.

Alluvial deposits along the Canadian River consist of unconsolidated silts, sands and gravels and were found to be 50 to 60 feet thick along the stream bed from the dam site to the US 54 bridge south of Logan. Older alluvial deposits are preserved on terraces above the river; the oldest locally identified terrace is widely developed north of the river but absent south of the river.

Bedrock exposed at the Ute damsite consists entirely of the Chinle and Trujillo Formations. The Chinle consists of interbedded red shale and light tan to red sandstone. The Trujillo is predominantly light gray to tan sandstone. The sandstone is massive, thick bedded and generally well-cemented.

Detailed cross-sections show that bedding dips generally eastward at a low angle. Several randomly oriented elongated domes and troughs modify the regional dip near the south abutment of the existing dam. Relief on the local folds is at least 70 feet. A high angle fault with dip slip of at least 190 feet (down to the north) was tentatively identified in borings north of the river at the existing dam site. The fault trace is buried under terrace deposits. It trends east northeast-to east through the immediate area of the dam, but its extent in either direction is unknown.

3. SYNTHESIS OF INFORMATION AVAILABLE ON THE BRINE AQUIFER

In our judgment, the project reports (and other references) discussed in Section 2 provide a limited understanding of the geologic and hydrologic conditions which fundamentally explain the brine inflow to the Canadian River which occurs in the Logan area. Consequently, it is difficult to use the available information and design with confidence a system to intercept and dispose of the brine. In this report, we have made an effort to interpret the literature in order to improve the understanding of the brine flow system. We believe this understanding is sufficient to direct more detailed studies of the system, including analyses of existing data and development of new data through field investigations.

3.1 Hydrogeologic setting

The starting point for understanding the brine problem is the regional hydrogeology. We need to know just where the brine comes from, and what kinds of rocks and structures it flows to en route to the discharge area near Logan.

<u>Background information</u>. **TAB 3** provides stratigraphic sections and **TAB 4** provides numerous maps which illustrate the regional geology and geomorphology of the area extending from the Rocky Mountains and central highlands of New Mexico eastward across the Texas Panhandle. Geologically, much of what is important today dates from the Permian when sediments which include salt deposits accumulated in a series of deep basins, such as the Palo Duro Basin; from the Triassic, when sediments (mostly redbeds) were deposited atop the Permian beds; and from relatively recent times, when the Canadian River cut its channel into the Triassic and Permian rocks and deposited modern alluvium.

In the Logan area, the primary aquifers are the Triassic sandstones, but since these units don't themselves contain a brine source, the salt water of the area must originate in the Permian evaporites. Consequently, a successful salinity control project near Logan requires an understanding of how salt flows originate in the Permian rocks, and how they reach the Canadian River through the Triassic units.

Salt dissolution from the evaporites. Rocks of middle to late Permian age are found at relatively shallow depths throughout the area from Logan to Lake Meredith, and are exposed at the land surface in many locations. The rocks are mostly red beds interlayered with varying amounts of limestone, dolomite and gypsum. At depth below the surface and distant from their surface exposures these rocks also contain thick bedded salt deposits. The salt beds originally extended through most of eastern New Mexico and the Texas Panhandle. Surface and ground water have dissolved much of the salt and the beds are now found mostly in the central parts of the major Permian basins. See TAB 5 for useful cross-sections, particularly Gustavson and Finley (1985, Figure 5).

Collins (1984) provides the most succinct description of the behavior of the evaporites with respect to ground water. Bassett and Bentley (1983) list the permeability of the undisturbed evaporite section at about 1×10^{-4} md (essentially impermeable). By comparison, the second most impermeable hydrostratigraphic unit in the region is 10,000 times more permeable than the evaporites. This low permeability restricts ground water movement, so that flow within the evaporite section occurs mostly through fractures in disturbed areas near faults and along the margins of the existing salt beds. Senger, et al. (1987) estimated the permeability in these marginal zones to be 100,000 times higher than in the undisturbed evaporite section. Along the evaporate margins, the ground water dissolves and removes the salts. Overlying and adjacent beds collapse into the resulting voids and the disturbed zone gradually advances into the salt beds.

According to Gustavson, et al. (1980a), dissolution within a single salt bed occurs at a relatively sharp and steeply sloping front. The sequence contains many individual salt beds, each with its own dissolution front. Dissolution has been active longer or at a higher rate in shallow salt beds and the dissolution fronts are more advanced than in lower beds. Salt dissolution in the Palo Duro Basin in Texas has been carefully studied. There the dissolution fronts in the youngest and oldest affected salt beds are separated in places by more than a hundred miles, but most of the dissolution fronts are found in relatively narrow bands along the Canadian River and the eastern escarpment of the High Plains (see maps in TAB 4, Part B.3).

The band of dissolution fronts along the Canadian River extends west to near Logan (see maps in TAB 4, Part B.3). Salt beds are present in the youngest part of the section (Salado and Seven Rivers of the Artesia Group) several miles south of the river and are absent north of the river. Salt beds occur in the middle of the evaporite section (San Andres Formation) under the river near Logan but are absent just north of the river and to the west. Salt in the Glorieta Formation appears to be only slightly more extensive than in the San Andres. Salt beds are still present in the oldest part of the evaporite section (the Yeso Formation) under the river and for several miles to the north and west.

As noted, removal of the salt beds causes collapse of the overlying units and subsidence at the surface. Collapse forms breccia blankets in the dissolved zone and fractures, faults and folds in the overlying rocks. Breccia pipes may form through overlying rocks where the collapse is sufficiently localized. Sink holes and closed drainage basins sometimes develop as surface expressions. The structures and sediment-filled sink holes and basins can be found wherever dissolution has occurred in the past, but the direct surface expressions remain only where dissolution is active or very recent.

Based on Gustavson, et al. (1980a) and Collins (1984), among others, the structural effects of collapse are found near Lake Meredith and upstream of the lake along the Canadian River in Potter County. Tertiary and younger

sediments north of the river obscure structural evidence for collapse. There, an offset on the High Plains surface and variations in the thickness of the Ogallala Formation are evidence for several hundred feet of relatively local collapse during Ogallala deposition and regional subsidence of 250 feet since Ogallala deposition.

Collapse structures were identified along the Canadian River east of Logan, but the absence of surface expressions suggests that dissolution and collapse there may not be active. Collapse features also are found in the Tucumcari area (Trauger and Bushman, 1964) and there the dissolution and subsidence may still be active.

Based on our inspection of topographic maps, several large depressions occur south of the Canadian River along a east-northeast oriented line from Tucumcari to at least section 30, T13N, R35E (i.e. only a few miles southeast of Logan). These depressions have not been investigated in the field. One, in section 5, T12N, R34E, shows clearly on the New Mexico State Highway Department geologic map of the Logan Quadrangle. This depression is shown in TAB 4, Part A. The flat floor of the depression is more than 20 feet below surrounding areas, roughly elliptical and covered with young alluvium. The ellipse is about 2/3 of a mile across on its northeast trending major axis. Bedrock is exposed on the sides of the depression and the surrounding area is covered with pediment deposits. The depression contains seven small seasonal ponds, all located on the lower side slopes and margins of the depression floor. This depression and other similar depressions may be evidence for active dissolution and subsidence.

Reflecting regional hydrogeologic gradients, ground water probably flows into the Permian section near Logan from the east and north. It moves through areas where salt dissolution and the structural disturbance associated with collapse have created open vugs and fractures. Upward flow from the pre-evaporite section also is possible but seems less likely since bedded salts (virtually impermeable) are locally still present near the base of the evaporite section. Ground water approaching a dissolution front must flow either within the evaporite section parallel to a dissolution front or upward to the overlying Triassic section.

Without doubt, dissolution of the Permian evaporites can account for the salts which are added to the Canadian River near Logan. The total chloride load, which is roughly 0.7 kilograms/second (see Surface Water Notebook) corresponds to a halite dissolution rate of only 593,000 cubic feet per year. If dissolution occurs only in a 350-foot thick salt section within the Artesia Group and upper San Andres Formation and only along a 21-mile front from Tucumcari to Logan, then this amount of salt could be supplied for 100 years, and the evaporites would dissolve back only 1.5 feet.

The high plains surface north of the Canadian River from Ute Dam to Sanford Dam is about 80 meters lower than it is south of the breaks. Studies in Texas explain this as the result of salt dissolution north of the river, and the same explanation is plausible for New Mexico. Salt is still present

along and north of the river in the lower San Andres and Glorieta Formations and in the Clear Fork Group. South of the river, salt is present in stratigraphic units as high as the middle of the Artesia Group (Seven Rivers Formation).

Currently, salt dissolution is most active along the eastern caprock escarpment and adjacent parts of the Rolling Plains. Dissolution also is occuring beneath the Canadian River, but at a lower rate.

Brine movement through the Triassic Section. The evaporite section is overlain by roughly 2,000 feet of Triassic age sedimentary rocks. In New Mexico the section is commonly divided into the Santa Rosa Sandstone and the Chinle Formation. See ${\sf TAB}$ 3.

The Santa Rosa Sandstone occupies the lower part of the Triassic section and is exposed along the Canadian River from Ute Dam to the State Line. It is generally covered north of the river by Tertiary and younger sediments. South of the river the Santa Rosa is exposed or near the surface in a band parallel to the river and along its tributaries. The outcrop band is about a mile across near Logan and expands to about 5 miles across near the State Line. South and east of the outcrop band the Santa Rosa is beneath the Chinle Formation and younger units. It is not exposed elsewhere in Quay County or adjacent areas in New Mexico. At Logan the Santa Rosa can be divided into two units and correlated to formations in Texas.

The lower Santa Rosa has been correlated to the Tecovas Formation in Texas, and we will use that term in this report, because the lower Santa Rosa appears to be the "brine aquifer". At DH-3, the Tecovas consists of about 220 feet of friable to well-cemented, light colored sandstones and interlayered clays. Based on geologic maps, the Tecovas at Logan is primarily sandstone but it becomes generally more fine grained to the east. In Texas the Tecovas is primarily shale. A prominent shale about 50 feet thick lies at the top of the Tecovas near Logan. The shale is reportedly absent to the east at the Dunes damsite and may be missing elsewhere. Exposures of the Tecovas Formation are limited to the immediate vicinity of the Canadian River. From Logan to near Rana Canyon the Tecovas is intermittently exposed along the lower canyon walls. Below Rana Canyon the exposure is more extensive.

The upper Santa Rosa has been correlated to the lowest sandstone unit of the Trujillo Formation in Texas. We will use the term Trujillo here, to distinguish this part of the Santa Rosa from the Tecovas brine aquifer. At DH-3, the Trujillo may be 300 feet thick and consists of well-cemented sandstone and conglomerate with some interlayered clay and coal horizons. The Trujillo Formation forms the prominent bluffs along the Canadian River canyon in New Mexico and underlies the river channel near Ute Dam and in some intervals from Ute Dam to near Rana Canyon.

The upper 1,500 feet (roughly) of the Triassic section contains the Chinle Formation. The Chinle consists of red to brown, sometimes variegated shale with interlayered siltstones. Sandstone similar to the Trujillo occurs in the

middle of the section. The Chinle is not exposed along the Canadian River below Ute Dam.

In the published literature, the Santa Rosa Sandstone is generally regarded as an aquifer. It lies between the underlying brine source in the evaporite section and the overlying confining beds of the Chinle Formation. It is the only bedrock unit directly connected with the Canadian River in the Logan area. Because of this geometry, the Santa Rosa is the major ground water outflow area for brine in the Logan area.

The Texas Bureau of Economic Geology has prepared two cross-sections to illustrate the regional geology: see TAB 5. The cross-sections are consistent with our understanding of the geology in the area, although they are not detailed enough to show the brine aquifer and they do not reconstruct probable flow lines for the brine.

Our interpretation is that ground water would enter the Santa Rosa at depth by upward flow from older rocks; along dissolution fronts in the underlying evaporites that upflow would be highly saline. The Santa Rosa near its outcrop probably receives downward flow of relatively fresh water from the overlying Chinle. Where exposed, the Santa Rosa would be recharged by percolating precipitation and runoff.

Ground water within the Santa Rosa at depth probably is influenced by both the prevailing eastward gradient in the underlying units and the convergent flow towards the Canadian River in overlying units. Thus flow at depth would probably vary from eastward to northeastward south of the river and southeastward north of the river. Where the Santa Rosa is exposed along the Canadian River the flow must converge toward the river.

Flow through the alluvium. Recent and active channel deposits of interlayered, unconsolidated silt, sand and gravel occur along the Canadian River. The alluvium is mostly silt and fine sand at the surface, with gravels and coarse sand found at depth. The alluvium is up to about 60 feet thick near Logan and the thickness increases downstream to about 100 feet near the State Line.

White deposits - probably left after the evaporation of water from the capillary fringe - are common in alluvium along the river banks 2 to 4 feet above the river level. The deposits are extensive in some areas where the valley floor is broad and flat and elevated only two or three feet above the river.

The alluvium receives ground water inflow from the Santa Rosa Sandstone and in some areas along the north side of the river from Tertiary and younger units. The alluvium also receives some recharge from precipitation.

Ground water in the alluvium flows, on net, to the river. The average gain under baseflow conditions from below Ute Dam to the state line is approximately 4 cfs. In detail the flows are variable, both spatially and

over time. The river gains from the alluvium when and where recharge and ground water flows to the alluvium exceed evapotranspiration from the alluvium. Otherwise the river loses to the alluvium.

3.2 A closer look at the brine aguifer

The USBR salinity control project for Lake Meredith relies on the capture of brine from the "brine aquifer". To evaluate the potential success of the project, it is important to state clearly what we do and don't know about the aquifer.

Stratigraphy and lithology. The nature of the brine aquifer is known from USBR drilling records. The brine aquifer identified at DH-1 and cored at DH-3 is positioned stratigraphically below the distinctive sandstone of the Trujillo Formation, above the salmon red shale in the Permian section and is lithologically similar to sandstones in the Tecovas Formation as described by Spiegel (1972a). The brine aquifer at DH-3 consisted of approximately 150 feet of sandstone with relatively minor interbedded claystone. The sandstone is overlain by 54 feet of shale and mudstone and underlain by 18 feet of shale. The overlying and underlying shales and mudstones probably should be included with the sandstone in the Tecovas Formation.

Brine composition Water in the brine aquifer is highly saline and dominated by chloride and sodium. The specific conductance of brine from OW-3 is usually within the range of 50,000 to 70,000 micromhos; at a conductance of 60,000 micromhos the brine would contain about 1.3 times the salinity of normal sea water (i.e. roughly 45,000 mg/l total dissolved solids). Sea water at this salinity would have a density of 1.035 grams/cm³ and a relative viscosity (compared to pure water) of 1.094. Water from DH-3 is considerably less saline. Its specific conductance of 36,000 micromhos is about 78% of normal sea water.

Chloride in brine from OW-3 averages more than 25,000 mg/l. The chloride concentration at OW-4, 19,700 mg/l, is slightly lower. At DH-3 chloride averages about 15,600 mg/l. Sulfate is much less prevalent than chloride; 2,790 mg/l at OW-3, 2,660 mg/l at OW-4 and 2,175 mg/l at DH-3. Sulfate/chloride ratios vary somewhat with the salinity; 0.139 at DH-3 (lowest salinity), 0.135 at OW-4 and 0.112 at OW-3 (highest salinity).

Variation in the salinity and the sulfate/chloride ratio probably reflects varying degrees of mixing between a chloride-dominated brine originating from halite dissolution in the evaporite section and relatively fresh water from the Triassic section with a higher sulfate/chloride ratio.

The sulfate concentrations listed above are substantially below the 9,000 mg/l concentration used in USBR (1979) to calculate the sulfate load originating from the brine aquifer. Using a concentration of 2,500 mg/l sulfate and a brine flow of roughly 0.9 cfs (see surface water notebook), the sulfate load would be less than 2,300 tons per year - less than half the 5,400

tons per year load calculated by USBR. The relative importance of the brine aquifer to the sulfate load entering Lake Meredith is correspondingly smaller; less than 9% of the total inflow of 26,910 tons per year rather than the 20% reported by USBR (and note that this ignores losses of salts which occur between Logan and Lake Meredith).

Permeability. The transmissivity of the brine aquifer was tested by USBR at 2,500 square feet per day at TW-1 and the hydraulic conductivity was calculated at 36 feet per day. Brine aquifer permeability was estimated for the USBR flow model at 25 feet per day. These values are extremely high compared to Bassett and Bentley's estimated permeability of 0.27-2.7 feet per day for the Triassic aquifers in Texas. A permeability of 25 to 36 feet per day is possible but it is roughly an order of magnitude higher than might be reasonably expected from a consolidated or semi-consolidated fine-grained sand; rather, it is typical of a good alluvial aquifer.

Permeability of the Trujillo Formation was estimated for the USBR flow model at 0.5 feet per day and a value near that might be considered more likely for the brine aquifer. However, this permeability value could not be reconciled with the transmissivity estimated from the pump test. What does this mean? We judge that either the pump test is flawed or the brine aquifer has exceptional permeability.

Storage. The storage coefficient of the brine aquifer in the TW-1 test was 1.5×10^{-4} . The specific yield of the Trujillo Formation was estimated for the USBR flow model at 10%. Both values are within the reasonable range of expectations, though the specific yield value might be higher than anticipated.

<u>Leakance</u>. Leakance, which is vertical conductivity divided by thickness of a confining layer, is a parameter frequently used in models. It exercises a controlling influence on vertical flows and influences the drawdown in wells, and must be reasonably well known to model (for example) the impacts of a brine control project. The vertical conductivity of the clay above the brine aquifer was estimated by calibration of the USBR flow model at 0.0011 feet per day. Given this value and the 50-foot thickness of the confining bed, leakance across the bed would be 2.2x10⁻⁵ per day.

It is important to recognize that the stream survey conducted by CRMWA in 1992 was the first to provide both water quality and flow data. Based on these data, it should be relatively easy to perform calculations which balance observed changes in the stream system with probable inputs from various ground-water sources, thus providing for calculation of variations in brine inflow rates in different reaches of the river. In turn, inflow variations could be interpreted in terms of variations in leakance, and variations in leakance could be interpreted in terms of geologic controls of brine flow. In our judgment, quantification of these factors is essential to evaluation of brine sources in the Logan area, and to evaluation of control strategies.

Structure. Bedding dips generally westward from the crest of an anticline that crosses the river in or just east of section 6, T13N R34E (i.e. near the confluence with Revuelto Creek) where, according to Spiegel's maps (Spiegel, 1957), the Tecovas is exposed along the river. The dip is interrupted by several small flexures. Work at the Ute Dam site (SEO, 1961) provided information on structural relief in that area but elsewhere there is little information to indicate the size or relative importance of the many small flexures that have been mapped.

Spiegel mapped the top of the Tecovas at the river level about 18,000 feet northeast of DH-1 at an elevation of 3650 feet. DH-1 encountered shale above the brine aquifer at an elevation of about 3485 feet. The apparent dip of the Tecovas from its exposure along the river to DH-1 is southwest at 48.4 feet per mile (0.5 degrees). The alluvium would lie directly on the Tecovas from its exposure southwest to near the Revuelto Creek confluence. The Tecovas would be well below the alluvium upstream from Revuelto Creek. However, small flexures between the Tecovas exposure and DH-1 could allow for considerable variation in the dip such that, for example, the brine aquifer could be at the base of the alluvium at points upstream of Revuelto Creek.

Small flexures probably also influence the position of the brine aquifer between Ute Dam and DH-1. The top of the Tecovas Formation at DH-3 was found 296 feet below the land surface at an elevation of 3485 feet; there is no apparent dip between DH-1 and DH-3. It is possible that the two wells lie directly on strike but that seems unlikely because it would require that beds dip almost directly to the north or south. It seems more likely that a synclinal axis passes between DH-1 and DH-3. The latter possibility is consistent with the presence of anticlines west of DH-3 at Ute Dam.

Faults have not been observed in exposures along the Canadian River but SEO (1961) did map a fault north of the river near Ute Dam. That fault juxtaposes the Trujillo and the Chinle sections. The fault would also offset the brine aquifer and might act as a boundary to ground water flow. The lateral extent and orientation of the fault are not clear from Spiegel's work. If the fault is extensive it could effectively bound the northern extent of the brine aquifer.

3.3 Characterization of the brine flow system

<u>USBR</u> ground water model. To the extent that the literature may characterize brine flow, that characterization is inherent in the USBR model, which represents a conservative interpretation of the hard data available to the modeller. However, some known or reasonably approximated features - such as the dip of bedding and the existence of fresh water recharge - were not used in the model.

Moreover, the model as constructed simulates unrealistic flow conditions. For the model, virtually all ground water originates in the brine aquifer.

The model shows the brine moving to the river without any mixing with fresh water, even though actual hydrologic conditions make it clear that fresh water in the Trujillo sandstone is a major component of the real-world flow system.

USBR concluded that the model results were undependable and that available information did not define even the fundamental geometry of the system. Additional information is needed before a model can be confidently constructed and applied. But the problems experienced by USBR could have been reduced by incorporating at least some of the omitted details and employing a more accurate construction.

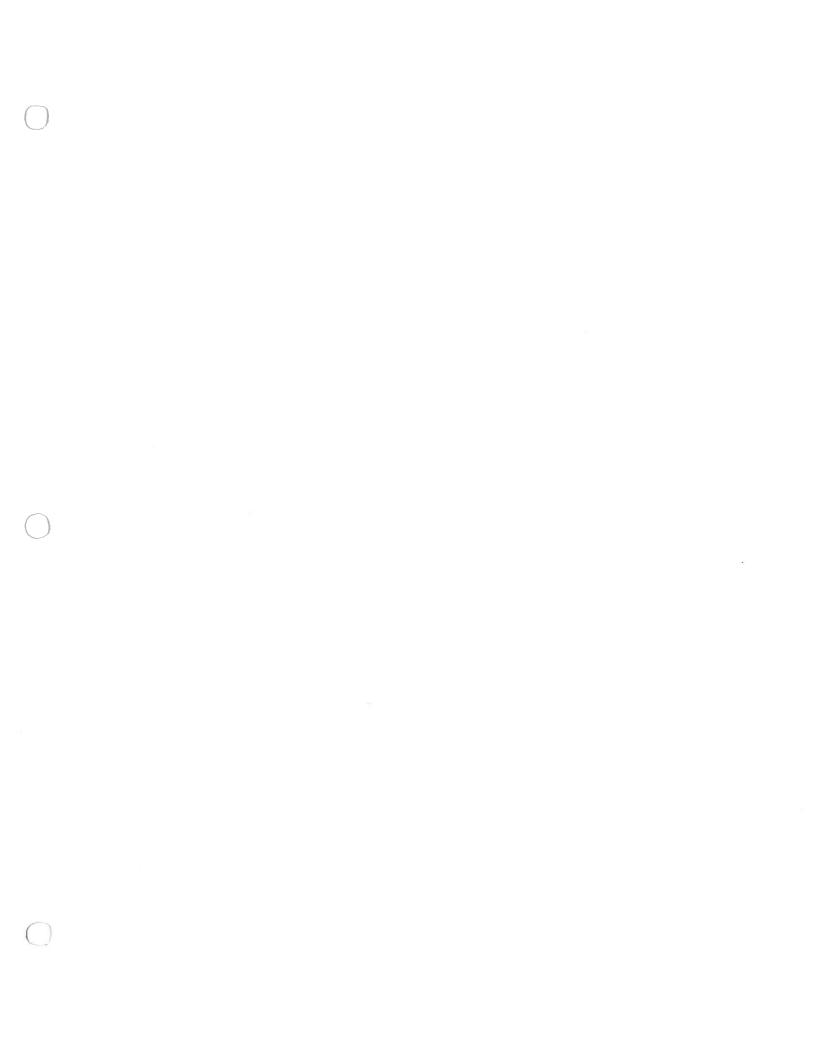
Synthesis of available information. The literature provides little insight to the brine flow system in the Logan area, except that the brines clearly originate in Permian units older than the brine aquifer, pass through the aquifer in the Logan area, and pass through other geologic units (e.g. alluvium) before reaching the river. It also is evident that river inflows include fresh water which must be passing through rock units above the brine aquifer. This fresh water is probably of relatively local origin (recharge) as the Santa Rosa Sandstone (Trujillo and Tecovas) at depths to the west and southwest is reported to contain only highly saline water.

Additional studies have been proposed to CRMWA which should greatly improve the understanding of the flow system. These studies would quantify at least two important components of the system: the response of water levels in the brine aquifer to storage changes in Ute Reservoir (TAB 11) and the hydraulic properties (leakance) of the confining layer which in some locations separates the brine aquifer from the Canadian River. Until such studies are completed, we have used the cross-section at the end of TAB 8 to summarize our understanding of conditions in the Logan area. Key points shown on the section include:

- geologic evidence (well logs, and mapping done by the New Mexico State Engineer) shows that the Tecovas brine aquifer is overlain by a confining layer in much of the Logan area, but that this layer is missing (and the brine aquifer is near the land surface) in the gravel pit reach;
- artesian pressures in the Tecovas force some brine upwards to mix with fresh water and discharge to the Canadian River in the east-west reach which is upstream of the gravel pit reach;
- however the primary flow in the brine aquifer is eastwards toward a discharge zone at or near the gravel pit reach;
- based on geophysical evidence, it is unlikely that much brine flow persists past the gravel pit reach.

The above summary will be greatly expanded in the next edition of the notebook, after the additional work recommended to CRMWA has been completed.





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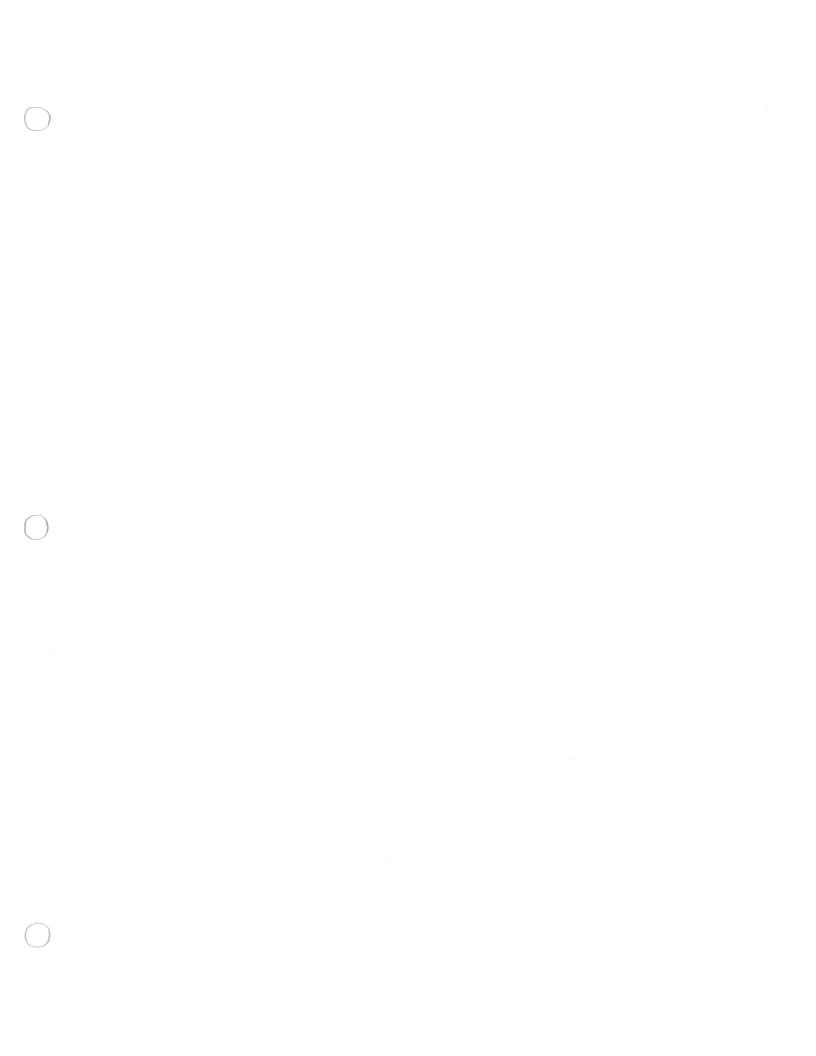
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Index to TAB 3: Stratigraphic columns of the study area

Collins, 1990, Table 1

HGC, 1984a, Figure 4

McGookey, et al., Table 1

Spiegel, 1957, Table 1

USBR, 1979, Figure 1

Table 1. General stratigraphic column and lithologies of the Texas Panhandle Palo Duro Basin (from Dutton and others, 1988). Alibates, upper Clear Fork, Tubb, lower Clear Fork, and Red Cave have been informally designated as formations.

System	Series	Group	Formation	Lithology and depositional setting	
	Holocene			Playa deposits	
Quaternary	Pleistocene Tule Blackwater Draw			Eolian and lacustrine clastics,	
			fresh-water limestone		
	Pliocene		Blanco	Fluvial, eolian, and lacustrine clastics	
Tertiary	Miocene		Ogaliala		
		Washita	Duck Creek	Marine sandstone, limestone,	
	Ī		Kiamichi	and shale	
Cretaceous	Comanchean	Fredericksburg	Edwards Walnut		
		Trinity	Paluxy	Transgressive sand and gravel	
Jurassic*	Upper		Morrison	Marine sandstone and shale	
Triassic	Upper (?)	Dockum	Trujillo Chinle Santa Rosa Tecovas	Fluvial-deltaic and lacustrine clastics	
			Dewey Lake		
	Ochoan		Alibates		
			Salado	Brine-pool salt, anhydrite, red beds. and	
			Tansill		
		Artesia	Yates		
	Guadalupian		Seven Rivers		
			Queen		
			Grayburg	peritidal carbonate	
		Pease River	San Andres (Blaine)	political of controls	
Permian			Glorieta		
	Leonardian	Clear Fork	upper Clear Fork		
			Tubb lower Clear Fork		
			Red Cave		
	Wolfcampian	Wichita-Albany	"Brown dolornite"	·	
	·	Cisco		Shelf and shelf-margin	
	Virgilian			carbonate, basinal shale,	
Pennsylvanian	Missourian	Canyon		and deltaic sandstone	
	Desmoinesian	Strawn	4	·	
	Atokan	Bend			
	Morrowan				
	Chesterian Meramecian	1		1	
Mississippian	Osagean	1		Shelf carbonate and chert	
	Kinderhookian'	1			
Devonian*	KINGGIRONAII				
Silurian*		Hunton		4	
CIICI PORT	Cincinnatian* Viola			Shelf carbonate and clastics	
Ordovician	Champlainian*	Simpson	1	Į	
	Canadian	Ellenburger			
Cambrian				Shallow marine sandstone	
				Igneous and metamorphic	

^{*} Unit absent or insignificant in the Palo Duro Basin

		NE NEW MEXICO	O STUDY AREA		PALO DURO BASIN	Z	
SYSTEM	SENIES	1		FUORB	FORMATION	General Lithology & Depositional setting	
	HOLOCENE				ollurium, dune sand Playe	·	•
QUATERNARY	PLEISTOCENE				Tahoka cover gands Tule / Plays Blance	Lacustrine electics & windblown deposits	
TERTIARY	NEOGENE		Ogallala	•		Firviol & locustrine clostics	
CRETACEOUS					undiferentiated	Merine shales &	*
TRIASSIC		DOCKUM		роским		Fluvial-delloic & locustring clastics	•
					Dowey Lake (Quartermaster)		••
3	OCHOA		Alibotes?		Alibotes	烧	
			*		Salado/Tonsill		
¥	15	×	•		roies		
	N.	ARTESIA	Bernal	ARTESIA	Seven Rivers	((4)) 9	
	GUADALUPE				Queen/Grayburg		est e
PERMIAN		2	San Andres		San Andres	sound sein, emperior, red beds, & perilidal dolomite	
			*		Glorieta *		
					Unger Clear Fort		
				CLEAR	7.46		a
	FONABO		Yeso	FORK	* 1		
	LEUNARD				Lower Clear Fork		
3					Red Cave		
				WICHITA			ε,
	WOLFCAMP		Abo	·			
	VIRGIL		Sangre de Cristos	CISCO		Shelf & shelf-merile	* 0
	MISSOURI	MAGDALENA	`	CANYON		carbonate, besinel shale, & dettale,	
PENNSYLVANIAN	DES MOINES		Madera	STRAWN		sandslene	
1903	ATOKA			BEND			
	MORROW						
	CHESTER						
MISSISSIPPIAN	MERAMEC					Shelf carbonate C	
e v	OSAGE						
ORDOVICIAN				ELLEN- BURGER		Shelf dolomite	
CAMBRIAN						Shallow merine (†)	
PRE	PRECAMBRIAN					Igneous & metamorphic	•1 - 15 X - 15 - 15 IS
*SALT PRESENT	ŀ		Gtratigraphic of	ຕວ່າກາກກ			
	e dure			7.	HYDRO GE	HYDRO GEO CHEM, INC.	

Eastern New Mexico Palo Duro Basin **Dalhart Basin** Anadarko Basin System Series Blackwater Draw Fm Blackwater Draw Fm Blackwater Draw Fm. Blackwater Draw Fm. Quaternary Ogallala Fm. Ogallala Fm Ogallala Fm. Ogaliaia Fm. **Tertiary** several formations, undifferentiated Cretaceous Exeter Ss. Jurassic Exeter Ss. Dockum Gp. Dockum Gp Dockum Gp. Triassic Dewey Lake Fm. Dewey Lake Fm. Dewey Lake Fm. Quartermaster Fm. Alibates Fm. Alibates Fm. Alibates Fm. Alibates Fm Salado Fm. Salado Fm. Salado Fm Tansili Fm. Cloud Chief Fm. Yates Fm. Ġ Artesia Gp. Artesia Gp. Artesio Seven Rivers Fm. Guadalupe Whitehorse Gp. Queen and Grayburg Fms. Dog Creek Blaine Fm. Blaine Fm. San Andres Fm. Son Andres Fm. **Permian** Flowerpot Fm Glorieta Sandstone Giorieta Fm. Glorieta Fm. Glorieta Sandstone upper Clear Fork Gp. 8 undifferentiated Clear Fork Clear Fork 2 Tubb Fm.2 Gp. lower Clear Fork Gp. undifferentiated Sangre de Cristo Abo Fm. Wichita Gp. Wellington Fm. Wichita Go.

Table 1. Stratigraphic chart, Permian to Quaternary strata, Palo Duro Basin and surrounding area.

(2) on salt flats that were isolated from open-marine water, and (3) on mud flats where salt crystals grew by displacing the mud matrix. Bein and Land's (1982) petrographic and geochemical study of the San Andres Formation in cores, on the other hand, indicates that evaporite rocks were deposited in a shelf basin or lagoon in which brine composition changed as CaCO₃, CaSO₄, and NaCl were successively precipitated. Carbonate strata in the

upper San Andres and Alibates Formations consist predominantly of dolomite.

Salt beds in the Palo Duro Basin range from a few feet to 200 ft (61 m) thick. Before dissolution of upper salt units in the northern Texas Panhandle and around the northern, eastern, and western margins of the Palo Duro Basin, the most widespread salt beds may have extended over the entire Palo Duro, Dalhart, and Anadarko Basins. Lower San Andres evaporite beds

¹Formation's lithology is not the same as the formally designated stratotype.

²The Tubb Sandstone member is informally designated Tubb Formation.

there and in Quay

Generally absent on the Bravo dome; no evidence elsewhere in this area. sandstone that weathers red is commonly present at the basc. Forms the Orange-rod (terra cotta) colored siltstone, fine-grained sandstone and green shale or friable white sandstone lenses; massive lens of friable Massive light tan sandstone and conglomerate with occasional maroon or canyon wall of the Canadian River from near the state line to about 3 Maroon, purple, and green shalv with bods of rod sandstone irregularly distributed vertically and horizontally (but generally with important sandstone beds at the base and middle). Present on uplands north and Very friable, coarse- and uniform-grained sandstone, somewhat conglo-TABLE 1. SUMMARY OF GEOLOGIC UNITS ALONG THE CANADIAN RIVER, QUAY south of the river, in local collapse depressions, and in the river Magnesium and calcium carbonate, slightly fractured and permeable. meratic at base; unconformable on Quartermaster formation. miles west of Logan, where it dips under the river. COUNTY, NEW MEXICO. canyon west of Sec. 16, T. 13 N., R. 33 E. DESCRIPTION Marcon shale Santa Rosa sandstone in Guadalupe County) Alibates dolomite Trujillo sandstone: sandstone member Tecovas formation: Chinle formation: (similar to the "Terra cotta" shale member upper member Quartermaster formation: member

About 50 at the state

Reported about 250

in Potter County.

75 at the State to 250 near Logan.

about 1000.

THICKNESS (feet)

Potter County, Texas.

Reported 0-25 in

in Potter County; not County, but observed to be similar to the reported section in Reported about 250 measured in Oldham Oldham County, but Unknown; not known to be exposed in probably present Potter County. shale, with scattered gypsum lenses in upper part, prominent gypsum bed Thick beds of white gypsum. Numerous collapse structures apparent throughout the Canadian River area have probably been caused by solution of this and other gypsum beds. in the lower third.

Bluff gypsum

GENERALIZED GEOLOGIC COLUMN QUAY COUNTY, NEW MEXICO LAKE MEREDITH SALINITY STUDY, TEXAS-NEW MEXICO

ERA PERIOD GROUP FORMATION THICKNESS		FORMATION THICKNESS (FT)	CHARACTER
CENOZOIC TERTIARY QUATERNARY		ALLUVIUM 0-100 0-50	Sand, silt, and gravel in present Canadian River channel. Silty sand, gravel, and clay of high terrace deposits with some aeolian sand and silt.
		OGALLALA 0-200	Sand, gravel, and caliche, with some silt and clay.
		CHINLE 0-865	Shale, siltstone, and silty sandstone with local thin beds of conglomerate and limestone. Three well-defined members may occur; upper and lower members of predominantly shale and siltstone, and a middle member similar to the Santa Rosa sandstone. Greenish-gray to bluish-gray but weathering to brown, red, or purplish red. In Texas, known as Trujillo formation.
NESOZOIC TRIASSIC	DOCKUM	SANTA ROSA 200-450	Silty to clean, fine to coarse-grained, massive to crossbedded gray to bluish-gray sandstone, locally conglomeritic, with thin to thick beds of red and bluish-gray shale and siltstone. Thin beds of carbonaceous shale or soft coal in upper member. In Texas, upper members known as Trujillo formation and lower fine-grained sandstone known as Tecovas formation.
PALEOZOIC PERMIAN	ARTESIA	BERNAL 200-500	Salmon, pink to orange-red to gray shale, siltstone, sandstone, limestone, dolomite, halite, and gypsum.
	NAN	ANDRES 500±	Gray limestone, dolomite, halite, gypsum, and anhydrite.
ū	CTOPTETA	0-80	Fine to medium grained well sorted, gray, tan, or white sandstone, usually cross-bedded with minor shale and siltstone.

Index to TAB 4: Geologic and geomorphic maps which include the study area

A. Surface geologic maps

BEG, 1983 (portion)

BEG, 1981 (portion)

Berkstresser and Mourant, 1966, Plate 2 (portion)

NMSHD, undated, Logan Quadrangle - 48 (portion)

SEO, 1961, Figure 3

B. Subsurface geologic maps

Structure

Collins, 1990, Figure 2

Collins, 1990 Figure A4-1

Collins, 1990, Figure A4-2

Collins, 1990, Figure A4-3

Collins and Luneau, 1986, Figure 4

Dutton and Orr, 1986, Figure 1

Foster, et al., 1972, Figure 10

Gustavson, et al., 1980a, Figure 7

HGC, 1984a, Figure 3

HGC, 1984a, Figure 7

HGC, 1984a, Figure 10

HGC, 1984a, Figure 16

HGC, 1984b, Figure 13

HGC, 1984b, Figure 14

HGC, 1984c, two unnumbered figures
SEO, 1961, Figure 4a and 4b, two each
Trauger and Bushman, 1964, Plate 2 (portion)

2. Isopach

Bassett and Bentley, 1983, Figure 3

Foster, et al., 1972, Figure 2

Foster, et al., 1972, Figure 3

Foster, et al., 1972, Figure 4

Foster, et al., 1972, Figure 5

Foster, et al., 1972, Figure 6

Foster, et al., 1972, Figure 7

HGC, 1984a, Figure 6

HGC, 1984a, Figure 8

HGC, 1984a, Figure 9

3. Salt dissolution

Gustavson and Finley, 1985, Figure 19
Gustavson, et al., 1980a, Figure 29
Gustavson, et al., 1980b, Figure 42
HGC, 1984c, unnumbered figure

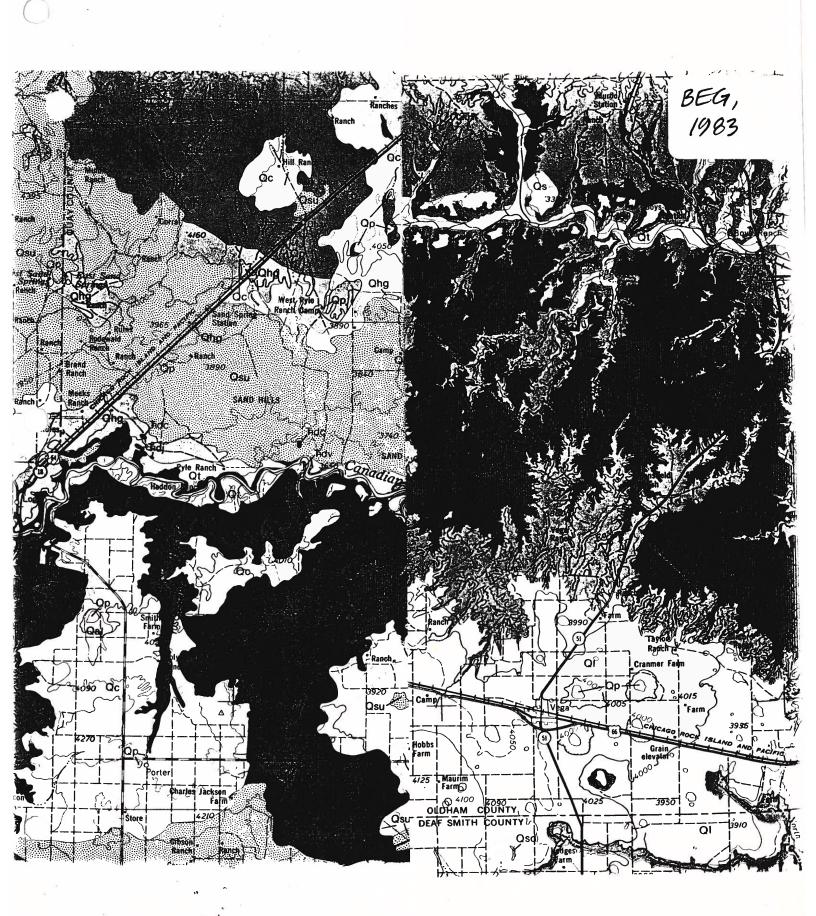
4. Lithology

Bassett and Bentley, 1983, Figure 4
Dutton and Simpkins, 1989, Figure 1

C. Geomorphic maps

Gustavson and Finley, 1985, Figure 2
Gustavson, et al., 1980a, Figure 1
Gustavson, et al., 1980a, Figure 9

TAB 4 Part A. Surface geologic maps

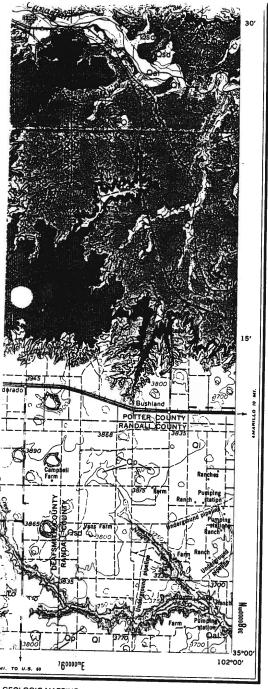


-Pqa ---Pqw

Quartermaster Formation Cloud Chief Gypsum and Whitehorse Sandstone

Cloud Chief Gypsum and Whitehorse Sandstone
Quartermaster Formation. Cloud Chief (hypsum, and Whitehorse Sandstone
undivided, Paw, sandstone, sand, silistane, shale, gypsum, and dolomite
interbedded. Sandstone and sand, fine-grained quartz, scattered to locally
abundant frosted and polished coarse quartz grains, silty, massive, frieble to
indureted, various shades of red and orange, orange-brown, and grayishgreen. Shale and siltsone, sandy in part, indistinctly bedded to massive, indureted, thin interbeds and wins of satin paper in upper part, various shades
of red and orange, reddish-brown, and grayish-green. Gypsum, white gray,
and pink. Alibates Dolomic. Pas. exparately mapped, comprises an upper and lower dolomite separated by shale, upper dolomite locally observed
dolomite locally replaced by chert, which is banded and mottled red, pink,
pale blue, pale purple, gray, brown, and black, forms tedges, a verage thickness 15 feet

Fault U. upthrown: D. downthrown side



F GEOLOGIC MAPPING

ea refers to item in bibliography in "Index to as. 1891-1961," by Brown, T. E. (1963) Bureau ersity of Texas, Austin, For New Mexico, area and Summerson, C. H. (1946) Geology of New Mexico, U.S. Geol. Survey, Oil and Gas net 1; for area B, see Baldwin, B., and eologic studies of Union County, New Mexico. 1 Mineral Res., Bull. 63, Plate 1d.

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Sand. silt. clay, gravel. and raticle. Sand, fine- to coarse-grained quirtz, silty in part. caliche nodules locally, cemented locally by calcite and by silica, locally crossbedded, various shades of gray, brown, and red. Minor silt and clay with caliche nodules, sandy in places, massive, white, gray, olive-green, brown, red. and maroon. Gravel, not everywhere present, composed of pebbles and cobbles of quartz, quartiste, minor chert, unnous rock, metamorphic rock, and clay balls in lower part. Caliche, not everywhere present, sandy, pisolitic, white, gray, pink, may include some caliche of Plesstocene age; thickness up to about 400 feet

Cretaceous

Upper Triassic Group

Dockum

Pebbles and cobbles of silic

:: J.b ::

QTg

ebbles and cobbles of siliceoux seed inentary rocks, intrusive igneous rock metamorphic rocks and sand; caliche zones locally; thickness 50 feet or Caps isolated high ridges west of lower Ute Creek (T14-15N, R31E)

Ogallala Formation

Older gravel deposits ceous sedimentary rocks, intrusive igneous rocks, and

Raton Basalt

Southward continuation of basalt mapped as undifferentiated Clayton Basalt! in nouthwestern Union County, Medium grained olivine basalt, contains pelloubroum olivine phenorysts 2 to 3 mm in diameter, stubby columnar jointing prominent



Dakota Sandstone and Purgatoire Formation ta Sandstone (not separately mapped), gray, conglomeratic, some yellowish-brown shale

yellowish-brown shale
Purpatoire Formation (not separately mapped), consists, from top down, of
Pajarito Shale, gray to green, locally sandstone at top, thickness 50 to 60 feet;
Mesa Rica Sandstone, fine- to coarse-grained, becomes finer grained upward, crossbedded, white to gray, locally reddish, thickness 50 to 100 feet; and
Tucumcari Shale, from top down, shale, limestone, sandstone, and conglomerate; shale, black, contains both Kiamichi and Duck Creek faunas, timestone,
argillacrous; basal sand and conglomerate unconformably rest on Morrison
Formation, thickness up to 20 feet; thickness of Tucumcari Shale up to 60 feet



Morrison Formation and Exeter Sandstone

Morrison Formation, Jm. Exeter Sandstone, Je. and Morrison Formation and Exeter Sandstone undivided, J

Morrison Formation. Jm. clay. shale, and sandstone, red to bluish, variegated; mainly shale; sandstone, light-brownish-yellow; thickness 250 feet

Exeter Formation, Je. Jine-grained, light-brownish-pellow to white to pale-reddish-brown; thickness 140 feet, Equivalent to Entrada Formation on New Mexico State Geologic Map



Chinle Formation

hale, siltstone, sandstone, thin limestone lentils, and mudstone; mostly dusky-red with thin greenish shales; thickness up to 1,200 feet

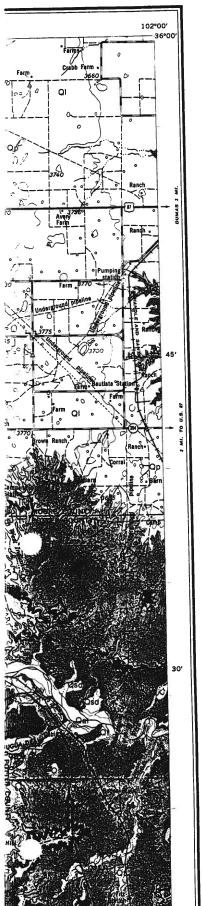


Trujillo Formation

Conglomerate. sundstome, and shale. Conglomerate. sandy, composed of granules and pebbles of quartz, timestome, sandstome, silustone, minor chert, and fragments of petrified wood, massive, gray, brown. Sandstome, conglomeratic, fine to carse grains of quartz and timestone, micaceous, calcareous locally, crossbedded to massive, gray, preenish-gray, brown. Shale, micaceous, ocurs as thin interbeds, gray and red. Forms scarp. Thickness 75 feet, truncated locally truncated locally



GEOLOGIC ATLAS OF TEXAS TUCUMCARI SHEET



EXPLANATION

Qal

Alluvium

Floodplain deposits, includes lowest terraces along Canadian River



Windblown deposits

In Texas, sand and silt, neets, Qs. locally modified by surface wash, and dunes and dune ridges, Qsd. locally: in New Mexico, sand sheets, dunes, and dune ridges undivided, Qsu



Pediment and other covering deposits

rediment and other covering deposits.

Pediment deposits, Qpd. silt, sand, and coarser debris derived chiefly from Triassic rocks and from minimal amounts of overlying rocks; rests on Triassic rocks; occurs in southwestern port of sheet. Other covering deposits, Qc. include collurium, allurial fans, and slopewash deposits in lowlands and along valley walls; in uplands mostly windblown sand



Fluviatile terrace deposits

Fluviatile terrace deposits

Gravel, sand, and silt. Gravel, sandy, composed of pubbles of quartz, quartzite, chert, igneous rock, metamorphic rock, and calishe. Sand, fine to coarse-grained quartz, crossbedded to massive, lenticular, reddish-brown, pink, gray, Silt. sandy, lenticular. Contiguous terraces of different ages separated by solid line

QI Loess

Windblown silt

Qcs

Blackwater Draw Formation

Elackwater Draw Formation
(Previously mapped as Windblown cover sand)

Sand, fine-to medium-grained quartz, sitty, calcureous, locally clayey, catiche
nodules, massive, grayish-red: distinct surface soil profile and buried paleonois: thickness up to 80 feet in northwestern Randall County, feathers out
locally (mostly Illinoian)



Playa deposits

Clay and silt, sandy, light-gray, in shallow depressions (Wisconsinan), mostly covered by thin deposit of Holocene sediment Note: water in depressions not shown



High terrace gravel

Reworked from Ogallala and older gravel deposits, with considerable addition of primary material in deposits nearest Use Creek and Canadian River, with volcanic ash lens northwest of Logan. New Mexico



Rita Blanca deposits

Upper part, annd, bentonitic lay, and thin-bedded calcareous sandstone; thick-ness 50 feet or more, lower part, dark distinctly lominated clay, some inter-beds of saud and thin uniform layers of non-marine dolomite; fossils are plants and small fish; thickness 30 feet or more



Older gravel deposits

Pebbles and cobbles of siliceous sed inventary rocks, intrusive igneous rocks, and
metamorphic rocks and sand; caliche zones locally; thickness 50 feet or more.

Caps isolated high ridges west of lower Ute Creek (T14-15N, R31E)



Ogallala Formation

Ogallala Formation

Sand, sill, clay, gravel, and caliche. Sand, fine-to coarse-grained quartz, silty in part, caliche nodules locally, cemented locally by calcite and by silica, locally crossbedded, various shades of gray, brown, and red. Minor silt and clay with caliche nodules, sandy in places, massive, white, gray, olive-gray, olive-prown, red, and maroon. Gravel, not everywhere present, composed of pebbles and cobbles of quartz, quartitie, minor chert, igneous rock, metamorphic rock, and clay balls in lower part. Caliche, not exemplacer present, sandy, pisolitic, white, gray, pink, may include some caliche of Pleistocene age; thickness up to about 400 feet



Raton Basalt

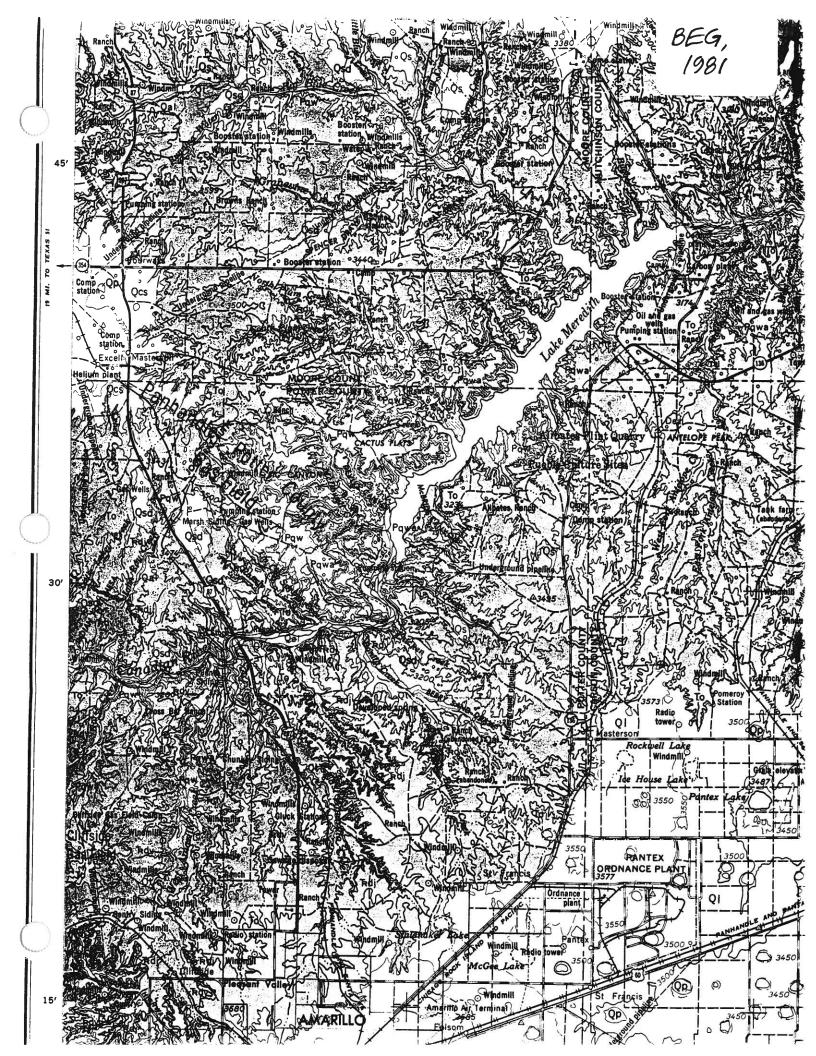
Southward continuation of basalt mapped as "undifferentiated Clayton Basalt" in southwestern Union County, Medium-grained olivine basalt. contains yellow-brown olivine phenocrysts 2 to 3 mm in diameter, stubby columnar jointing prominent



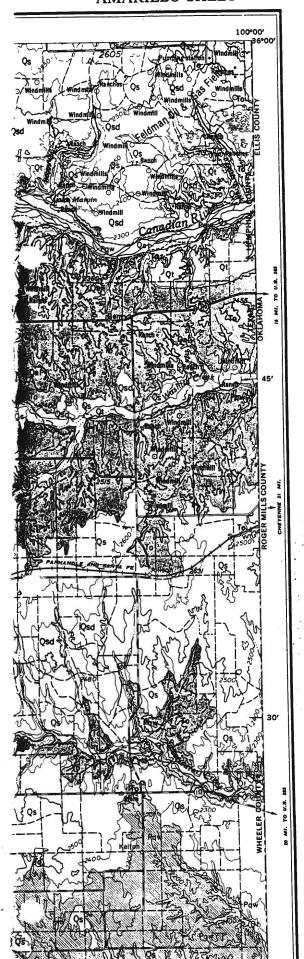
Dakota Sandstone and Purgatoire Formation capped), gray, conclusions

ons

QUATERNARY AND TERTIARY



GEOLOGIC ATLAS OF TEXAS AMARILLO SHEET



EXPLANATION

Sedimentary rocks

Qai

Alluvium

Flood-plain deposits; includes lowest terrace al

Os

Windblown sand
Sand and silt, in sheets, Qs. locally modified by surface wash; dunes and dune ridges, Qed, locally

Ot

Fluviatile terrace deposits

Gravel, sand, and silt. Gravel, sandy, composed of pebbles and cobbles of quarts, quartizite, chert, igneous rock, metamorphic rock, caliche, and rare abraded Gryphaea, Sand, fine to coarse-grained quarts, cross-bedded to massive, lenticular, reddish brown, pink, gray. Silt, sandy, terraces of different ages separated by solid line

Qp

Playa deposits

Clay and silt, sandy, gray in shallow depressions, usually covered by thin deposit of Recent sediment; weathers light gray. (Wisconsinan) Note: Water

QI

Windblown silt

Qcs

Blackwater Draw Formation

d, fine to medium-grained quarts, eilly, calcareous, caliche nodules, mas-sive, pink to grayish red, reddish brown, olive gray; distinct soil profile locally; thickness 25 fest, feathers out locally. (Mostly Illinoian, may include younger deposits)



Ogallala Formation

Ogallala Formation

Sand, silt, clay, gravel, and caliche. Sand, fine to coarse-grained quarte, silty in part, caliche nodules locally, comented locally by calcite and by silica, locally crose-bedded, various shades of gray, brown, and red. Minor silt and clay with caliche nodules, sandy in places, massive, white, gray, olive-green, brown, red, and maroon. Gravel, not everywhere present, composed of pebbles and cobbles of guarts; quartsile, minor chert, igneous rock, metamorphic rock, limestone, clay balls in lower part, and abraded Gryphaen in intraformational channel deposits and in basel conglomerate. Caliche, not everywhere present, sandy, pisolitic, white, gray, pink, comprises four or five bads up to 18 feet thick in upper part, forms ledges and caprock. Maximum thickness-550 feet, thins westward. (Locally includes Ogallala sand which has moved downslope covering older formations)



Dockum Group undivided and

Trujillo and Tecovas Formations

Trujillo and Tecovas Formations
Dockum Group undivided, Rd, in Palo Duro Canyon. Thickness 250 feet. (Elsewhere Dockum is divided into Trujillo and Tecovas Formations.)
Trujillo Formation, Rdj. conglomerate, sandstone, and shale. Conglomerate, sands, composed of granules and pebbles of quarts, limestone, sandstone, siltetone, minor chert, and fragments of petrified tood, massive, gray, brown. Sandstone, conglomeratic, fine to coarse grains of quarts and limestone, minor chert-revue locally, cross-bedded to massive, gray, greenish gray, and brown. Shale, micaceous, occurs as thin interbeds, gray and red. Forms accurs. Thickness 20 fest transacted locally

gray, and brown. Shale, micaceous, occurs as tens reservous, gray and re-forms scarp. Thickness 30 feet, truncated locally was Formation, Rdv. shale, clay, silistons, and sand. Shale, clay, and sili-stone, sandy in places, micaceous, calcareous locally, reddish brown, various shades of red, marcon, gray, greenish gray, yellow, and purple. Sand, fine to medium-grained quarts, locally large petrified logs, unconsolidated, massive, lenticular, white, and light gray. Thickness 275 feet, truncated

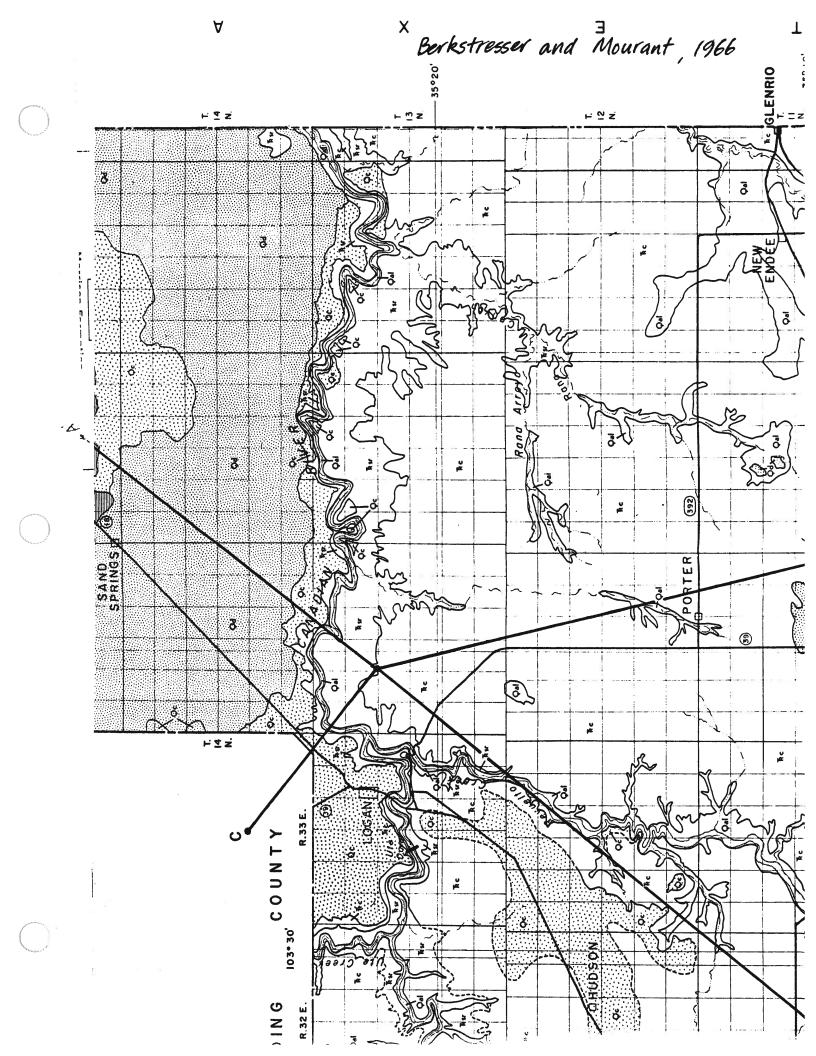


Quartermaster Formation

Cloud Chief Gypsum and Whitehorse Sandstone artermaster Formation, Pq. mapped separately in Palo Duro Canyon and along Mulberry Creek. Shale, siltstone, sandstone, and gypsum, interbedded. Shale and niltstone, sandy, indurated, evenly bedded, thin interbedded and veins of satin spar, various shades of red, reddish brown, and reddish orange. Sandstone, fine-grained quarts, silty, scattered frosted and polished grains, red, reddish orange. Gypsum beds thin and discontinuous. Maximum exposed thickness ISO feet

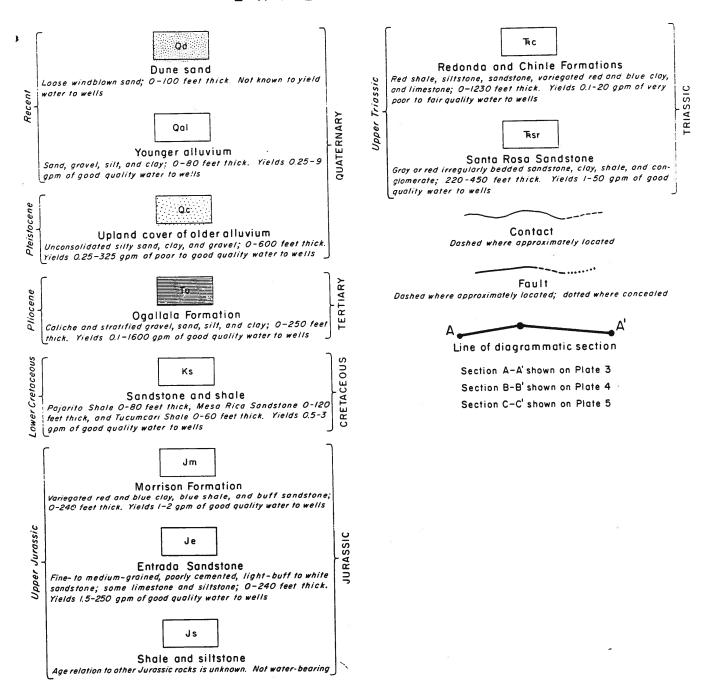
mum exposed increases to Josel

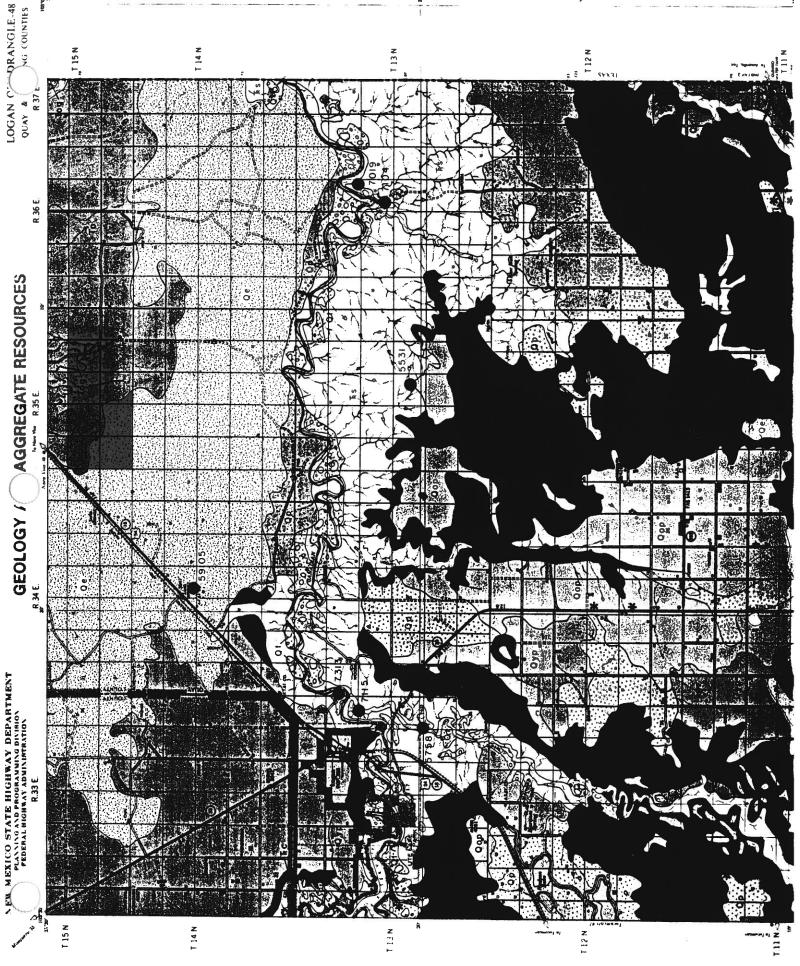
Quartermaster Formation, Cloud Chief Gypsum, and Whitshorse Sandstone
undivided, POW, sandstone, sand, siltstone, shale, gypsum, and dolomite
interbedded. Sandstone and sand, fine-grained quarts, scattered to locally



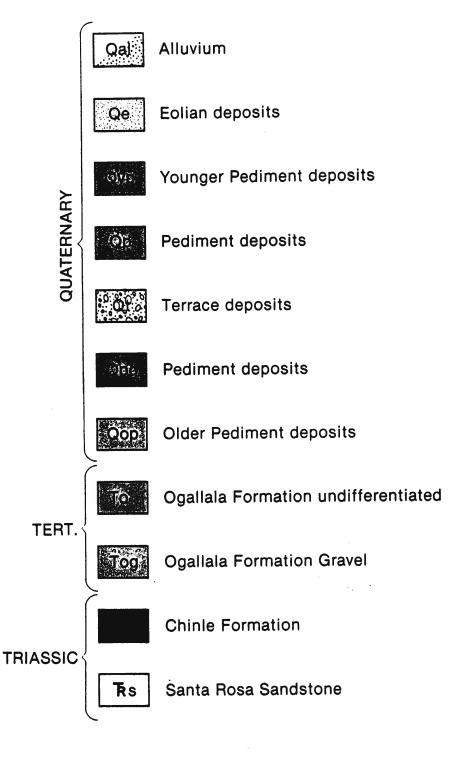
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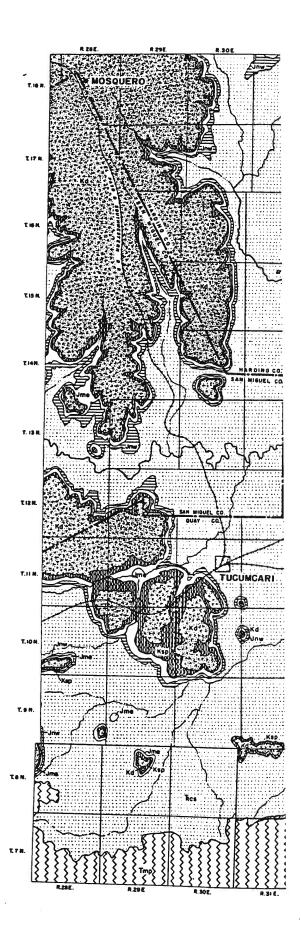
EXPLANATION





EXPLANATION





Regional Geologic Setting

The area investigated encompasses the Palo Duro Basin and adjacent regions within the Texas Panhandle, eastern New Mexico, and the Oklahoma Panhandle (fig. 2). The general stratigraphic section of the Palo Duro Basin is shown in table 1. The

Palo Duro Basin is bordered on the north by the Amarillo Uplift and Bravo Dome, on the south by the Matador Arch, and on the west by the Roosevelt positive and Tucumcari Basin (fig. 2). North of the Bravo Dome and Amarillo Uplift are

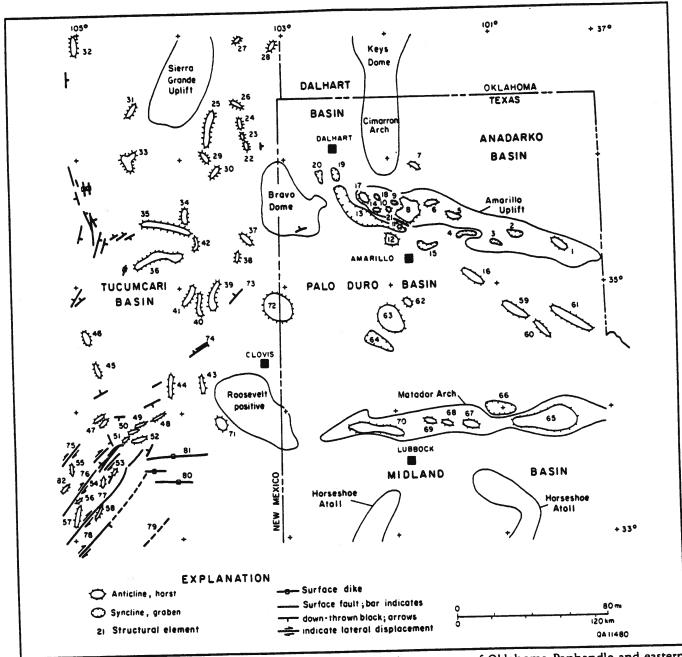
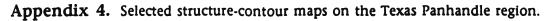


Figure 2. Major structural features of the Texas Panhandle and adjacent areas of Oklahoma Panhandle and eastern New Mexico. Numbered structural elements are listed in appendix 2.



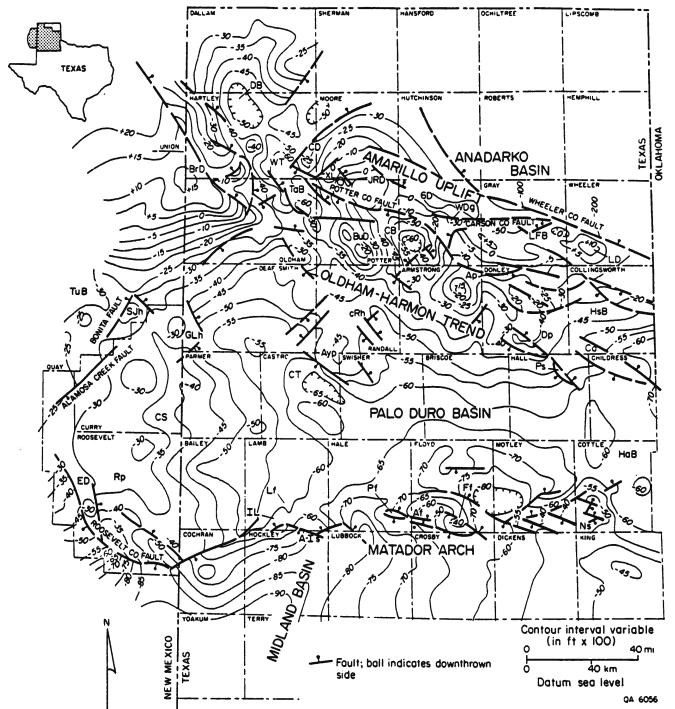


Figure A4-1. Structure-contour map on the top of basement, southern Texas Panhandle (from Budnik, 1989). The structures are abbreviated as follows: 6D = 6666 Dome; Af = Arick field; A-Is = Anton-Irish structure; Ap = Armstrong positive; Ayp = Arney positive; BrD = Bravo Dome; BuD = Bush Dome; Ca = Childress anticline; CB = Carson Basin; CD = Channing Dome; CS = Clovis Sag; CT = Castro Trough; CRh = Carcia Randall high; CRh = Carcia Lake high; CRh = Carc

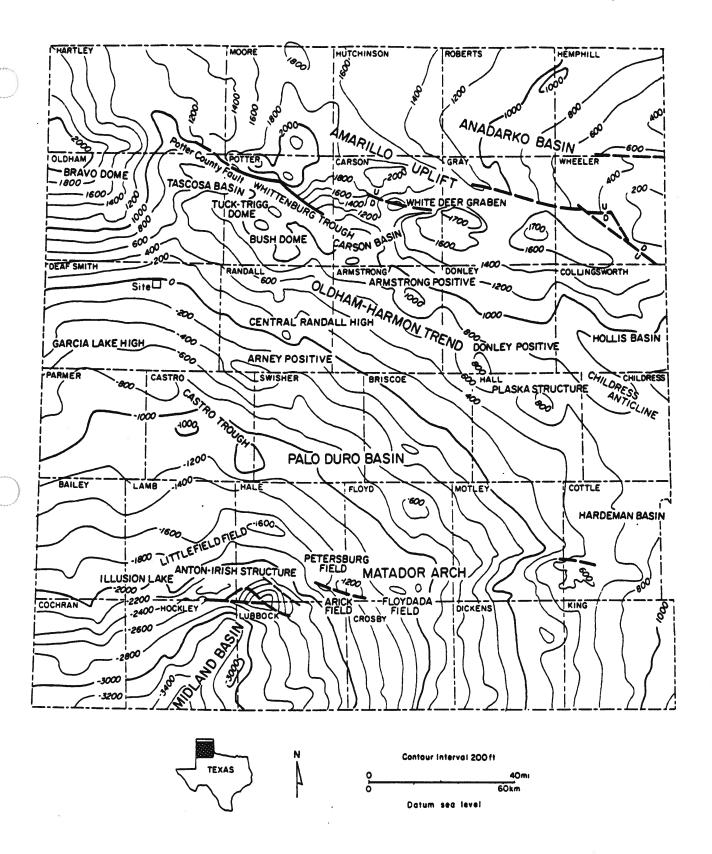


Figure A4-2. Structure-contour map on the top of the Tubb interval, Palo Duro Basin. From Budnik (1989).

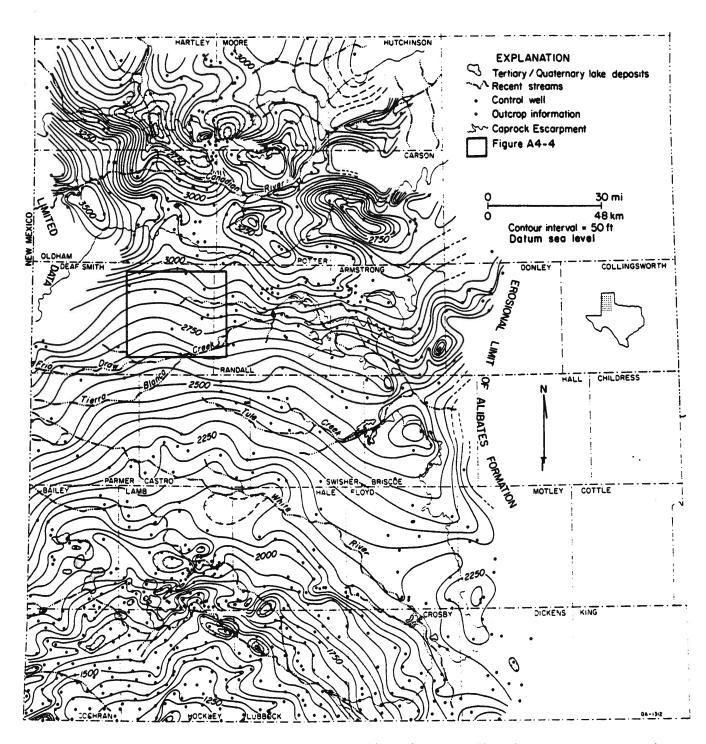


Figure A4-3. Structure-contour map on the top of the Alibates formation. Note that structures are complex and well defined in areas of sufficient data but show little structural detail in areas of sparse data. From Gustavson and Finley (1985).

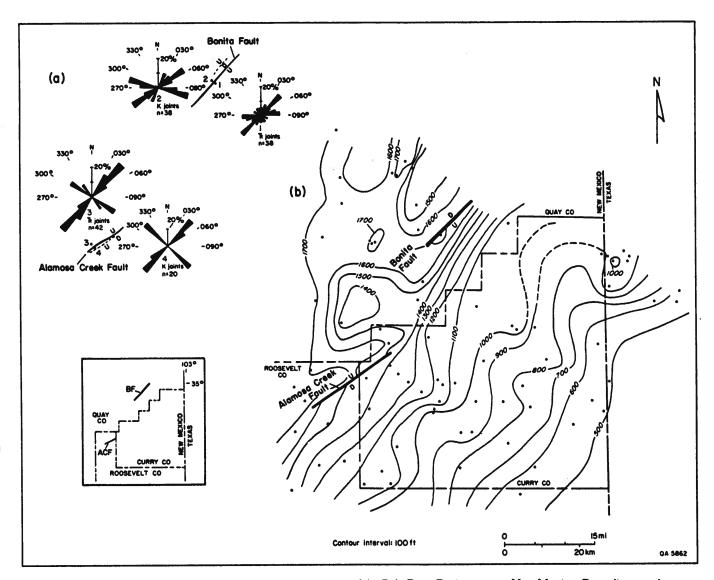


FIGURE 4. (a) Strikes of faults and joints at the western margin of the Palo Duro Basin, eastern New Mexico. Rose diagram data are plotted as percentages of total number of measurements (n) for 10° intervals. (b) Structure-contour map on the base of the Permian San Andres Formation, Quay, Curry, and Roosevelt Counties, eastern New Mexico.

flank of John Ray Dome is a fault exposed at the surface that displaces Permian against Triassic rocks (Eifler, 1969). The strike of this fault is 295°-310°.

Along the southwestern flank of John Ray Dome, joints in Permian Quartermaster (Dewey Lake), Triassic Dockum, and Tertiary Ogallala strata were analyzed using azimuth versus traverse distance plots (AVTD) to detect variability in the strike and in the occurrence of joints along a traverse (Wise and McCrory, 1982). The locations of three traverses used for data collection are shown in figure 5b. The strike of a representative joint from each set was measured at intervals along

each traverse. The joints are almost perpendicular to the sandstone beds; thus even in the gently dipping strata the joints are nearly vertical. Wise and McCrory (1982) describe the methods for plotting and contouring data for AVTD plots.

In the Permian strata, joints occur in two sets (fig. 6a and d). The predominant set strikes 300°-320°. The AVTD plot (fig. 6a) indicates that this set is well defined. Other joints, striking 050°-100°, are less common and appear irregularly along the traverse.

Two joint sets are also in the overlying Triassic Dockum sandstones (fig. 6b and e). Most of the joints strike 300°-320°; this set is well defined along

INTRODUCTION

This report discusses the origin of water and the flow potential of water in low-permeability carbonate rock in the San Andres Formation. The San Andres Formation makes up approximately one-third of an evaporite confining system in the Palo Duro Basin (fig. 1). The 1,800-ft- to 5,000-ft-thick (550-m- to 1,520-m-thick) confining system includes halite, anhydrite, red siltstone and mudstone, limestone, and dolostone (fig. 2) of Leonardian to Ochoan (Permian) age. The research

was part of studies to characterize possible sites for a high-level nuclear waste repository in bedded halite (U.S. Department of Energy, 1984a, 1984b). Hydrologic tests and studies of chemical composition of San Andres brine in the Palo Duro Basin were conducted simultaneously with a regional study of San Andres hydrogeology. The regional picture supports our interpretation of two water samples and of hydrologic tests at six wells in the San Andres Formation in the Palo Duro Basin.

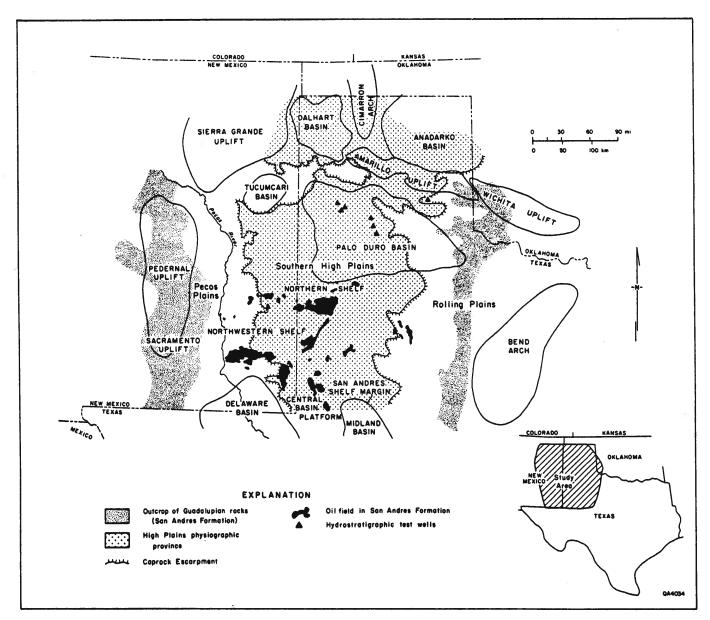
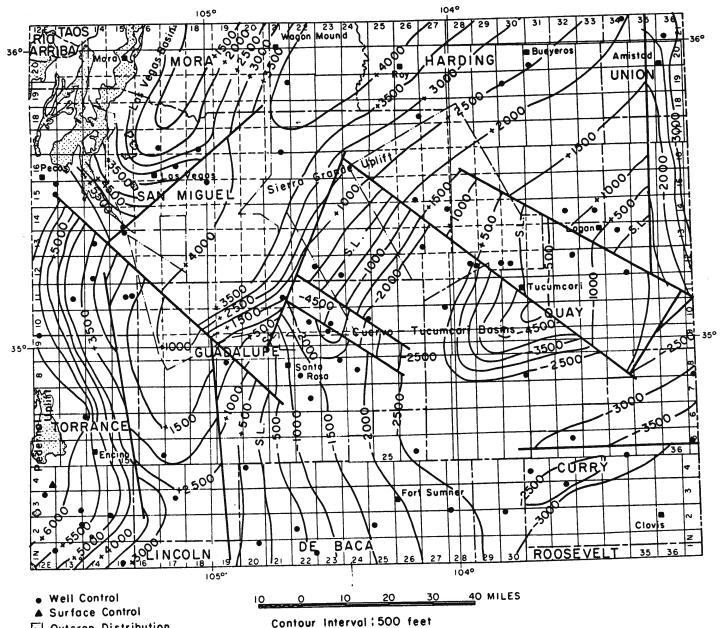


Figure 1. Study area, lying between outcrops of Guadalupian rocks in eastern New Mexico and the Texas Panhandle, includes the Palo Duro Basin, Northern Shelf, Northwestern Shelf, and the northern parts of the Midland Basin, Central Basin Platform, and Delaware Basin.



- Outcrop Distribution
- Faults

Figure 10. Structure Contour Map on Precambrian Surface.

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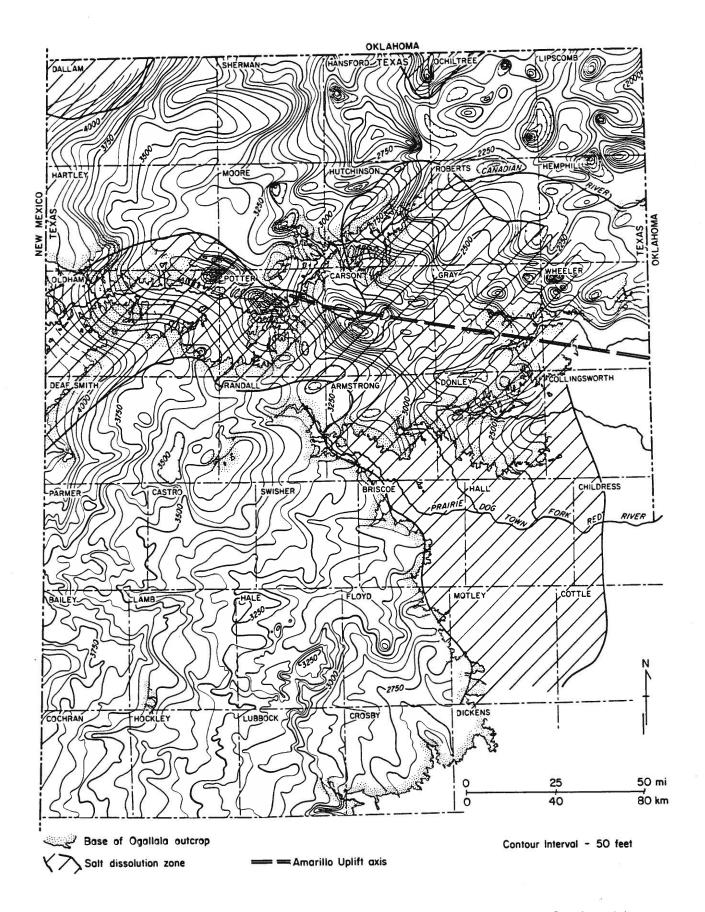
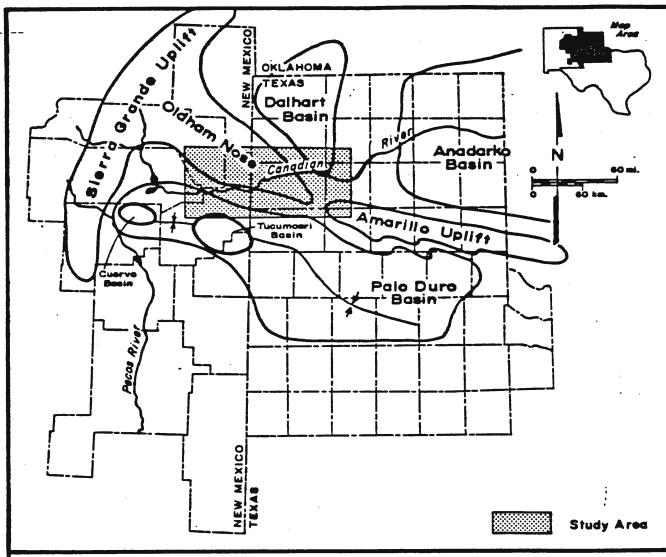


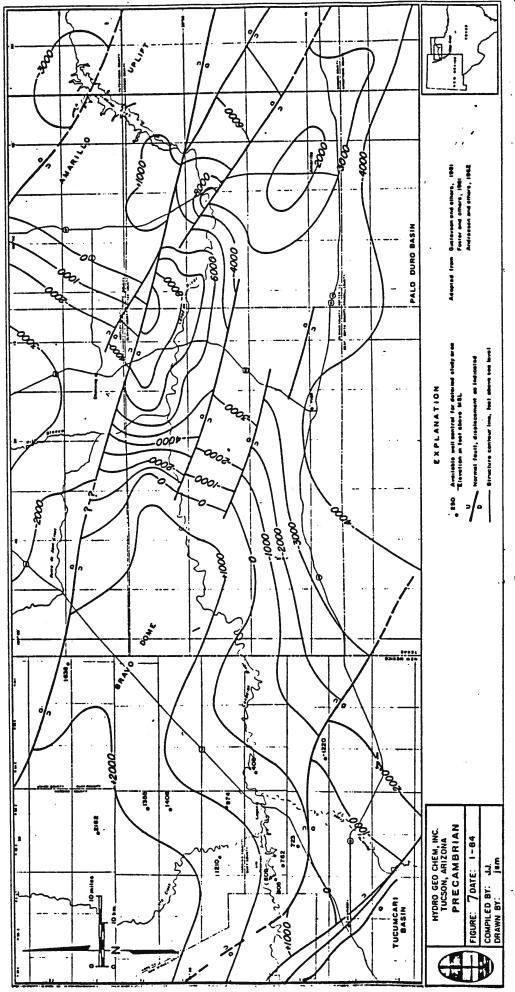
Figure 7. Structure-contour map on the base of the Ogallala Formation (in part from Cronin, 1961). Map also indicates the active salt dissolution zone for the Salado, Seven Rivers, San Andres, and Glorieta Formations.



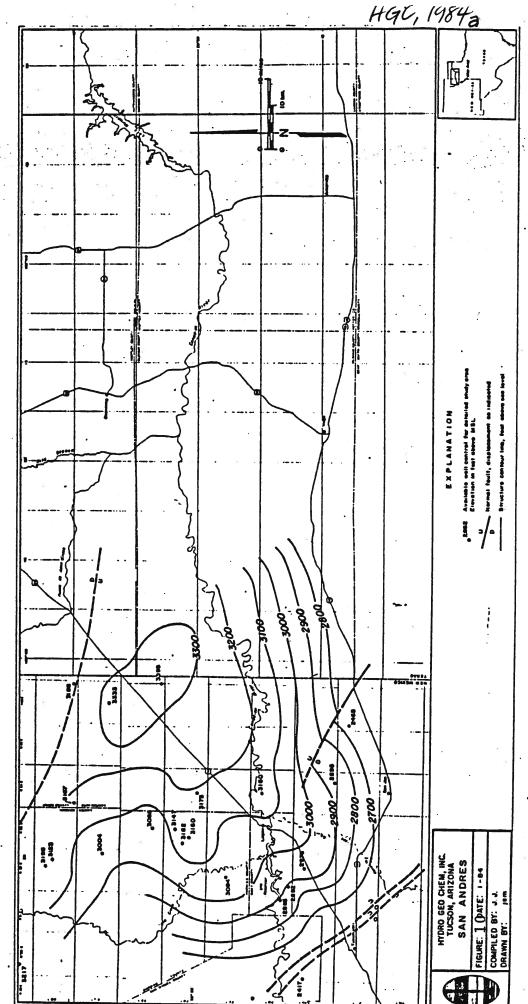
Adapted from Nicholson, 1960; Gustavson and others, 1982

Figure 3. Structural elements in the vicinity of the study area

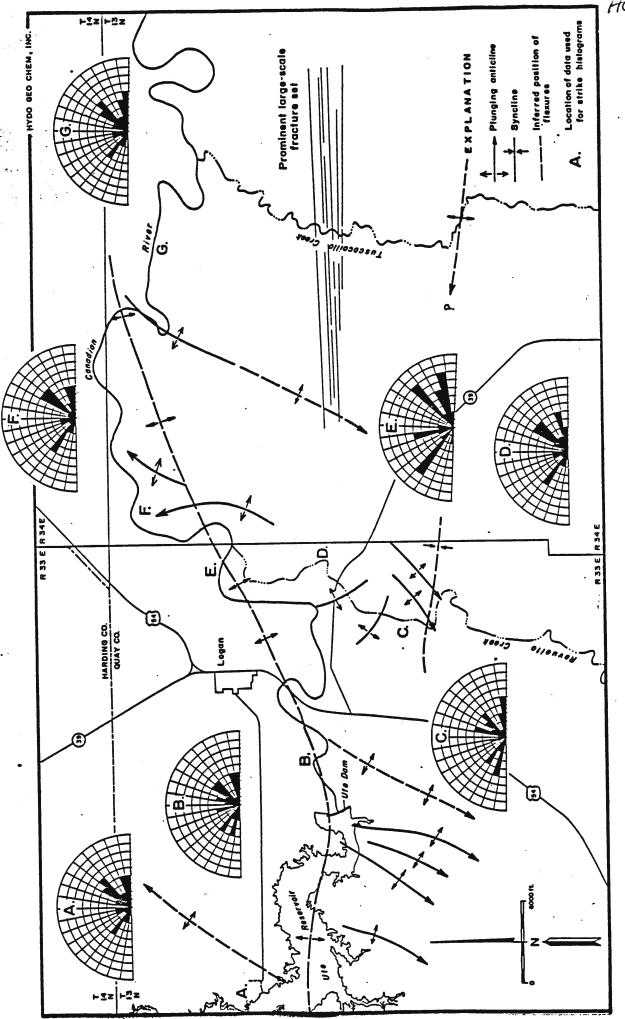
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Pre-cambrian surface Structure contour map of Figure 7.

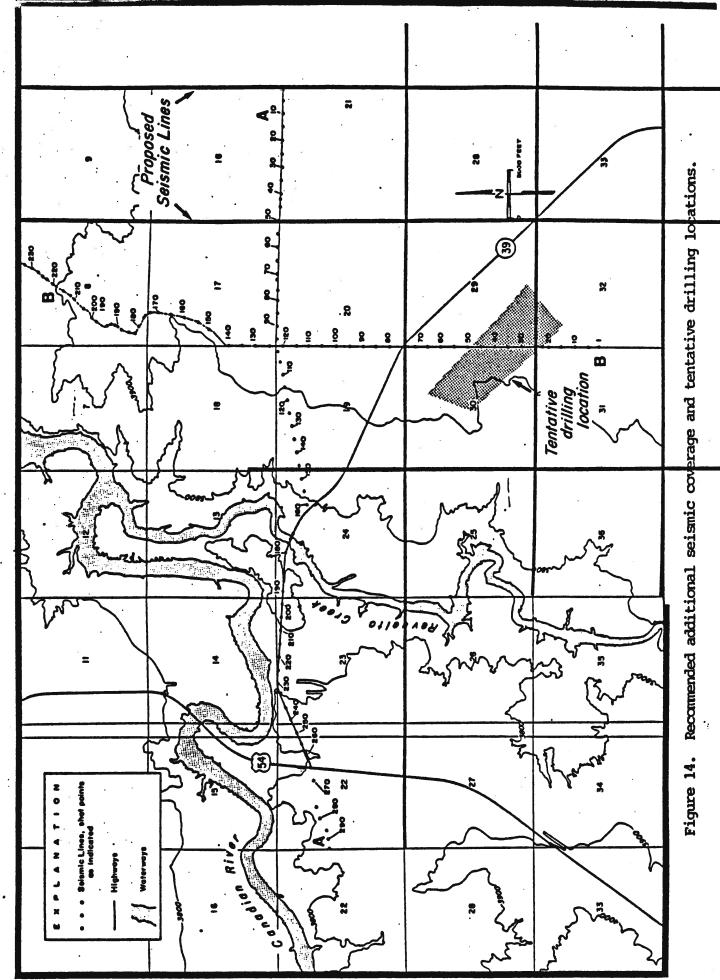


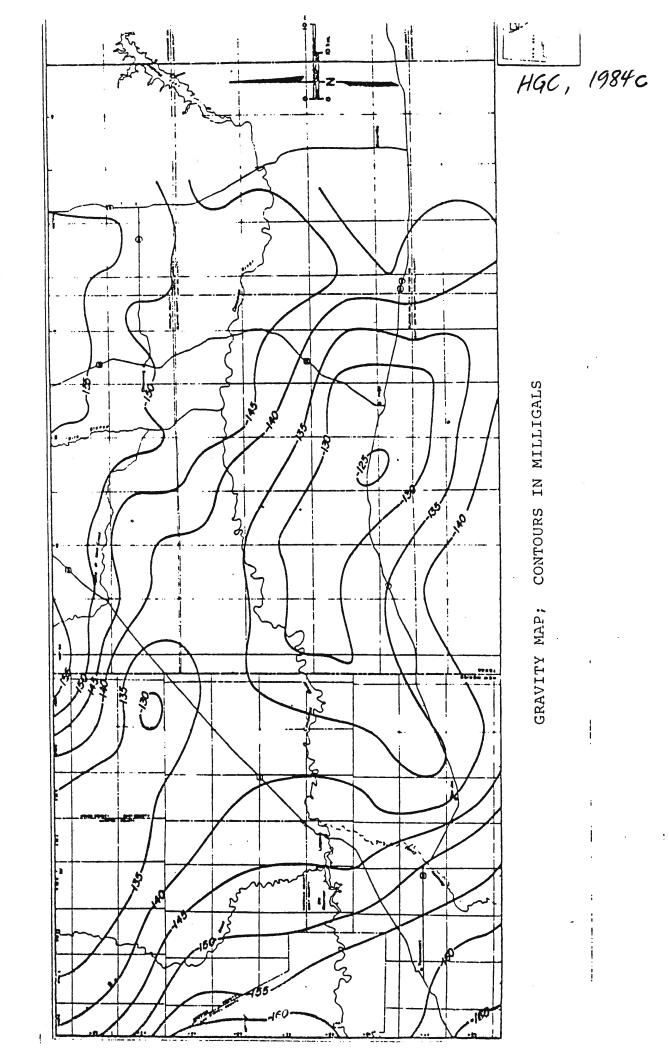
San Andres Formation Structure contour map of Figure 10.



Map showing structural features in the detailed study area Figure 16.

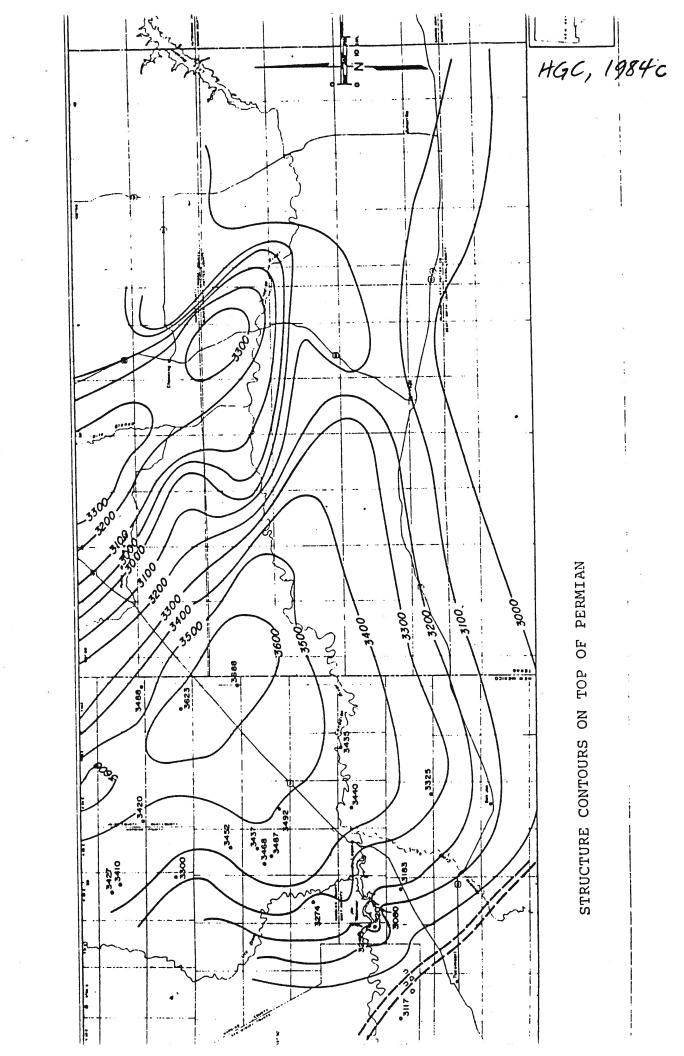
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1600 800 1200-400 800. Worth hest of Res FIGURE 40. GRAPHIC SECTIONS SHOWING CORRELATION OF TRIASSIC, PERMIAN, AND PENNSYLVANIAN STRATA Sangre THAT UNDERLIE NORTH-CENTRAL QUAY AND ADJACENT PORTIONS Location of drill holes and line of section is shown on figure 4b de In Res. Ymiles nest of Pari Santa San Limestone Cristo Stalk Andres COUNTIES, NEW MEXICO Bluff formatton note: Rosa Dam is based on Yeso Pre-Cambrian formation Santa Rosa, uppe . member) formation sandstone member. East of Dam -OF SAN MIGUEL B Southeast of Dans 84 EXPLANATION FOR WELL LOGS Minite or light gray

Measure or light gray

Light rad or pink EXPLANATION FOR MAP Drill hole, operator, number lease, and total depth. SYMBOLS FOR LITHOLOGY Medium or light gray SYMBOLS FOR COLOR Gibson I Parks

3502 Cranic Wells FIGURE 4b. MAP OF PARTS OF QUAY, SAN MIGUEL, AND CURRY COUNTIES, NEW MEXICO, SHOWING DRILL HOLES AND LOCATION OF GRAPHIC SECTION SHOWN ON FIGURE 4a. R29E R30E ROOSEVELT 00. 2200 INT 2300 2400 ETRININ THE ore "-the-00 1 Rinories 3685 2600 Sero-2800 interred from R.32E 2900 area. Salph 14 DAMSITE-2600 CURRY C 000 CO. Den Cois: R.33E. 2500 1 Dr.11,119. 3/00 Red 24 1095 top cf 2340 atson FIGURE 4a and 4b 16987 مرق 1000 DEAF SMITH COUNTY Andres 7200 30.24 \$ 24.7.45 7200 2600 1700 7000 2200 2500

* San Andres tormation Santa Limestone Chinle CH Bluff Rosa member formation sandstone EXPLANATION FOR WELL LOGS MANGUEL CO. 2800 Muchaca 3200 3300 DAMSITE 3400 RANGE STATE OF THE STATE OF THE

dun

FIGURE 40. GRAPHIC SECTIONS SHOWING CORRELATION OF TRIASSIC, PERMIAN, AND PENNSYLVANIAN STRATA Sangre THAT UNDERLIE NORTH-CENTRAL QUAY AND ADJACENT PORTIONS OF SAN MIGUEL de Cristo Pre-Cambrian formation

Union or light gray

Puller Williams

SYMBOLS FOR COLOR

40 Otal

Danube 1218 Maurae 1218 Quay Co. Den 1322 Mailoce

310

ZOA TEXAS

からい。

Uph 746 or pint Middum or light gray

EXPLANATION FOR MAP

Drill hale, operater, number lease, and total depth.

3400

B-ETMON 3700 EDIVE

CURRY CO

3300

3200

\$100

DEAF SMITH COUNT

ROOSEVELT OD. R.30E

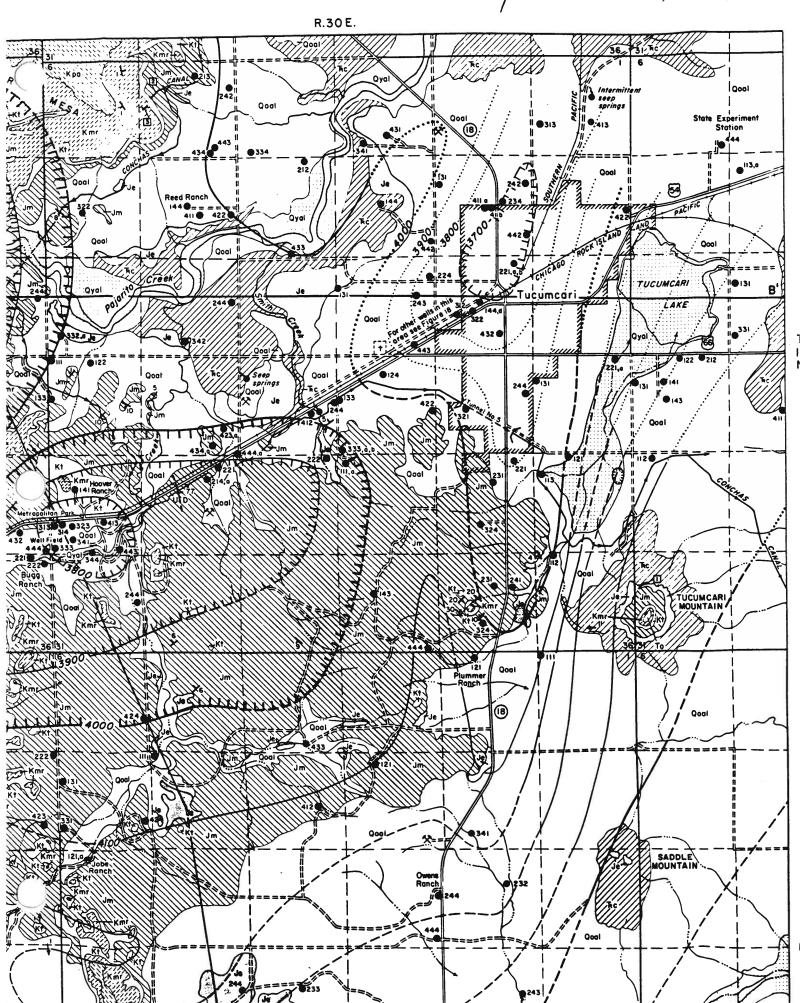
R32E

COUNTIES, NEW MEXICO

Location of drill holes and line of section is shown on figure 4b

FIGURE 4b. MAP OF PARTS OF QUAY, SAN MIGUEL, AND CURRY COUNTIES, NEW MEXICO, SHOWING DRILL HOLES AND LOCATION OF GRAPHIC SECTION SHOWN ON FIGURE 4d. of Chark Bluff

s.thotuo deep Wells. ep Well. preceding map for top Andres - Inferred from dogs interred from logs st a. Watson



Qoal

Older alluvium

Cobbles, pebbles, and sand, some clay; generally unconsolidated but locally firm to moderately well cemented. Water bearing along principal drainageways, and locally where it occurs as thick deposits under the plains or upland basins

То

Ogallala Formation

Caliche caprock on Tucumcari Mountain, gravel, sand, and conglomerate elsewhere. Not water bearing



Pajarito Shale

Sandy shale, and sandstone, locally conglomeratic near base; generally soft. Not water bearing



Mesa Rica Sandstone

Fine- to medium-grained, generally well cemented, cliff-forming. Not water bearing

Kt

Tucumcari Shale

Mostly a shaly sandstone or siltstone, commonly calcareous and fossiliferous. Not water bearing



Morrison Formation

Sandstone, shale, and clay; mostly hard, fine- to coarse-grained sandstone in the upper part, shaly in the middle section and clayey at the base. Yields water to wells locally but generally in small quantities



Entrada Sandstone

White fine- to very fine-grained, soft, generally unconsolidated. Principal aquifer of the area

Pleistocene

Pliocene

Lower Cretaceous

Trauger & Bushman, 1964

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-

QUATERNARY

TERTIARY

CRETACEOUS

JURASSIC

Upper Jurassic

Upper Triassic

Mis. , a sharp solusion & a shis. a z, commonly calcareous and fossiliferous. Not water bearing



Morrison Formation

Sandstone, shale, and clay; mostly hard, fine- to coarse-grained sandstone in the upper part, shaly in the middle section and clayey at the base. Yields water to wells locally but generally in small quantities



Entrada Sandstone

White fine- to very fine-grained, soft, generally unconsolidated. Principal aguifer of the area



Redonda Formation and Chinle Formation

Mostly red shale and clay in the upper part; sandstone and siltstone reported in drill cuttings from the lower part. The Chinle locally yields small quantities of water to wells

-4200--....

Structure contour

Drawn on top of the Chinle Formation; dashed where approximate; dotted where based on inadequate control. Contour interval 100 feet. Datum is mean sea level

Geologic contact

Fault, showing dip

Dashed where approximate. U, upthrown side; D, downthrown side

Strike and dip of beds

Direction of prevailing dip

Measured section

Line of geologic section

●²⁴¹

Well or bore hole -

Number indicates location within section (see figure 2)

Spring

TAB 4 Part B.2 Subsurface geologic maps: isopach maps

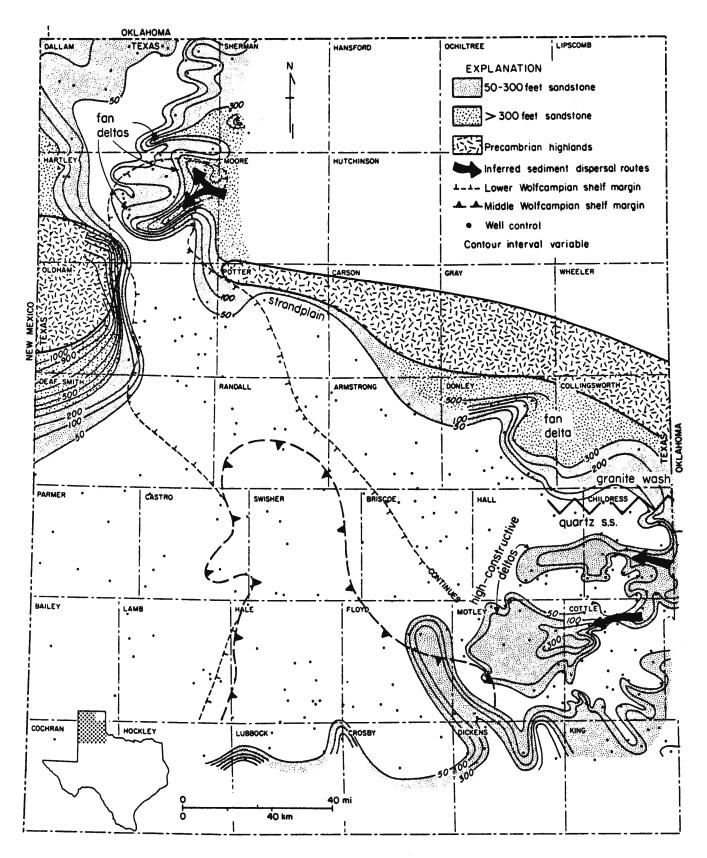


Figure 3. Net-sandstone map of Wolfcampian Series (from Dutton and others, 1979).

Exeter Sandstones. He described the Exeter (Ocate) as a gray, medium to massive-bedded cross-laminated sandstone about 50 feet thick. Wanek (1962) observed that it consists of massive beds of conspicuously cross-laminated fine-grained grayish-orange to reddish-orange sandstone again about 50 feet thick. Wanek suggests the origin is eolian based on cross-lamination and uniform lithology. From well data a range in thickness from 40 to 228 feet is indicated although the thickness is normally less than 100 feet. It is about 175 feet thick on the north face of Tucumcari Mountain. The maximum thickness of 228 was noted in well logs in Metropolitan Park west of Tucumcari (Trauger and Bushman, 1964, p. 149).

TRIASSIC SYSTEM

Rocks of Triassic age are extensively exposed in east-central New Mexico.

Figure 2 is an attempt to restore the thickness of Triassic

rocks left following pre-Exeter and pre-Cretaceous erosion cycles. The interval thickens to the south, and there is a suggestion of the beginnings of the thinning south of the overlap of the Jurassic by the Cretaceous. In western New Mexico the thinning is very abrupt south of the Cretaceous overlap. The absence of Cretaceous rocks beneath much of the Llano Estacado probably will limit the determination of events of similar nature in this area.

In general the Triassic has been subdivided into the Santa Rosa Sandstone at the base and the Chinle Formation above, or the entire interval is referred to as the Dockum Group. In the Tucumcari area Dobrovolny and Summerson (1946) named a contrasting upper interval the Redonda Member. Later Griggs and Read (1959) elevated this interval to formational status. In addition to these more formalized units a middle sandstone-conglomerate interval has been recognized in much of the area.

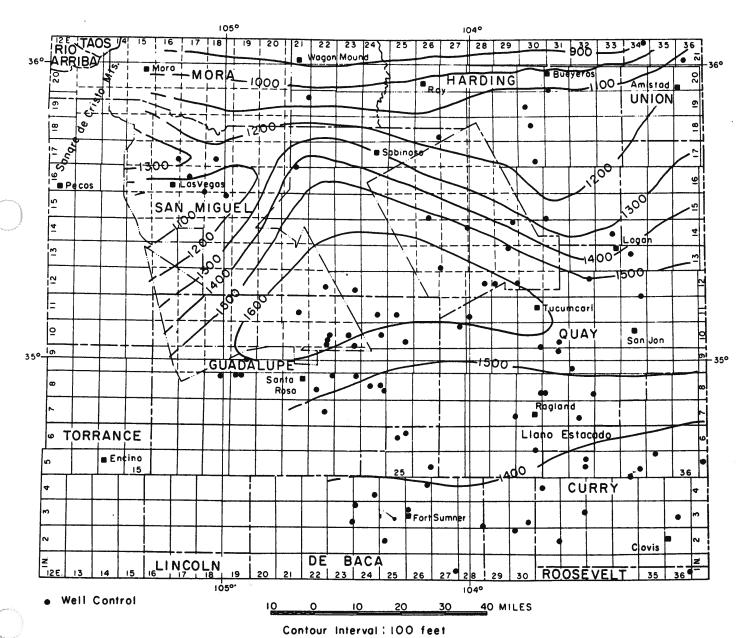


Figure 2. Isopach Map of Triassic Rocks (Restored).

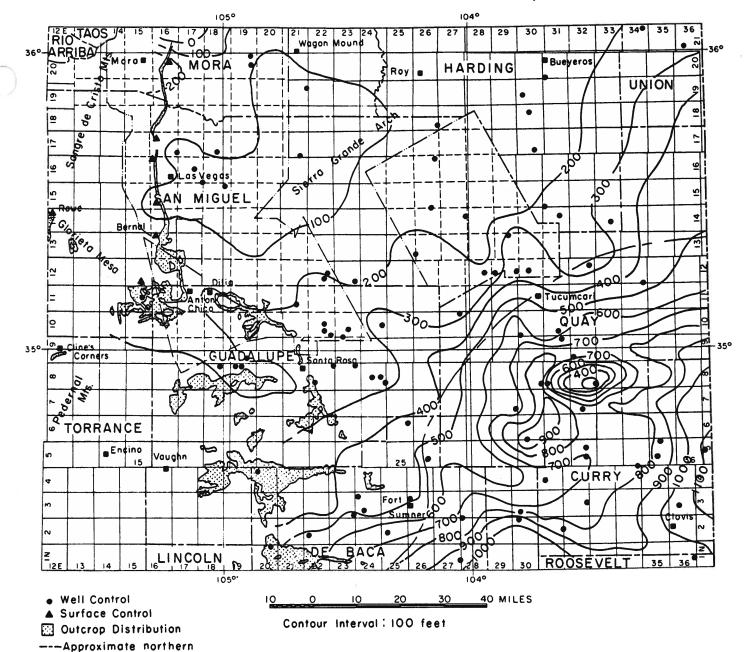


Figure 3. Isopach Map of Bernal-Artesia Group.

Glorieta-like sandstones. Beds of salt extend west to the Cuervo-Fort Sumner area and as far north as the vicinity of Logan. The extent of salt is shown on the isopach map.

limit of solt

Similar to the Artesia Group sediments the thickest sections of the San Andres occur in the southeastern part of the area where it locally is as much as 1,200 feet thick. From here it thins to the north and west with some rather striking areas of local thinning. The thin areas southeast of Santa Rosa and northwest of Tucumcari probably reflect continuing adjustments along the structural features discussed later. Whether this represents continued movement along the suggested faults or compaction in basinal areas interspersed with structurally stable high areas is not known. In any event the potential for combined stratigraphic-structural traps would appear to be good. The high northwest of Tucumcari also is reflected in the isopach maps of underlying units but the same is not true of the thin area of San Andres-Glorieta south of Santa Rosa. In

part of course this may represent a facies change into Yeso-like sediments and thus the apparent thinning may be misleading as far as the over-all Permian thickness is concerned. In any event the typical facies of the San Andres does thin in this area.

At Rowe the Glorieta consists of 190 feet of fine-grained, cross-bedded sandstone with minor dark-red and greenish-gray shale at the top. West of Las Vegas Northrop et al. (1946) show a range in thickness from a little over 100 feet at Montezuma to 290 feet at La Manga. Sections north of Montezuma are in excess of 200 feet thick so a general northward thinning cannot be demonstrated in this area. At the Mora River Gap the Glorieta consists of 205 feet of fine-grained sandstone with minor thin shales and some siltstone. In the Ocate area Bachman (1953) describes the Glorieta as gray to light-brown, medium-grained, cross-laminated sandstone. In a section four miles northwest of Ocate he measured 256 feet but notes that it is absent in the Cimarron Mountains

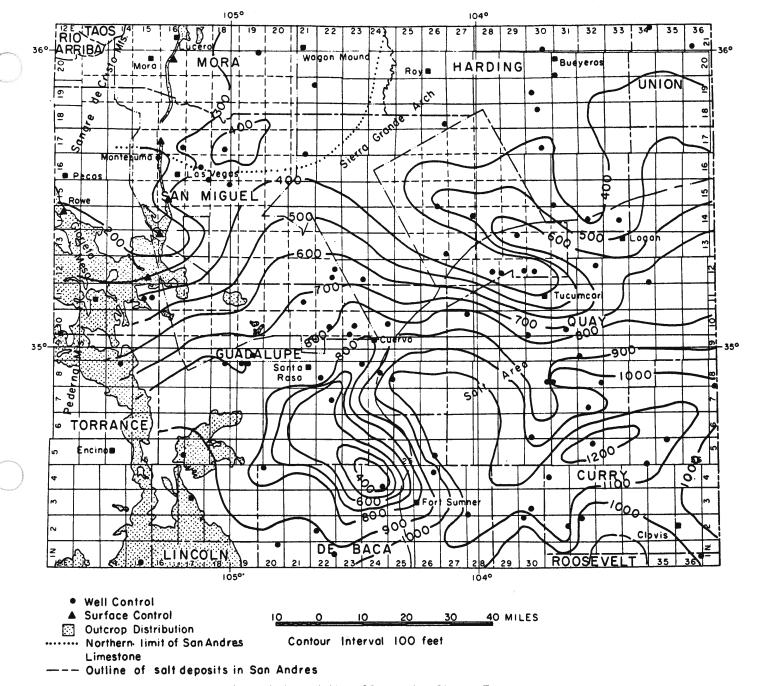


Figure 4. Isopach Map of San Andres-Glorieta Formations.

15 miles north of Mora County. He suggests the possibility that the Glorieta intertongues to the north with the Sangre de Cristo Formation. Baltz and Bachman (1956) note that north of the Mora River Gap the Glorieta coarsens in its lower part and in the vicnity of Lucero Gap contains some conglomerate at the base. They state that the Glorieta grades laterally into the Sangre de Cristo Formation between Guadalupita and Black Lake (north of the map area).

In the Tucumcari area of Quay County the Glorieta is thin or absent. As observed by Bates (Dobrovolny and Summerson, 1946) the eastern extent of Glorieta-like sandstone is approximately R. 31 E. Locally to the east there are sandstones resembling the Glorieta in the approximate position of this

sandstone but these may represent isolated development of offshore bars beyond the general limit of Glorieta deposition.

Yeso Formation

For the most part the Yeso Formation has not been subdivided into the various members used in the oil field areas or those used to the west in central New Mexico. Meseta Blanca and San Ysidro members were used by Read et al. (1945) and others on Glorieta Mesa and locally in the subsurface of eastcentral New Mexico. Some geologists have used Clear Fork, Tubb, Cimarron and Fullerton for various parts of the Yeso indicating possible correlations with subsurface and surface sections to the south and northeast. The Cimarron Anhydrite is a useful marker bed in most of the area.

As can be seen on the isopach map (Fig. 5) the thickest sections are to the southeast and the unit thins quite uniformly to the north and northwest. The thinning is considerable from 3000 feet in the southeast to only 160 feet in a well drilled west of Wagon Mound. On the higher parts of the Sierra Grande arch it is less than 400 feet in thickness. Earlier structural elements also are reflected in the Tucumcari-Logan areas and west of Santa Rosa.

The Yeso Formation crops out in the slopes of Glorieta Mesa, along the foothills west of Las Vegas and extensively around the Pedernal Mountains where it overlaps the Abo Formation onto the Precambrian. At the Rowe section it consists of orange to dark-red sandstone, siltstone and shale. Sandstones are mostly very fine-grained with coarser grains

occurring in some beds and apparently increasing in amount toward the base of the section. In the upper part of the Yeso there are three thin silty carbonate intervals, the lower of which contains chert nodules. The section remains quite uniform in exposures to the east along Glorieta Mesa and continues to contain some carbonate in the upper part. North of Agua Zarca the unit thins rapidly but still contains one or two thin beds of dolomite. At Mora River Gap dolomites are lacking. The upper part contains very fine- to fine-grained sandstone with thin intervals of dark-red shale. Grain size increases in the lower part and the basal thirty feet contains small pebbles of quartz, quartzite, and feldspar. Baltz and Bachman (1956) note that a short distance north of Mora the Yeso grades laterally into arkose and red siltstone that cannot be distinguished from the Sangre de Cristo Formation.

The Yeso is 546 feet thick at Rowe; 230 feet at

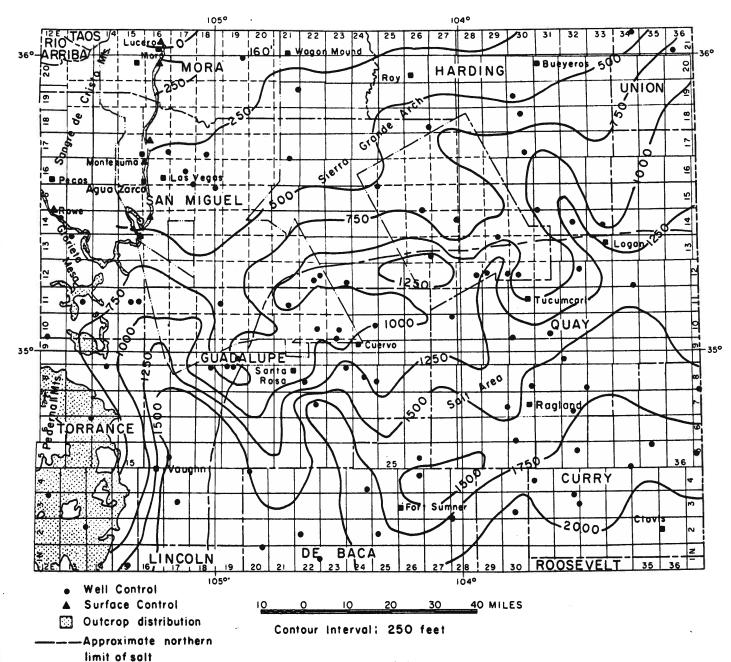


Figure 5. Isopach Map of Yeso Formation.

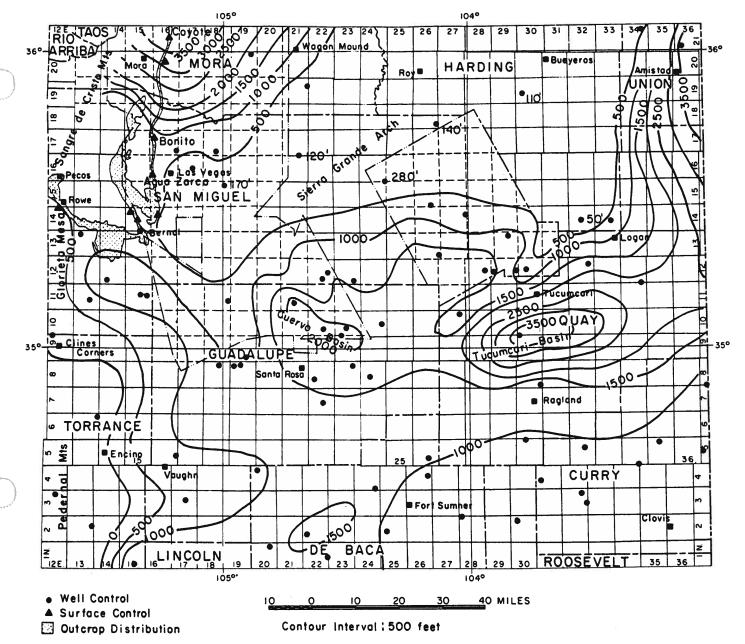


Figure 6. Isopach Map of Sangre de Cristo-Abo Formation.

the La Pasada thickens abruptly and the entire section is dominated by terrigenous material (Flechado Formation). Limestones are sparse in the lower part of the Flechado but similar to those of the Pecos Canyon section.

The La Pasada-Flechado formations are Morrowan to middle Desmoinesian in age. The source of the clastic material was the Uncompaghre highland to the west. Sutherland suggests that the source area consisted of Precambrian metasediments because of the low percentage of feldspar in this part of the Pennsylvanian. The La Pasada was deposited on a marine shelf named the Pecos shelf by Sutherland. This shelf was flanked on the north by the Taos trough (Rowe-Mora basin of previous workers), that extended east of Las Vegas and west at least as far as Santa Fe. The shelf was bordered on the south by the Pedernal uplift; a positive element that apparently remained above sea level during most or all of Pennsylvanian time. The lower part of the Flechado Formation

consists of a series of alluvial deposits alternating with minor near-shore marine units. The upper part consists of anastomosing alluvial fans, that thin eastward, and include a few interbeds of shallow water near-shore marine deposits.

The Alamitos Formation consists of arkosic sandstones and conglomerates, limestones, shales, and siltstones. Almost all the limestones are sandy biosparudites to biosparites and in addition to quartz almost all contain some feldspar. The Uncompaghre highland continued to be active during deposition of the Alamitos Formation and remained high during the rest of Pennsylvanian time and at least locally well into Mesozoic time.

Extensive granite outcrops supplied the coarse arkosic material of the Alamitos Formation. The lower part of the interval consists of alluvial fan and deltaic deposits with periodic shallow marine invasions. The upper part represents the final major marine invasion of the southern part of the

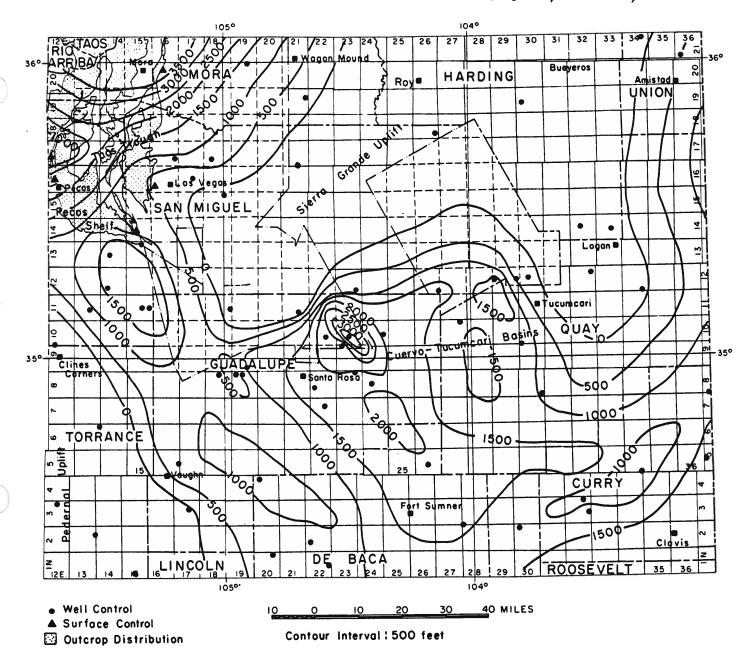


Figure 7. Isopach Map of Pennsylvanian Sediments (Restricted).

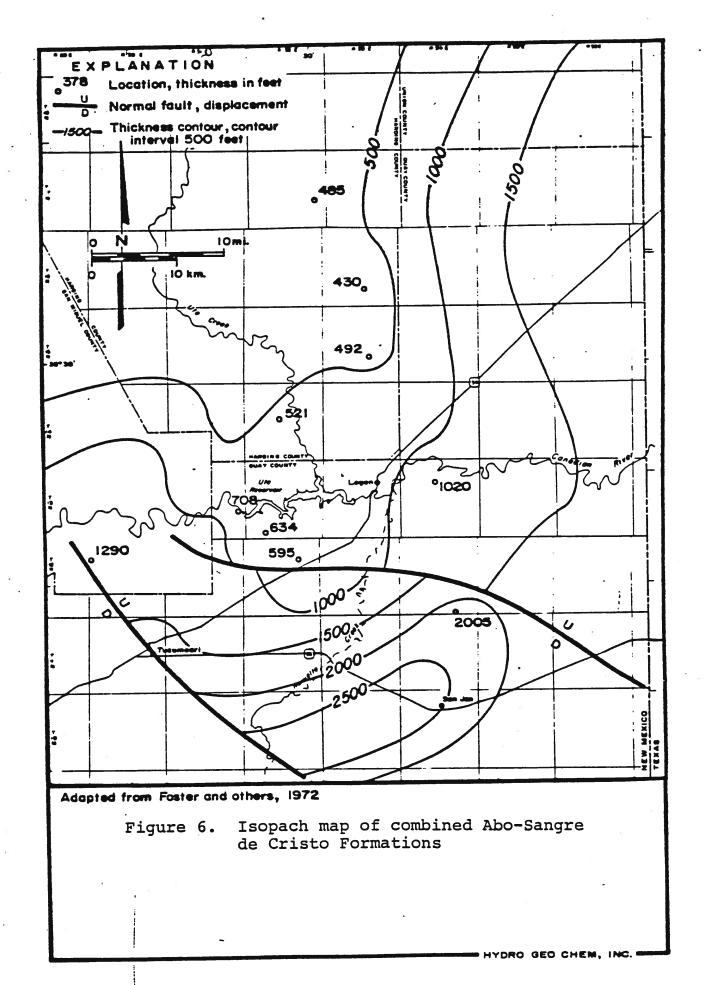
of this interval in the Pecos Valley. The section thins to the east in the subsurface of Mora County. Total thickness of known Pennsylvanian strata are 3,655 feet in the Mora River Gap section but part of the sequence may have been cut out by thrusting. The total thickness measured by R. L. Bates and R. W. Foster in Pecos Canyon is 2,303 feet.

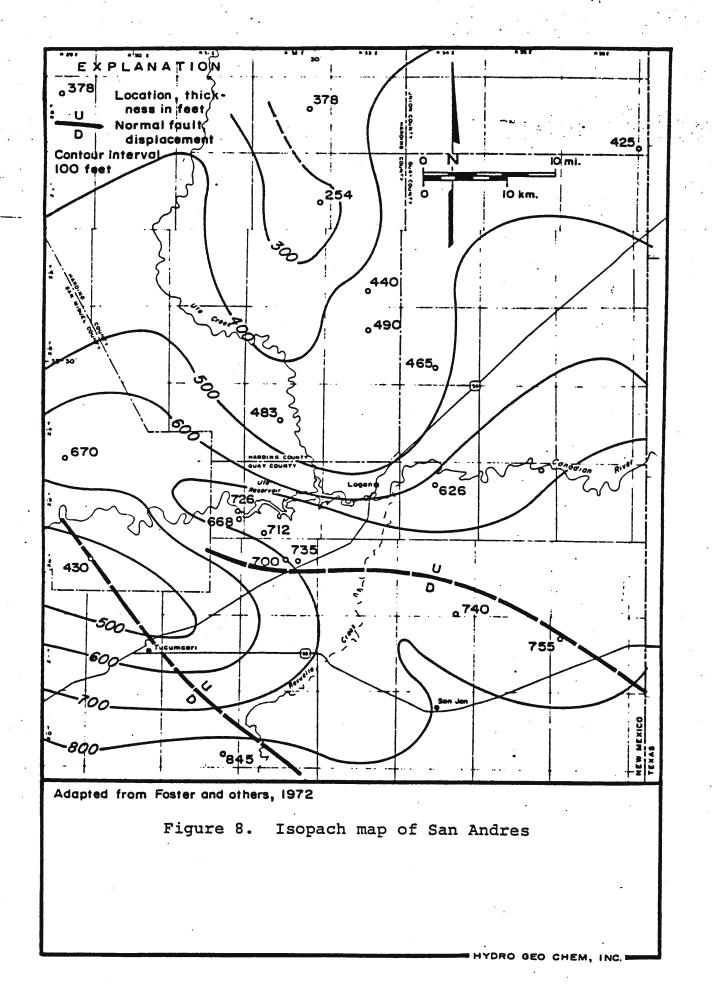
MISSISSIPPIAN SYSTEM

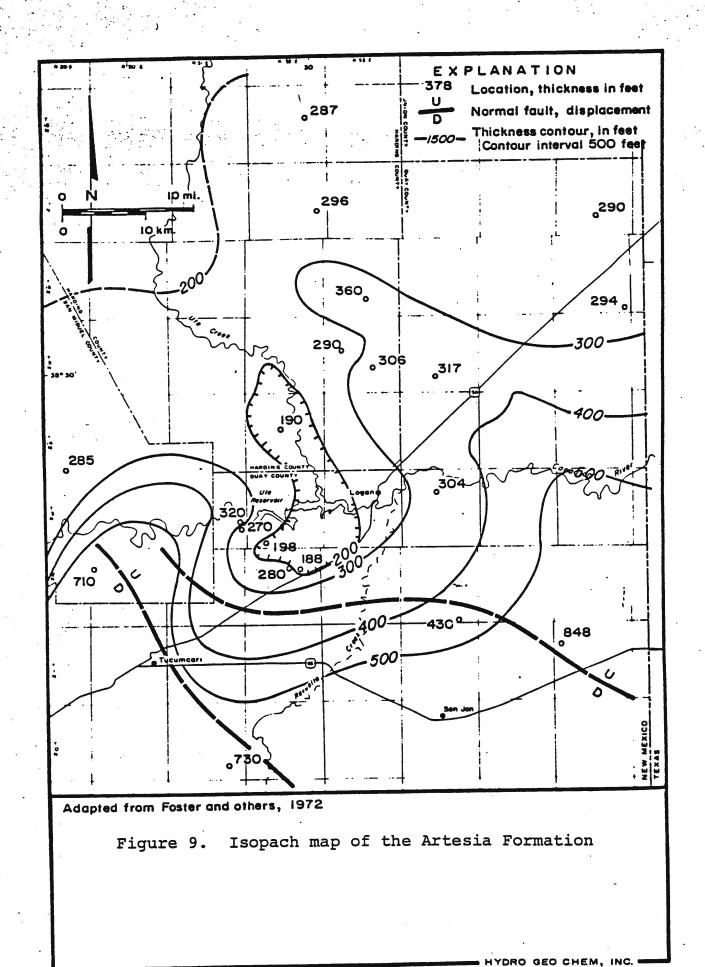
Rocks of Mississippian age crop out in the Sangre de Cristo Mountains and have been penetrated in numerous wells in east-central New Mexico. The area underlain by these rocks is shown in Figure 8. Mississippian strata consist for the most part of carbonate with minor shale and sandstone. In the area from Tucumcari to Santa Rosa and south to Fort Sumner rocks of this age range in thickness from approximately 25 to 170 feet. From Santa Rosa west the thickness varies from

about 20 to slightly over 140 feet. Thus throughout the area of preserved Mississippian strata about 100 feet of section can be expected although normally the interval would be somewhat thinner.

Mississippian rocks were originally assigned to the basal member of the Sandia formation and considered Pennsylvanian in age. Many early geologists realized that this member likely was of pre-Pennsylvanian age and Henbest (1946) reported *Endothyra baileyi* (Hall) from the lower member of the Sandia in the Sangre de Cristo Mountains and on the basis of this foraminifera correlated this member with the Leadville Limestone of the San Juan Mountains. Armstrong (1955, 1958, 1967) correlated the Mississippian of the Sangre de Cristo Mountains with his Arroyo Penasco Formation with type locality in the Nacimiento Mountains, and on the basis of foraminifera considered the rocks to be late Osage and early







TAB 4 Part B.3 Subsurface geologic maps: salt dissolution

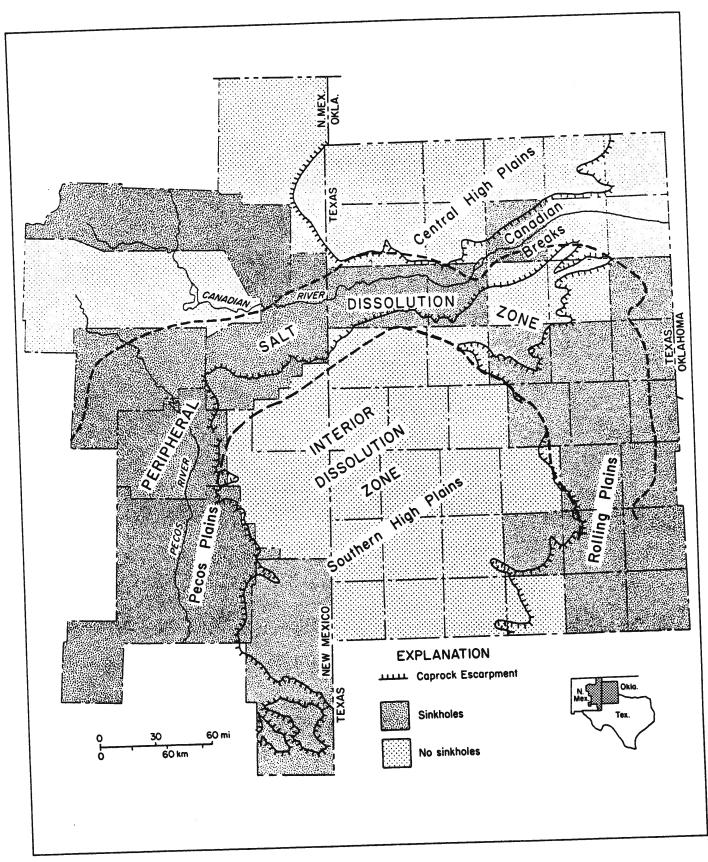


Figure 19. Texas and New Mexico counties, from which sinkholes and fractures have been reported, lying mostly within a peripheral dissolution zone that encompasses the Pecos Plains, the Canadian River Breaks, and the Rolling Plains. Sinkholes and open fractures have not been reported from the Southern High Plains or the Central High Plains within eastern New Mexico or the Texas Panhandle.

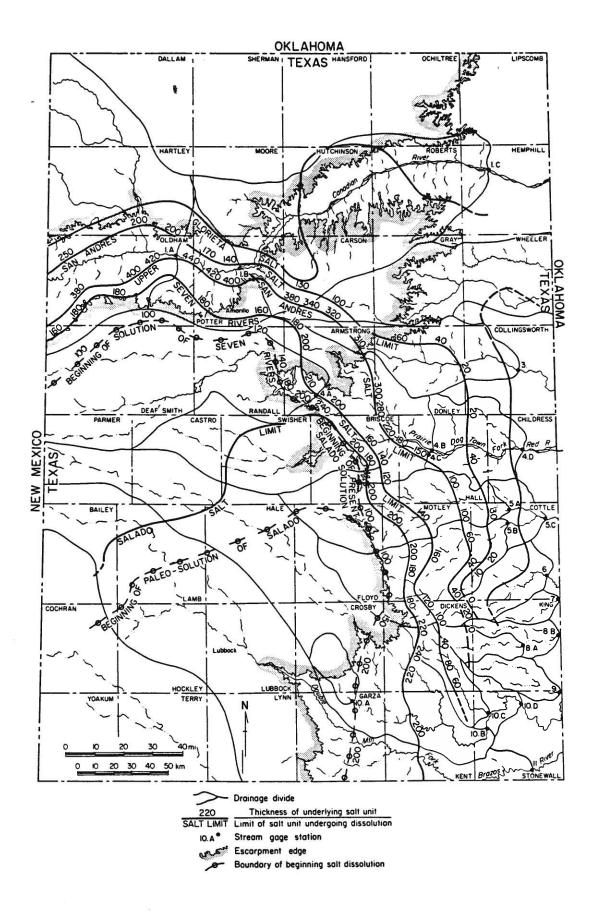


Figure 29. Watersheds and associated water quality sampling stations. Lines marked "beginning of dissolution" indicate the limit of the complete salt section for the Salado and Seven Rivers Formations. Salt dissolution occurs north and east of these lines.

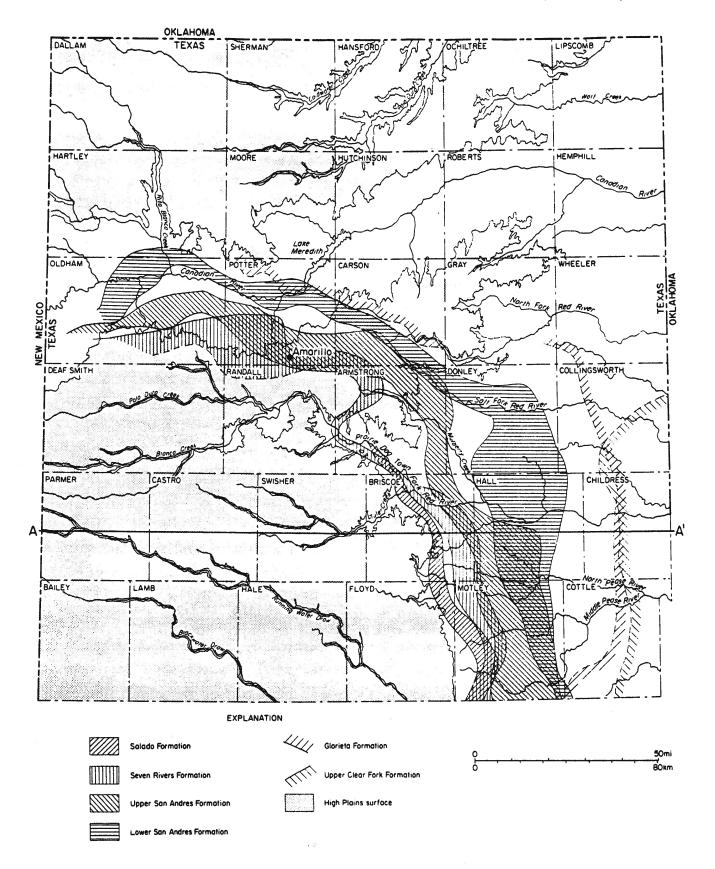
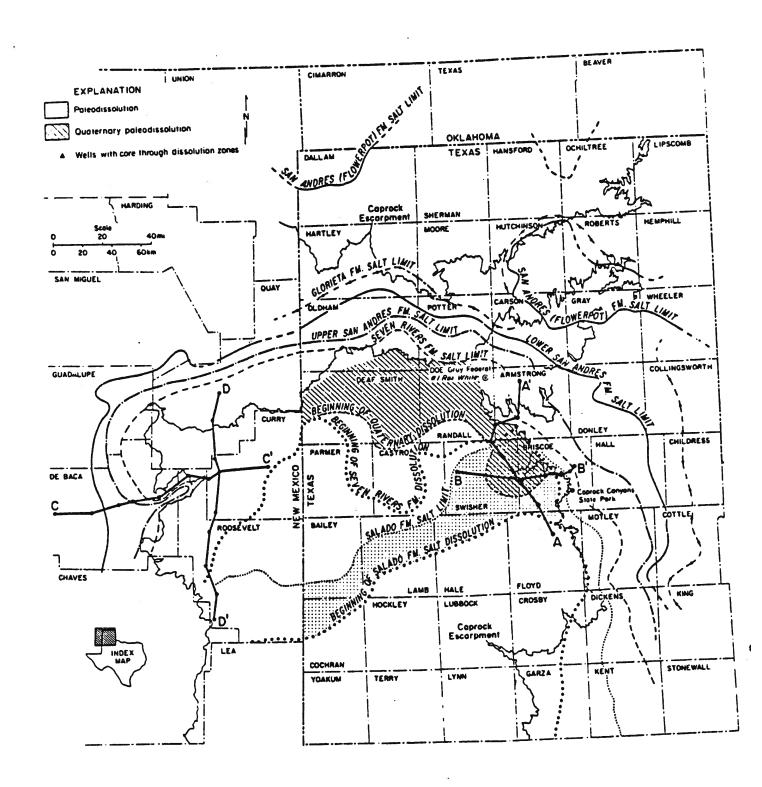


Figure 42. Salt dissolution zones, Texas Panhandle.



TAB 4 Part B.4 Subsurface geologic maps: lithology

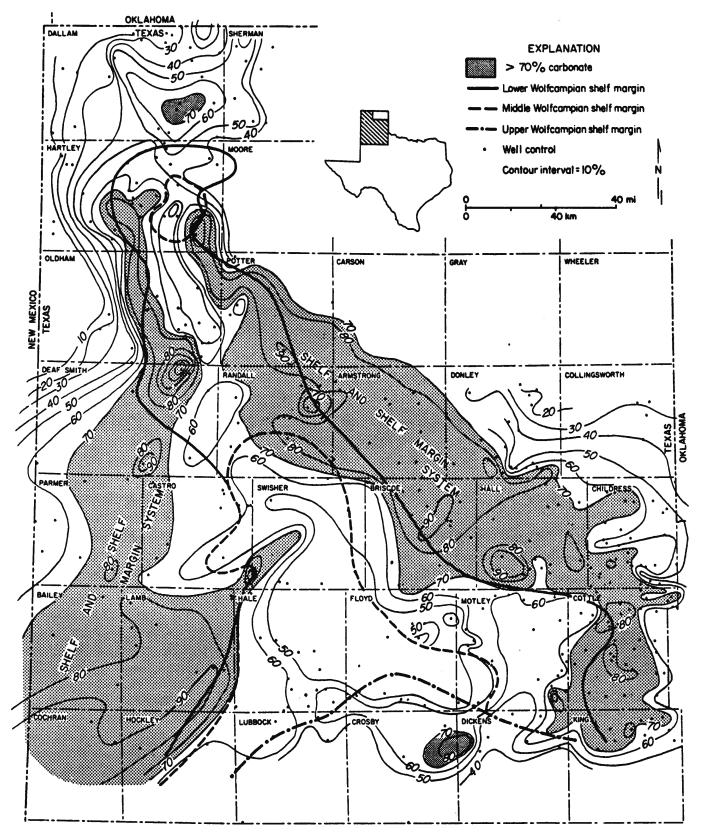


Figure 4. Percent carbonate map of Lower Permian Wolfcamp strata. Position of the shelf margin is shown by dashed line (from Handford, 1980).

Isotopic evidence for paleohydrologic evolution of ground-water flow paths, southern Great Plains, United States

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Bureau of Economic Geology, University of Texas, Austin, Texas 78713

RSTRACT

A confined aquifer in Triassic Dockum Group sandstone beneath the southern Great Plains was isolated from hypothesized paleorecharge areas in eastern New Mexico by Pleistocene erosion of the Pecos and Canadian river valleys and formation of hydrologic divides. Truncation of the flow system left meteoric water in the confined squifer with mean δD and $\delta^{18}O$ values that are $17^{\circ}/_{\circ \circ}$ and $2.0^{\circ}/_{\circ \circ}$ respectively, lighter than those in the overlying High Plains aquifer. Thick upper Dockum mudstone retards downward flow from the High Plains aquifer, which has been recharged by isotopically heavy precipitation during the Holocene. Recharge to the confined aquifer occurred at altitudes of 1600 to 2200 m in proximal Dockum sandsone facies since eroded in eastern New Mexico, at a mean temperature 3 °C cooler than present temperature across the southern High Plains. Effects of Pleistocene climatic change on isotopic composition of Dockum ground water could be superposed over geomorphologic effects.

INTRODUCTION

Conceptual hydrogeologic models of regional aquifers seldom consider whether paleohydrologic conditions are responsible for observed hydraulic-head and solute distributions. However, hydraulic-head distribution can gradually change during geologically long durations in response to physiographic evolution (England and Freeze, 1988). If ground water has a long residence time, stable isotopes and dissolved solutes can reflect a preceding flow regime even when hydraulic head is equilibrated with techarge and discharge rates.

The purpose of this paper is to discuss the significance of isotopically light δD and $\delta^{18}O$ compositions of ground water in the Triassic Dockum Group in the Texas Panhandle and eastern New Mexico (Fig. 1). Physiographic and climatic changes are obvious possible explanations of the sotopic depletion. The Pecos and Canadian river valleys were incised during the late Tertiary and Pleistocene (Gustavson and Finley, 1985; Gustavson, 1986). Isotopically light ground waters in other basins have been interpreted as Pleistocene in age and related to climatic variation (Clayton et al., 1966; Perry et al., 1982; Friedman, 1984; Rozanski, 1985). Whether the Pleistocene climate in the southwestern U.S. was generally cooler and wetter than the Holocene climate is uncertain (Barry, 1983; Friedman, 1983), particularly in the study area (Holliday, 1987).

HYDROGEOLOGIC SETTING

The Dockum Group is composed of more than 610 m of conglomersandstone, and mudstone deposited in fan, fluvial, deltaic, and lacusde environments in a continental basin (McGowen et al., 1979).

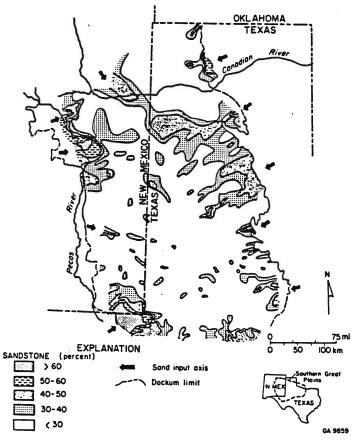


Figure 1. Percent sandstone in lower Dockum Group (modified from J. H. McGowen et al., unpublished map).

Sandstone and conglomerate beds in the lower Dockum Group are discontinuous and occur mainly around the basin perimeter (Fig. 1). In the muddy upper Dockum Group, which acts as a confining unit, thickest sandstone deposits are in southeastern New Mexico and the south-central part of the Texas Panhandle and are not extensive elsewhere (McGowen et al., 1979).

The essentially flat surface of the southern High Plains is inclined to the southeast and bounded by an erosional escarpment (Fig. 2) that has a relief of as much as 265 m. Beneath most of the southern High Plains, the Dockum is overlain by the Miocene Ogallala and Pleistocene Blackwater Draw Formations. Cretaceous Edwards and Trinity Groups subcrop beneath the younger formations in the southern part of the southern High Plains. Similarity of hydraulic head in the Ogallala Formation and Cretaceous rocks suggests that hydrostratigraphic units in these formations are interconnected in the so-called High Plains aquifer (Gutentag et al., 1984; Nativ and Smith, 1987). The Dockum Group overlies Permian evaporites

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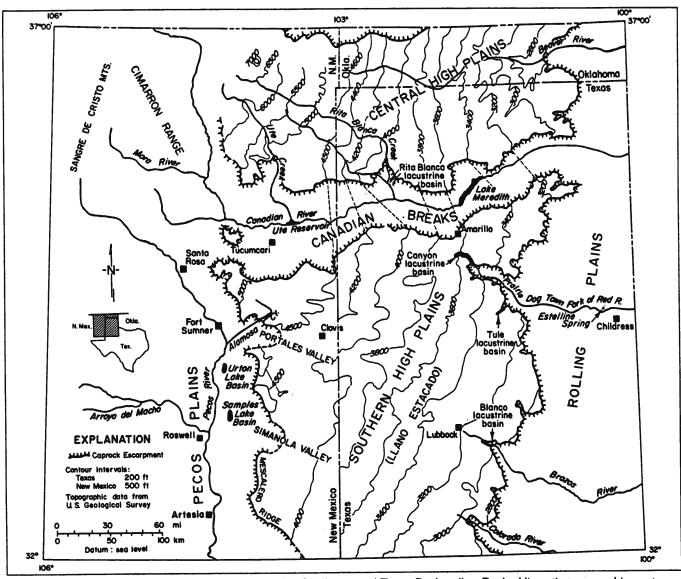


Figure 2. Physiography of eastern New Mexico and the Oklahoma and Texas Panhandles. Dashed lines tie topographic contours across the Canadian Breaks. If the strike of contour lines on the Southern High Plains is projected across the Canadian River valley, it becomes apparent that the northern side of the valley is approximately 80 m (250 ft) lower in elevation than the south rim of the valley.

PHYSIOGRAPHIC DEVELOPMENT OF EASTERN NEW MEXICO AND THE TEXAS PANHANDLE

Methods

The following discussion of the geomorphic evolution of eastern New Mexico and the Texas Panhandle is presented as case studies of drainage systems including the Pecos, Canadian, and Prairie Dog Town Fork Rivers, Tierra Blanca Creek, and other streams draining the Southern High Plains and the Rolling Plains. For each of these areas, field studies were integrated with subsurface studies based on interpretation of geophysical logs and

seismic reflection profiles, hydrologic and water quality data obtained from stream-gauging stations, interpretations of aerial photographs and Landsat images, and topographic maps.

Evidence of Dissolution

Regional salt dissolution and the subsequent collapse of overlying strata affected substantial parts of the Texas

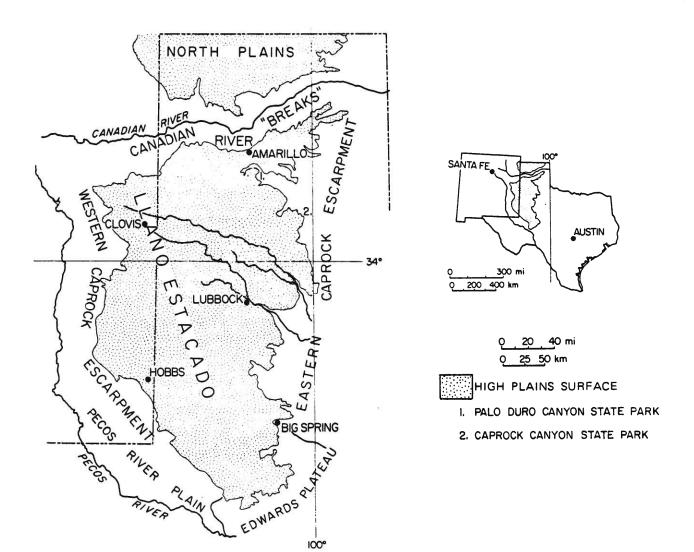


Figure 1. Physiographic units of the Texas Panhandle and eastern New Mexico.

River that is called the Canadian River Breaks. The Great Plains south of the Canadian River are also called the Southern High Plains or the Llano Estacado. To the east and west, the High Plains are truncated at the Caprock Escarpment, an erosional scarp where relief locally exceeds 500 m (1,500 ft). East of the escarpment lie the Rolling Plains. The High Plains of the Texas Panhandle are developed on the Tertiary Ogallala Formation, the remnants of a large alluvial apron that spread eastward as a result of uplift and erosion of the southern Rocky Mountains of New Mexico. Several tens of feet of Pleistocene eolian sediment cap the Ogallala in many areas. Drainage is poorly developed on the Southern High Plains and is mostly internal into the thousands of playas that cover its surface (Woodruff and others, 1979). Integrated drainage exists mainly as a series of extremely elongated rectilinear draws. The Caprock Escarpment is supported by a massive caliche horizon that marks the top of the Ogallala Formation and to a lesser extent by well-indurated sandstones that occur in the upper part of the Triassic Dockum Group. Eastward from the Caprock Escarpment, the Rolling Plains are developed on structurally disturbed Permian red beds. The Pecos River Valley and the Pecos River Plain lie westward of the Southern High Plains.

The major positive tectonic elements that surround the Anadarko, Dalhart, and Palo Duro Basins in the Texas Panhandle (fig. 2) have been described by Nicholson (1960) and Johnson (1976). Tectonic activity that created the series of arches, domes, and uplifts defining the basins occurred primarily during the Pennsylvanian Period and was largely completed by the end of that period. Minor movements have occurred since the Permian, but they may have resulted from differential compaction of basin sediments or from post-tectonic adjustments in the earth's crust.

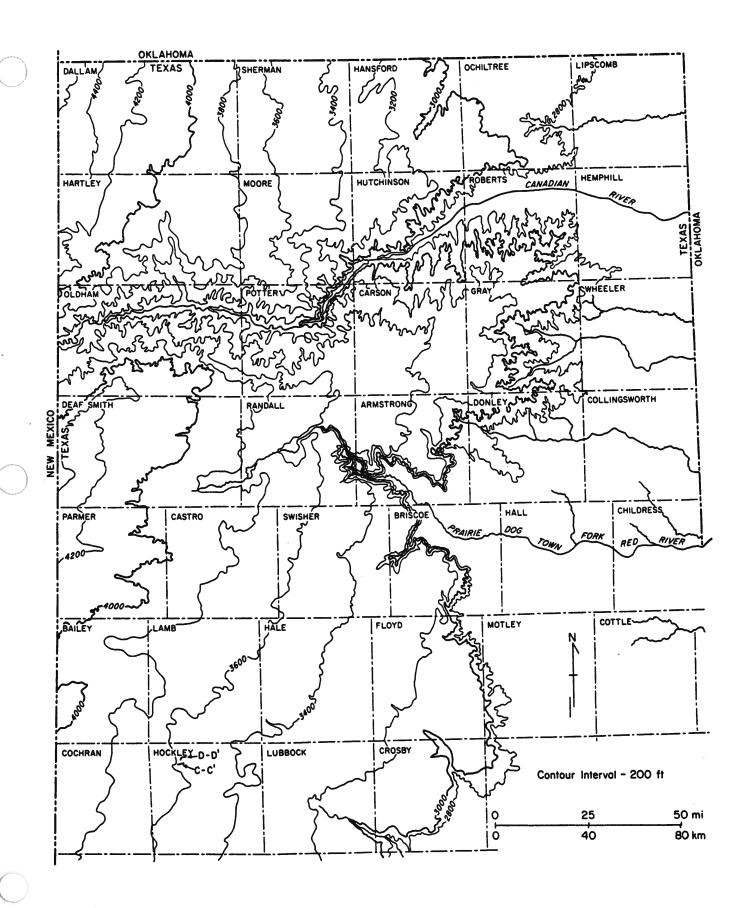


Figure 9. Topography of the High Plains, Texas Panhandle.

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Index to ${\sf TAB}$ 5: Geologic cross-sections of the study area

Gustavson and Finley, 1985, Figures 5 and 3

Gustavson, et al., 1992, Figure 2

Gustavson, et al., 1992, Figure 3

Gustavson, et al., 1980a, Figure 25

Gustavson, et al., 1980a, Figure 26

Gustavson, et al., 1980b, Figures 5 and 4

Gustavson, et al., 1980b, Figures 44 and 45

HGC, 1984a, Figures 13-15 and 12

HGC, 1984b, Figures 7, 8 and 6

SEO, 1961, Figure 6a

SEO, 1961, Index map to cross-sections, from reduced Figures 6b through 6m

SEO, 1961, Explanation of log symbols

SEO, 1961, Figures 7a through 7g

USBR, 1985, Figure II-1

USBR, 1979, Figure 4

USBR, 1979, Appendix C, Figure 2

USBR, 1979, Appendix D, unnumbered figure

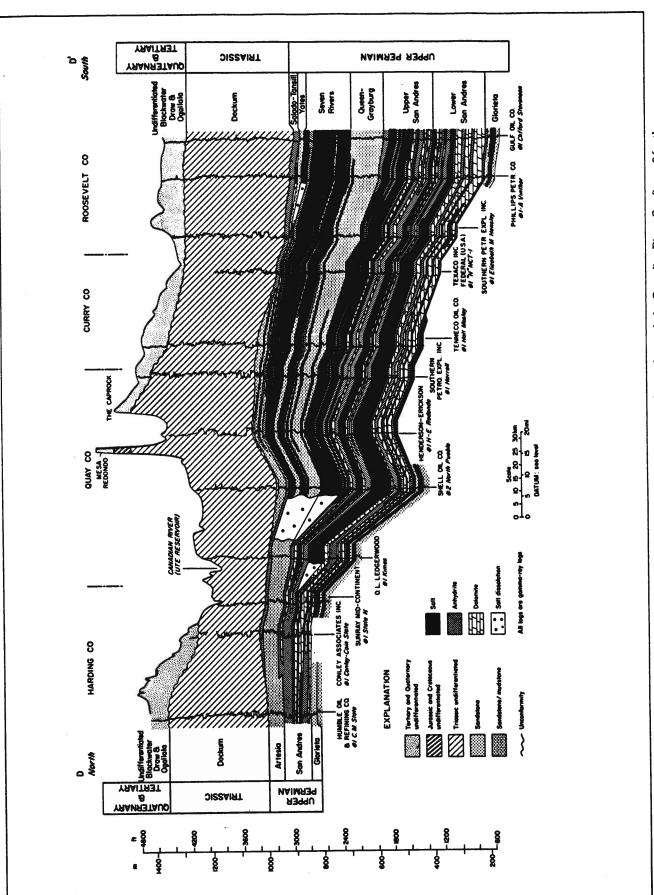


Figure 5. Stratigraphic cross section illustrating salt dissolution and collapse of strata beneath the Canadian River. See figure 3 for the location of section D-D'.

Figure 3. Zones of salt dissolution in eastern New Mexico and the Texas and Oklahoma Panhandles. Lines indicate the present extent of salt in the study region. In stratigraphic succession upward from the Glorieta to the Salado Formation, increasing amounts of salt are preserved toward the southwest corner of the Texas Panhandle. Some San Andres Formation salts are preserved in northwestern Dallam County, and some Glorieta and San Andres Formation salts are preserved near Hutchinson County. A peripheral dissolution zone is approximately located by salt-limit lines of the Seven Rivers, San Andres, and Glorieta Formations. An interior dissolution zone is approximated by the area overlain by the Southern High Plains. Wells with core through strata from which salt has been dissolved are indicated by numbered triangles: (1) DOE-Gruy Federal Rex H. White No. 1. (2) DOE-Gruy Federal Grabbe No. 1. (3) Stone and Webster Engineering Corp. Sawyer No. 1. (4) Stone and Webster Engineering Corp. Mansfield No. 1. (5) Stone and Webster Engineering Corp. Detten No. 1. (6) Stone and Webster Engineering Corp. G. Friemel No. 1. (7) Stone and Webster Engineering Corp. Zeeck No. 1. (8) Stone and Webster Engineering Corp. J. Friemel No. 1. (9) Stone and Webster Engineering Corp. Harman No. 1. (10) Stone and Webster Engineering Corp. Holtzclaw No. 1. Line A-A' is figure 7, line B-B' is figure 6, line C-C' is figure 4, and line D-D' is figure 5.

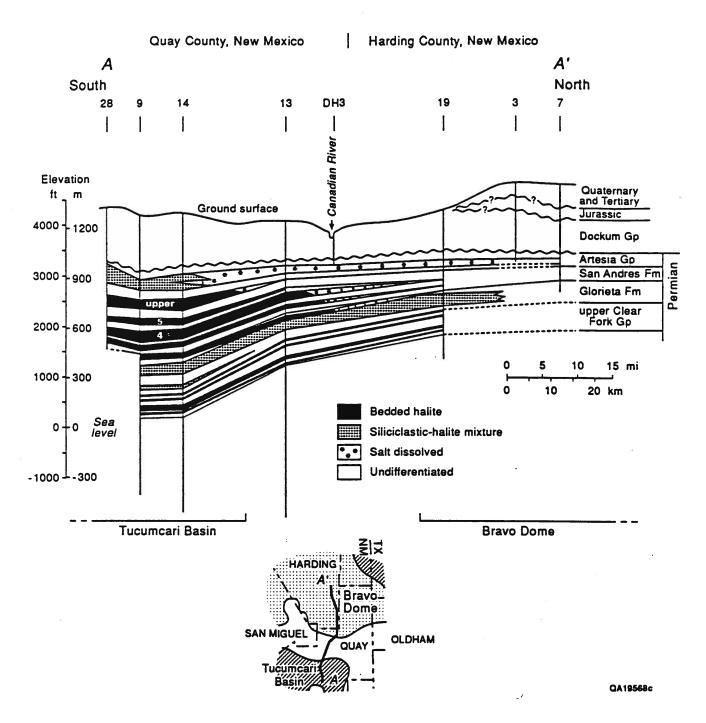


Figure 2. North-South structural cross section through the Ute Dam area.

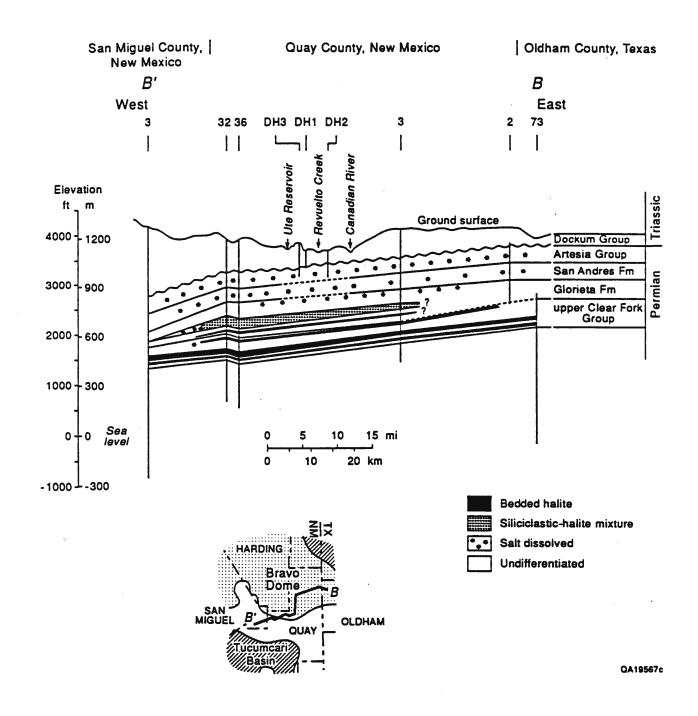


Figure 3. East-West structural cross section through the Ute Dam area.

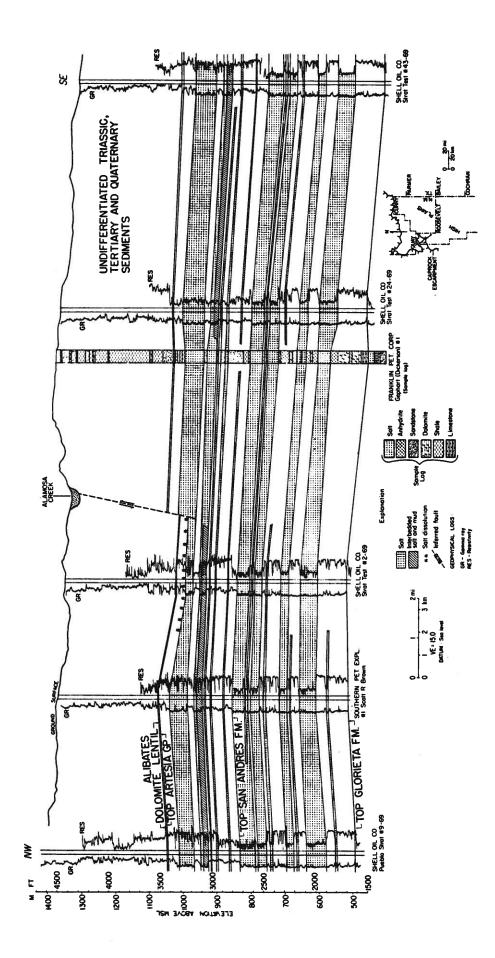
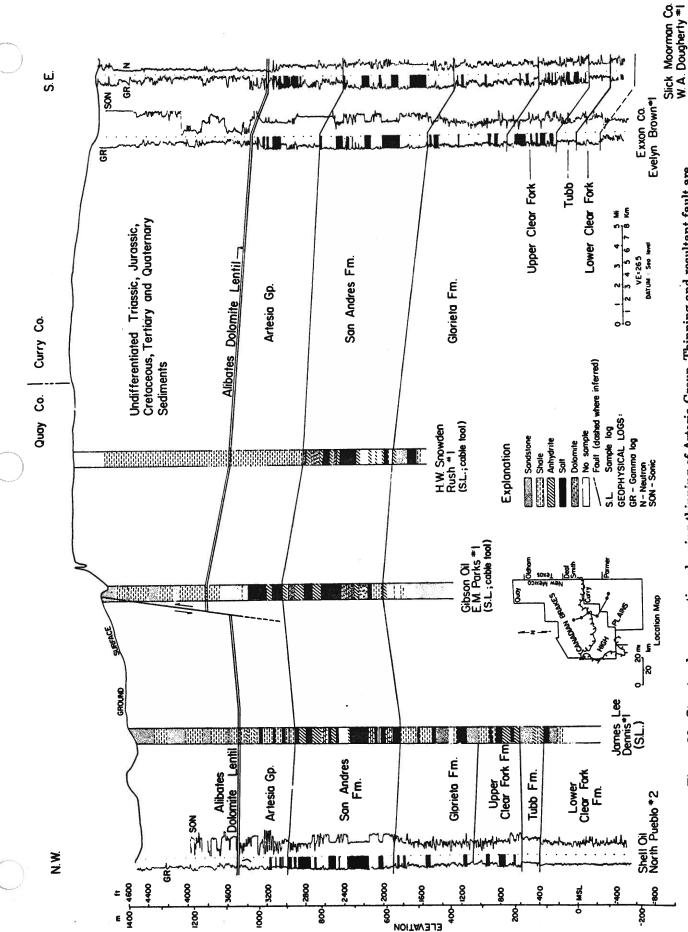


Figure 25. Structural cross section showing thinning of Artesia Group. Thinning and resultant faults are probably due to salt dissolution in the Artesia Group northwest of the fault. Cross section location is given in the inset map in this figure.



probably due to salt dissolution in Artesia northwest of the fault. Cross section location is given in the inset Figure 26. Structural cross section showing thinning of Artesia Group. Thinning and resultant fault are map in this figure.

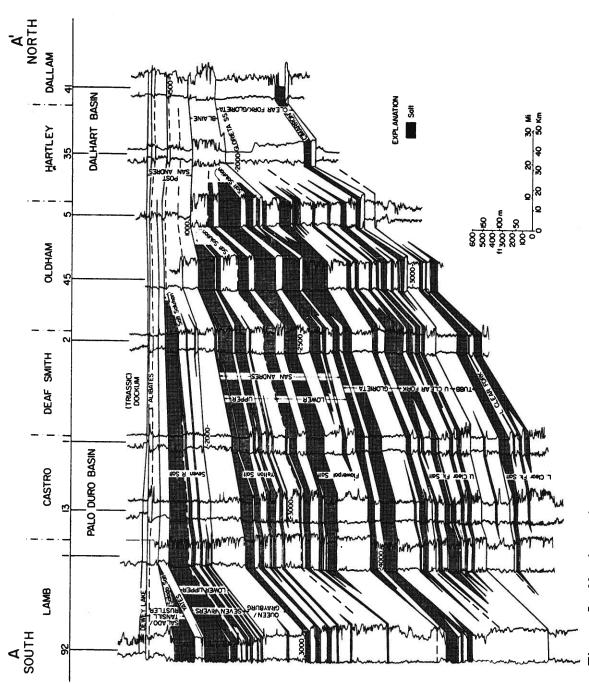


Figure 5. North-south cross section, Upper Permian salt-bearing strata, Texas Panhandle. Generalized salt units are correlated. Location of section in figure 4.

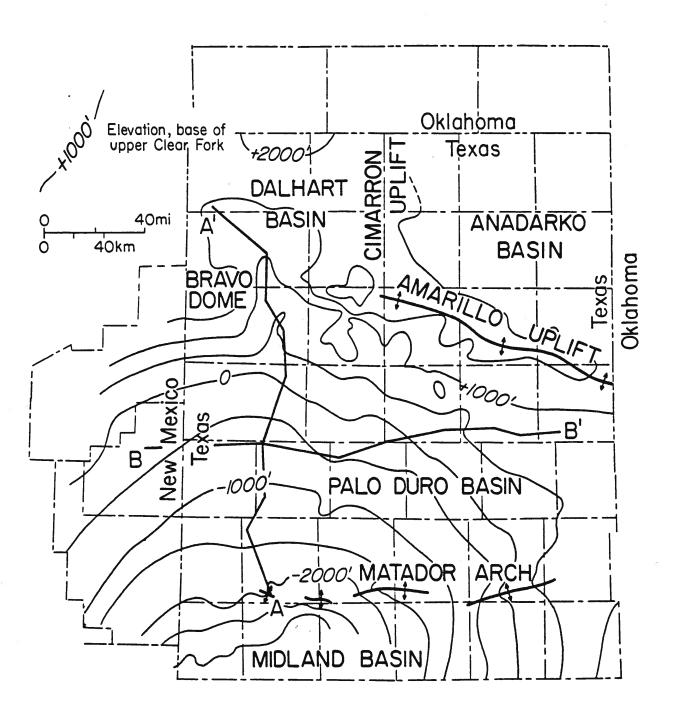


Figure 4. Regional structural setting of the Palo Duro and Dalhart Basins.

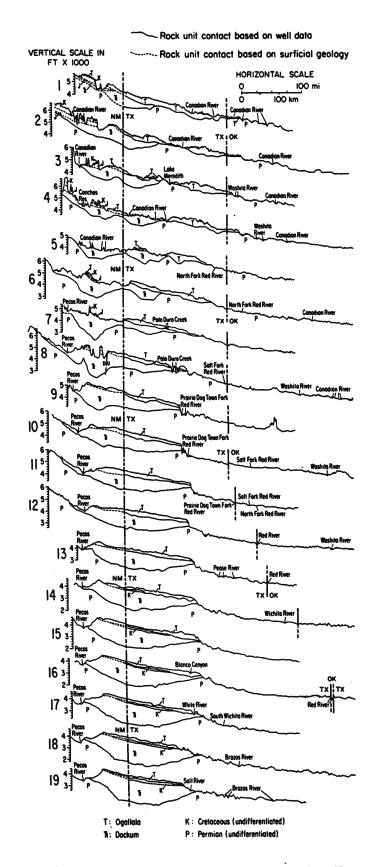


Figure 44. Topographic and geological profiles across the Texas Panhandle and adjacent parts of New Mexico and Oklahoma.

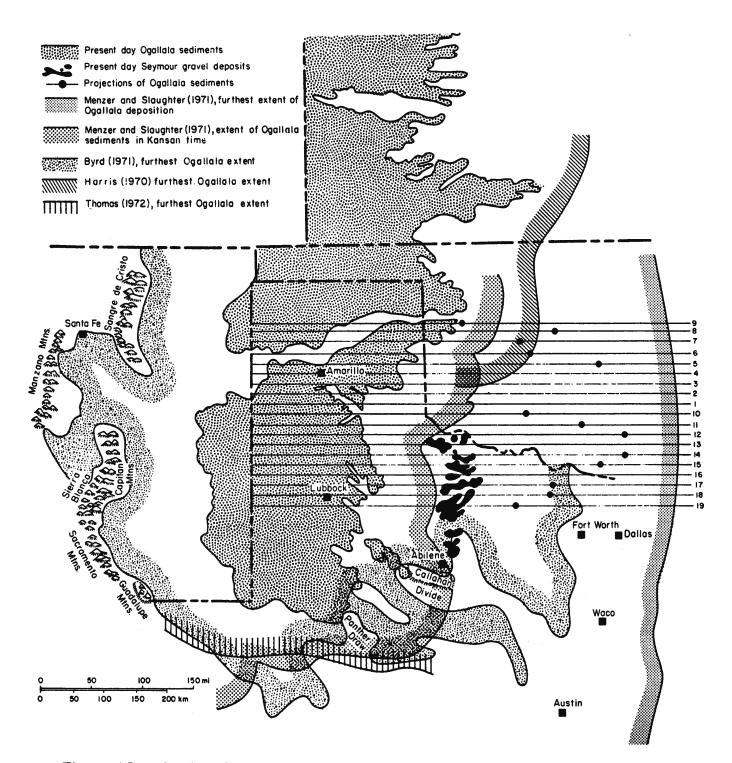


Figure 45. Combined interpretations of the easternmost extent of the Ogallala Formation (Harris, 1970; Menzer and Slaughter, 1971; Byrd, 1971; Thomas, 1972). Position of topographic profiles (9-19) in figure 44 are shown. Black dots show points of intersection of projected Southern High Plains surface, with the surface of the Osage Plains in Oklahoma.

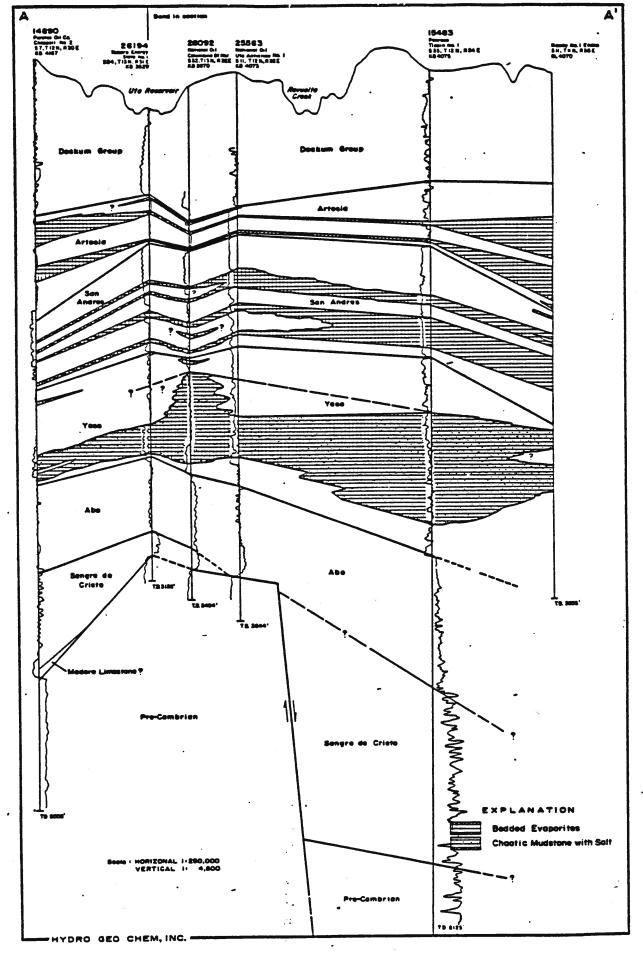


Figure 13. East-West geologic section

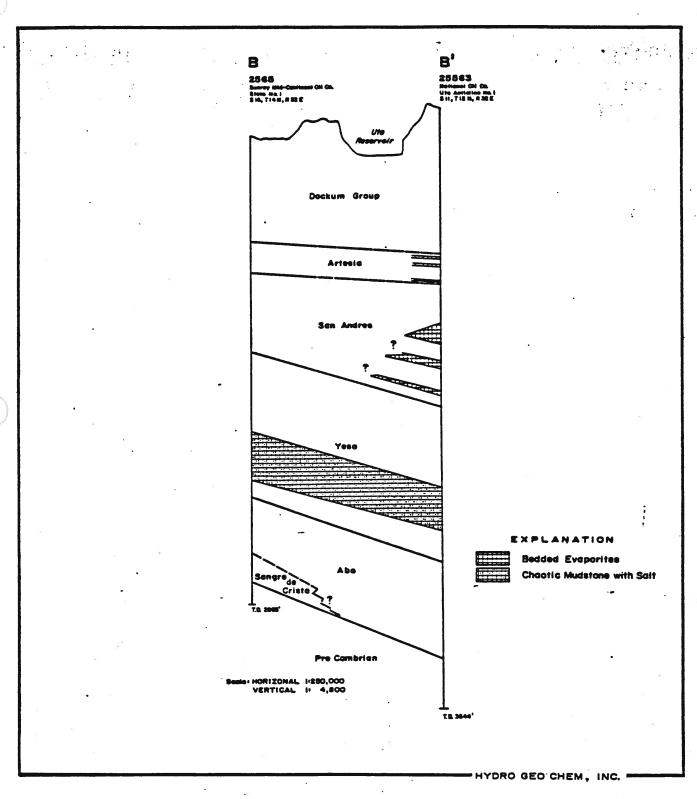


Figure 14. North-South geologic section through Ute Reservoir

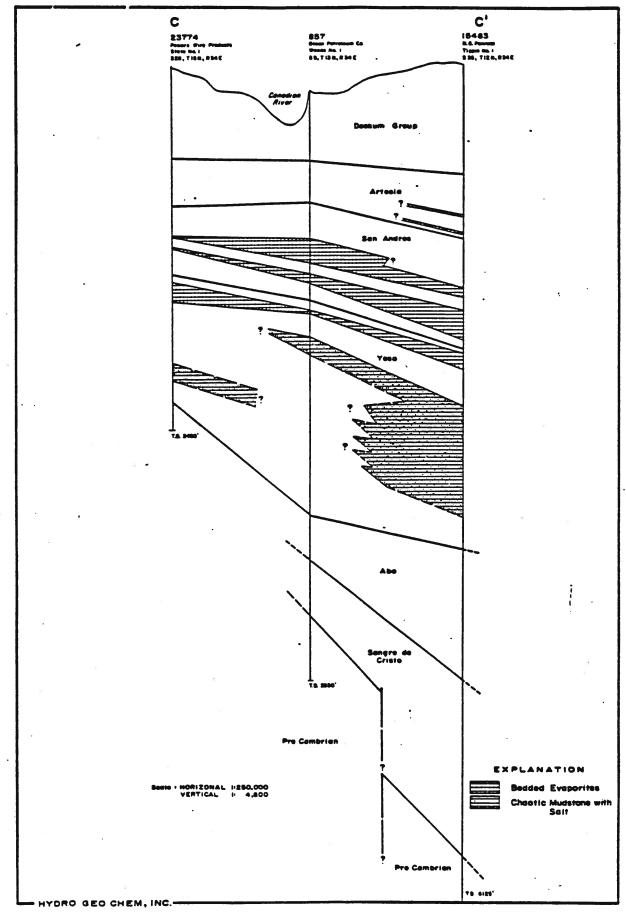


Figure 15. North-South geologic section through Canadian River

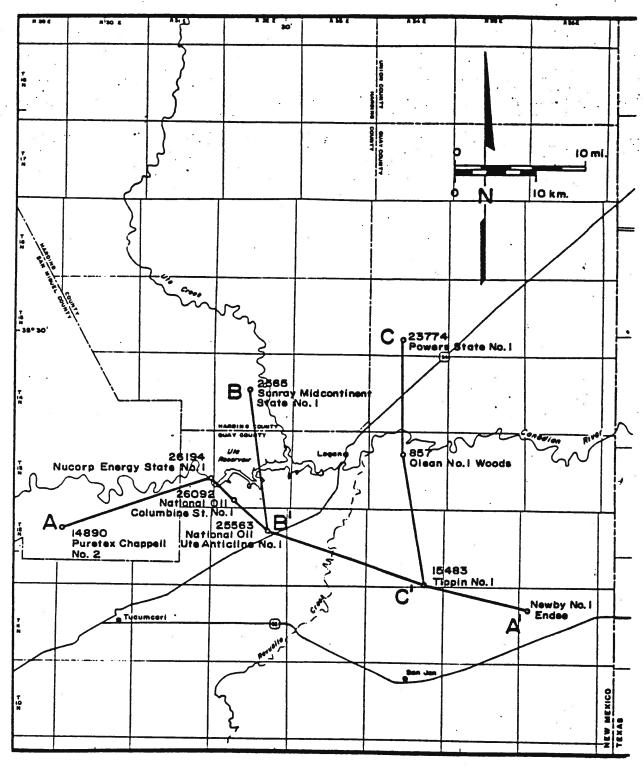
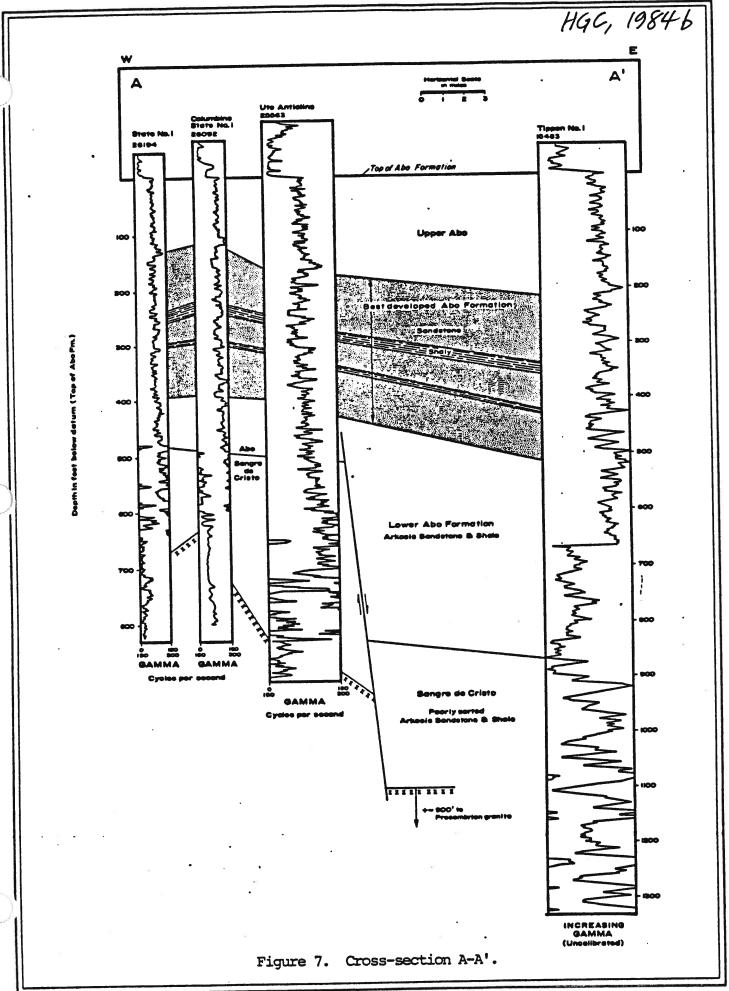
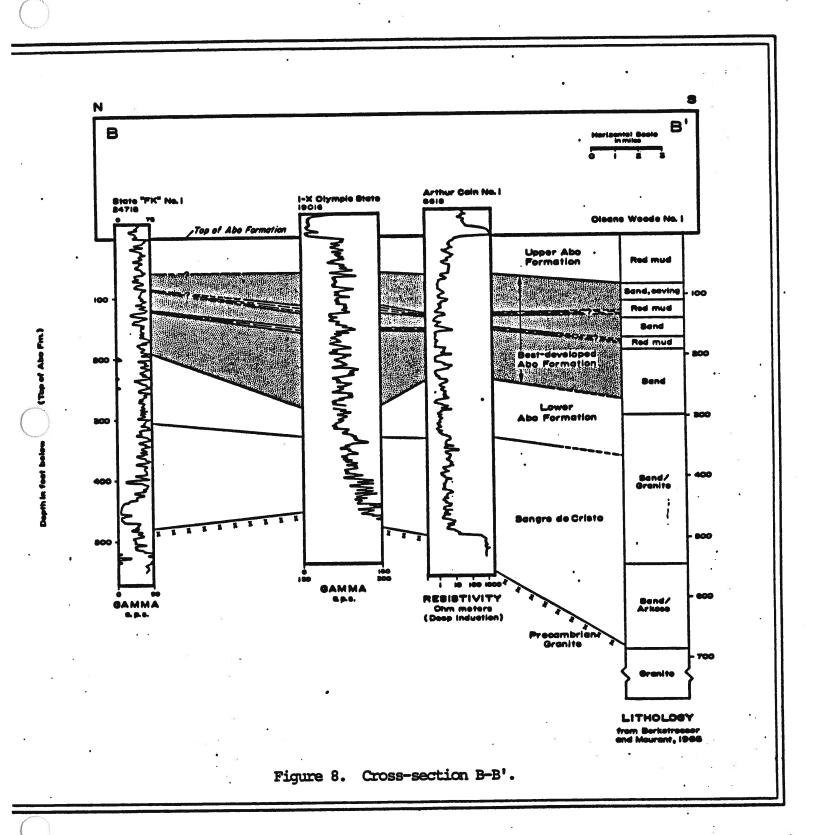


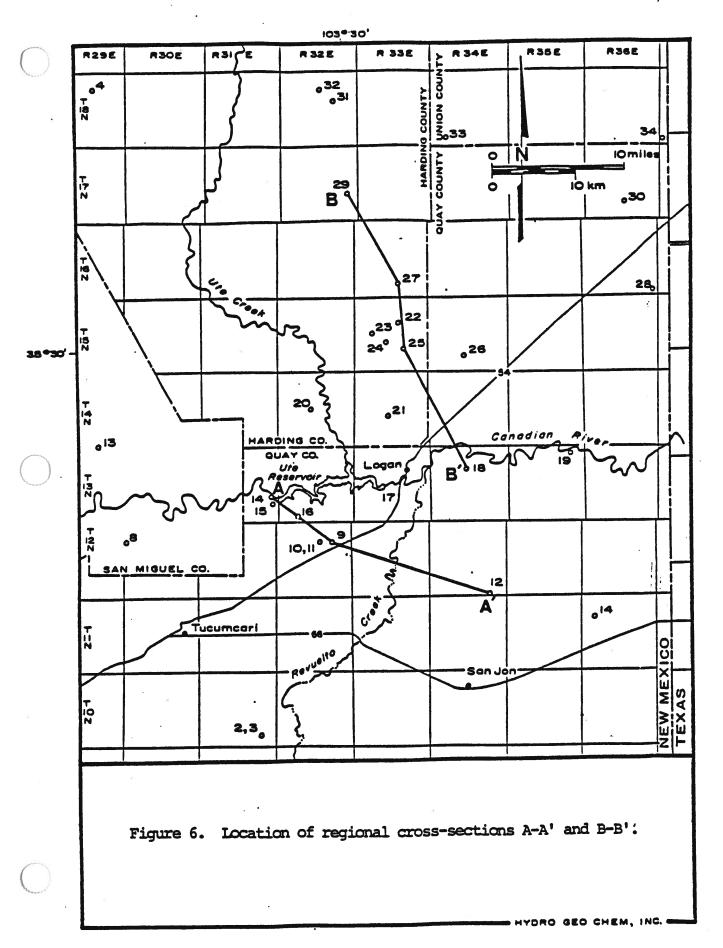
Figure 12. Map showing locations of cross-sections

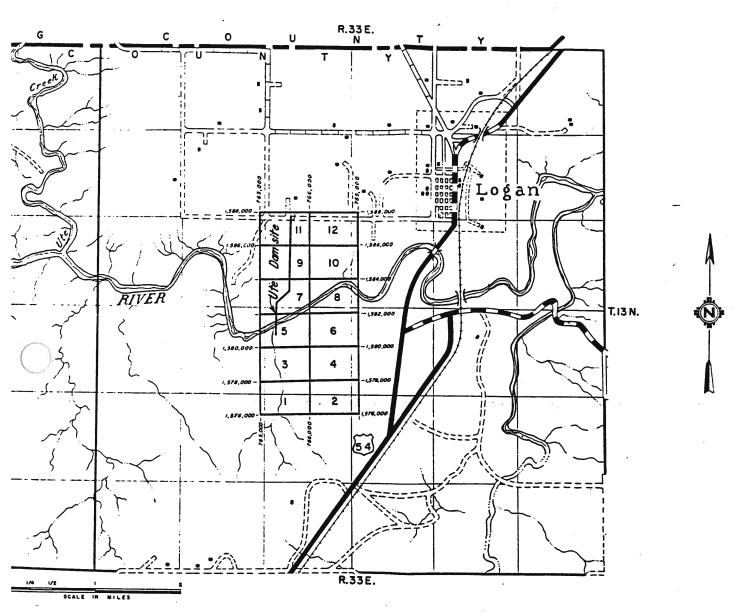
EXPLANATION

O 14890
Puretex Chappell
No. 2
Well number
Operator, name









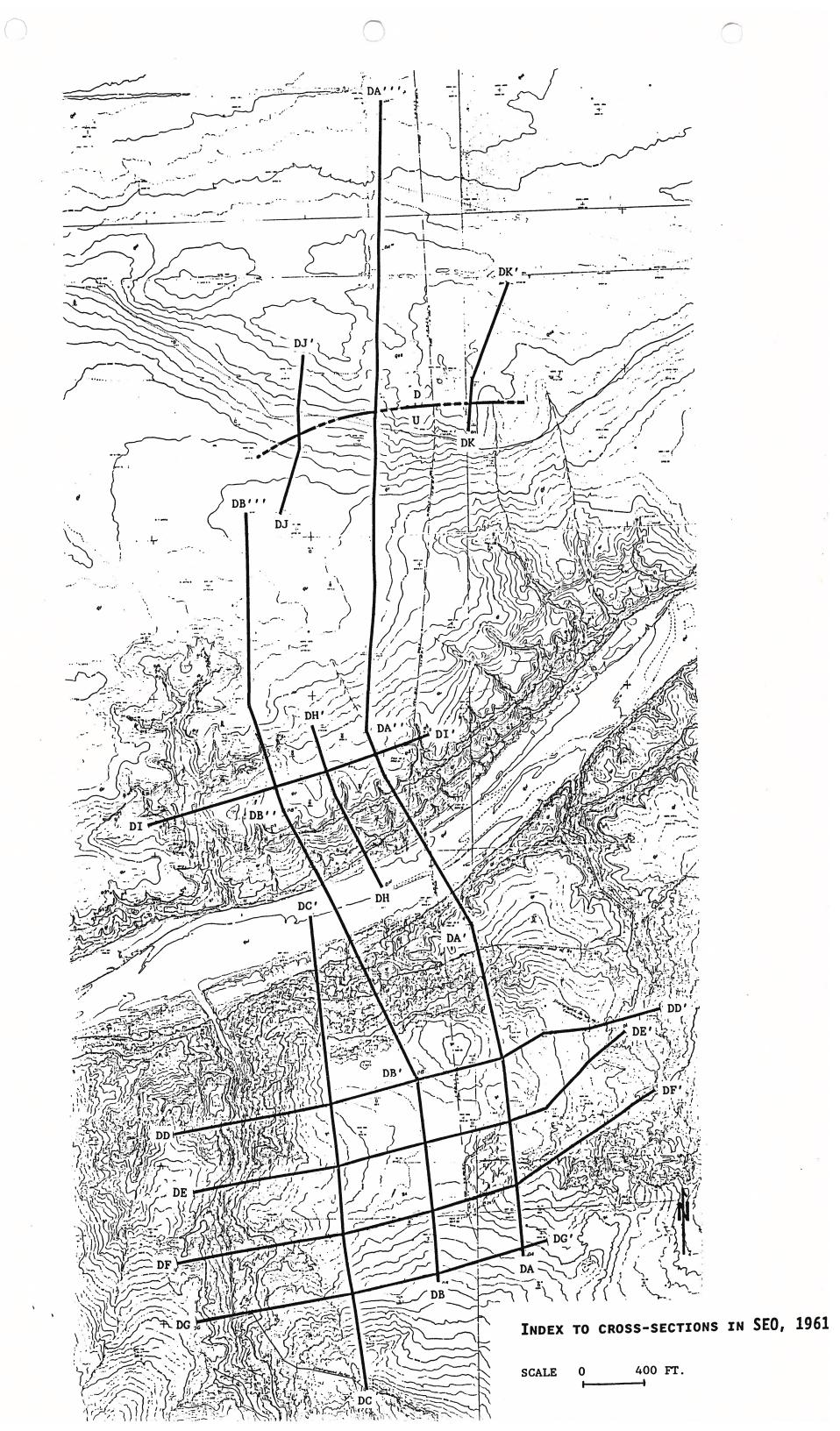
EXPLANATION

TRIANGULATION POINT (BRASS CAP) TRIANGULATION POINT (PIPE) 1-26 HOLE NUMBER; E AT HOLE REFERENCE MARKER YT, NUMBER INDICATES PERMANENT BENCH MARK I LEFT, CENTER OR IS IDENTIFIED WITH THE SHEET NUMBER AND LOCATION OF UTE KELSH ELEVATION OF HIGHEST POINT

T-X

IN VICINITY

FIGURE 6a



EXPLANATION OF LOG SYMBOLS (Figures 7a through 7g)

DH - 4 Elev. 3778.5 Exploratory drill-hole number and elevation of land surface at hole

RA - 20 Elev. 3764 Reconnaissance auger-hole number and elevation of land surface at hole

Grvl

Gravel

Bldrs

Boulders

Sd & Grvl

Sand and gravel

Congl

Conglomerate

Soil

Soil

Sd

Sand

Slt

Silt

Clay

Clay

Ss

Sandstone

Slts

Siltstone

Cs

Claystone

Ms

Mudstone

Sh

Shale

Caliche

Caliche

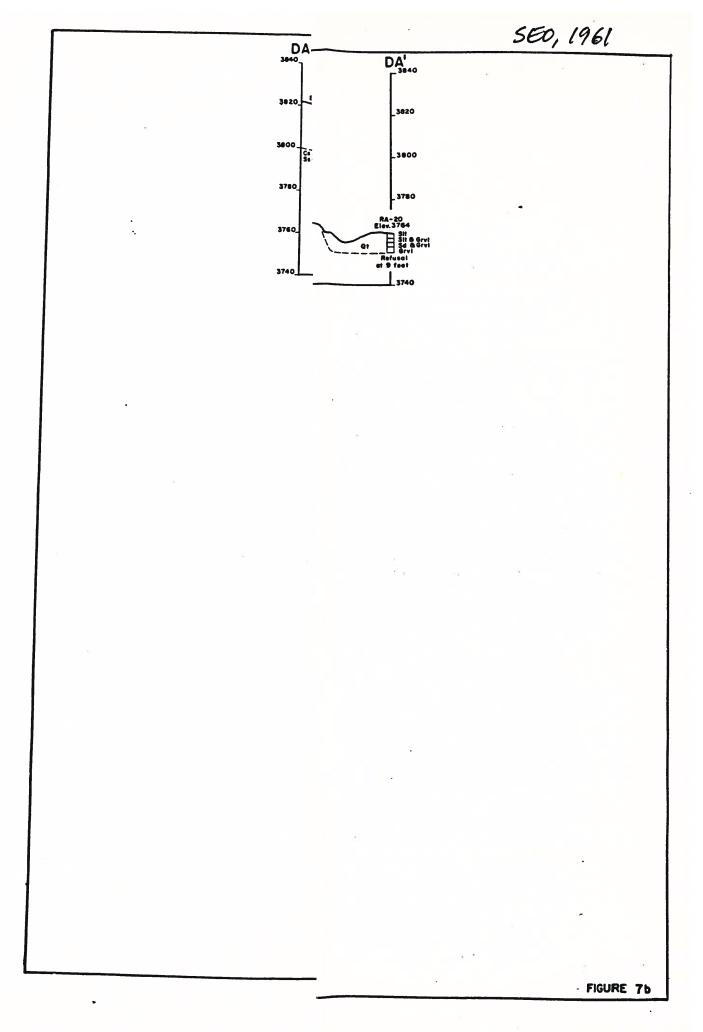
Ls

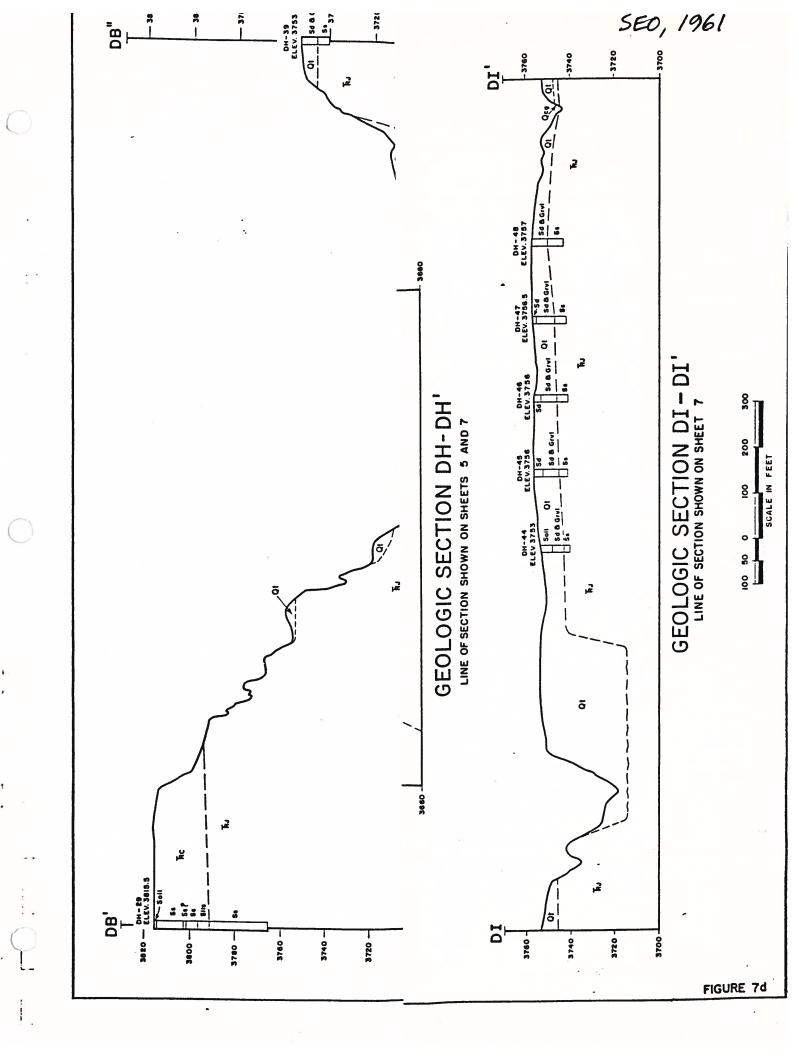
Limestone

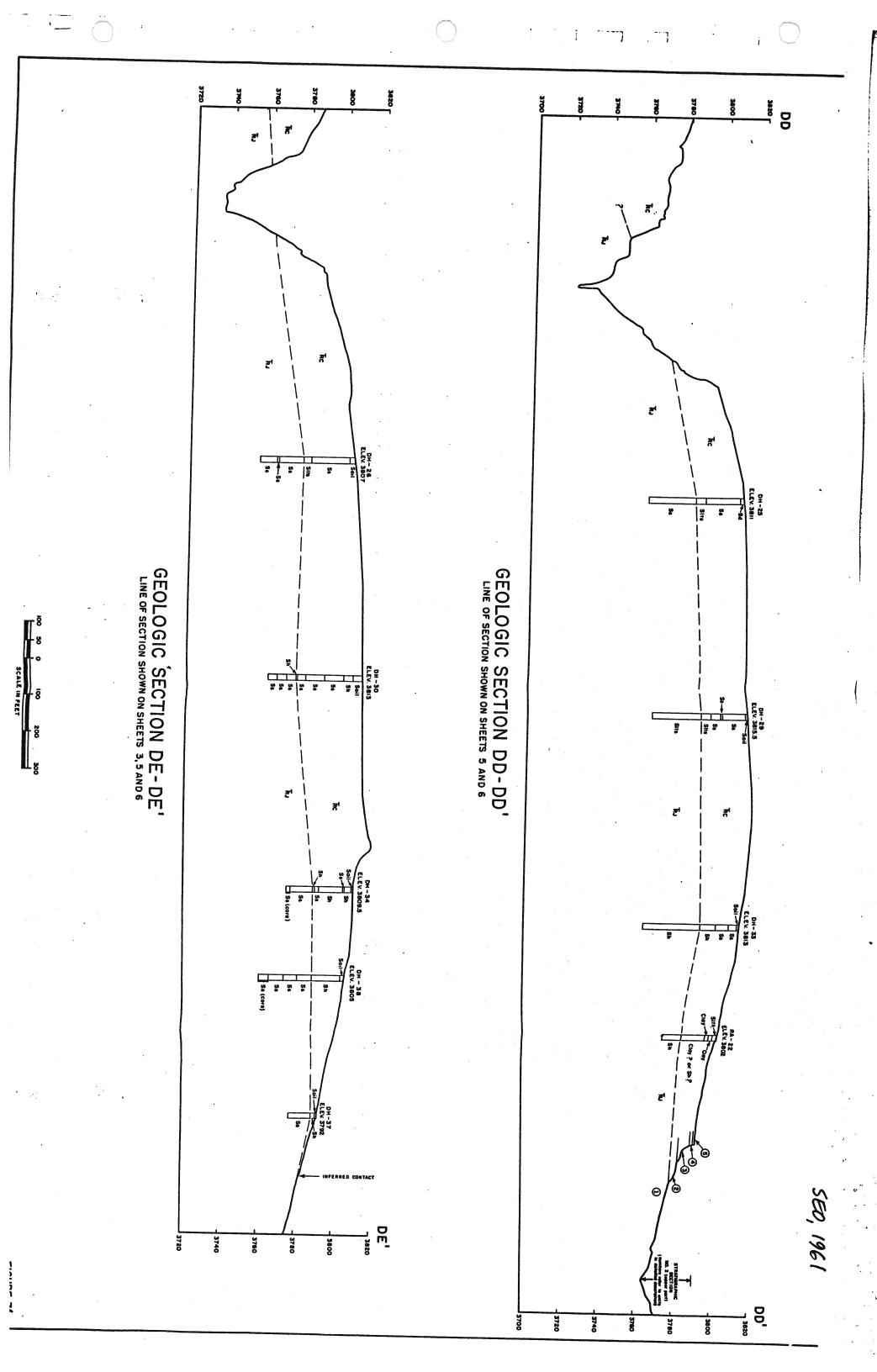
Cored interval

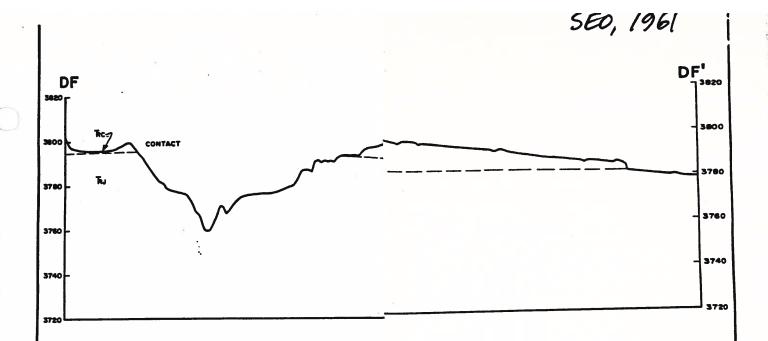
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CANADIAN RIVER INVESTIGATION
GEOLOGIC SECTIONS - UTE DAM SITE

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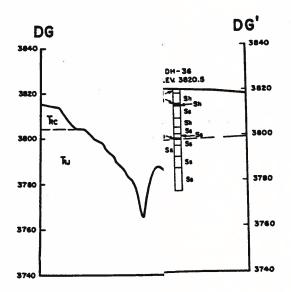
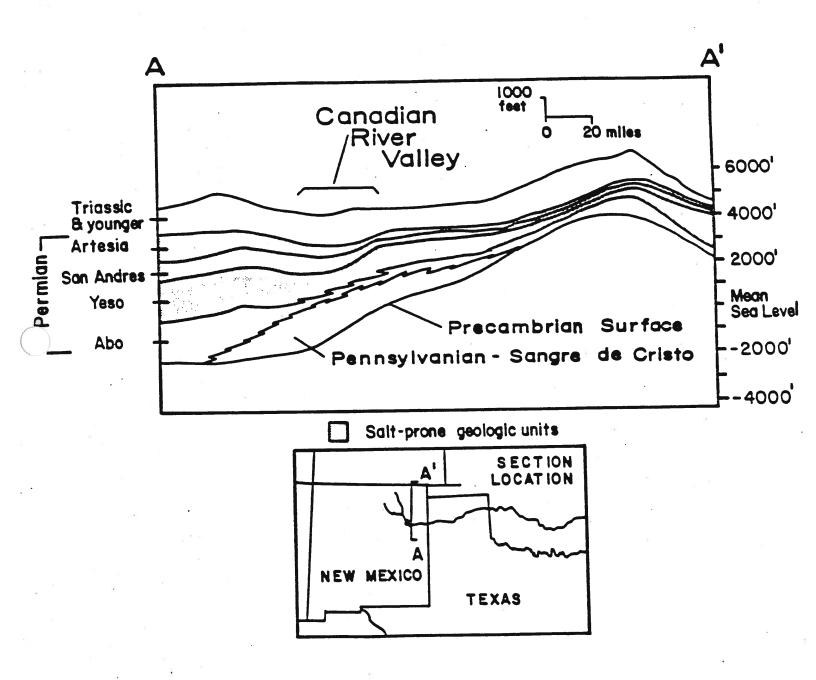
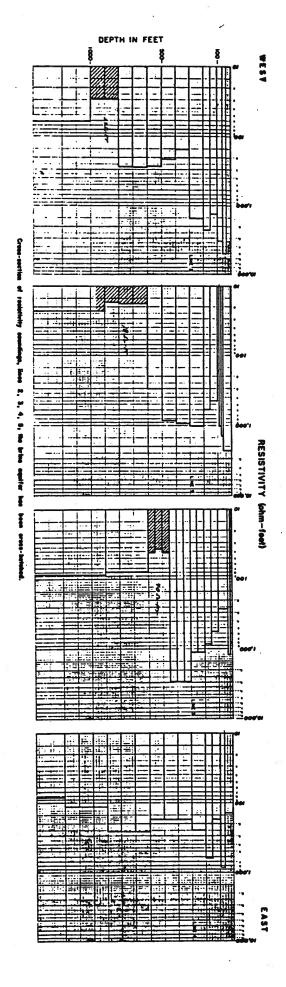


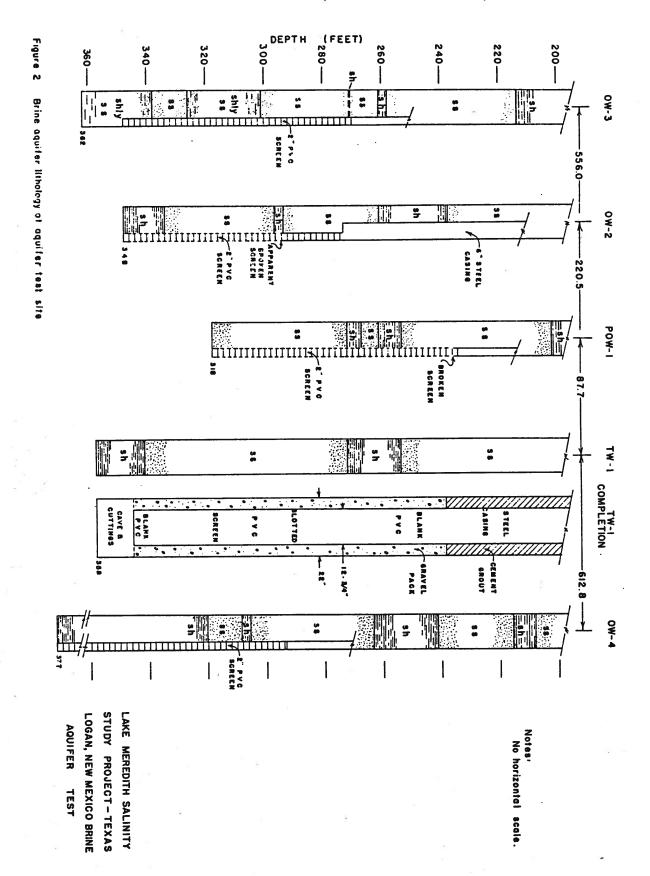
FIGURE 79



SUBSURFACE GEOLOGIC FORMATIONS Logan, New Mexico Area



FIGURE



USBR, 1979 Appendix D

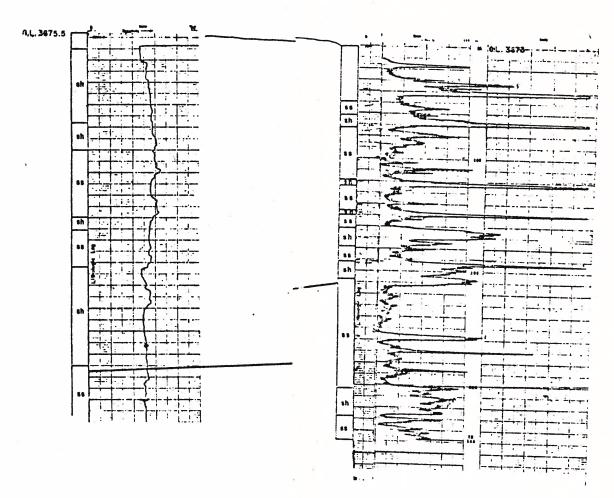
OW NO. 4

STAPP-HAMILTON ASSOC. Quay County, New Mexico

Date First Reading Last Feating 3:1 Size Resistivity Scale Recording Speed Resistivity 12-13-77 326 0 4 3/4 20 ohm m/m² 30 FPM OW NO. 3

STAPP-HAMILTON ASSOC.

| Commo Roy | Comm



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LAKE MEHEDITH SALINITY STUDY-TEXAS, N.M.

GEOLOGIC SECTION LOGAN AREA

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Index to TAB 6: Hydrostratigraphic columns of the study area

Bassett and Bentley, 1983, Table 1

Senger, et al., 1987, Table 1

Wirojanagud, et al., 1986, Table 2

Wirojanagud, et al., 1986, Table 3

Table 1. Generalized stratigraphic column, depositional environment, and general hydrologic properties, Palo Duro Basin.

System	Series	Group	General lithology and depositional setting	Hydrogeologic element	Hydrogeologic unit	Approximate permeability (gpd/ft²)	Relative permeability
Quaternary			Fluvial and				
Tertiary	~~~	~~~	locustrine clastics	Ogallala - aquifer	Upper	₂₀₀ @	108
Cretaceous	~~~	~~~~	Nearshore marine clastics		aquifer		
Triassic	~~~	Dockum	Fluvial deltaic and lacustrine clastics and limestones	Dockum aquifer		2-50(3)®	10 ⁶ -10 ⁷ (?)
~~~	Ochoa	~~~					
	Guadalupe	Artesia	Salt, anhydrite,		Evaporite	.0	
Permian		Pease River	red beds and peritidal dolomite	Evaporite aquitard	aquitard	2×10-6 ©	'
rermian	Leonard	Clear Fork					
		Wichita					
	Wolfcamp			Wolfcamp carbonate aquifer	Deep-basin		,
Pennsylvanian			Shelf and platform carbonates, basin shale and deltaic sandstones	Pennsylvanian carbonate aquifer Upper	brine aquifer		
			Basin shale aquitard	Paleozoic granite wash aquifer	>	2 x 10 - 2 ⁽¹⁾ Basin le aquitard	104
Mississippian			Shelf limestone and chert	Lower Paleozoic	I \		
Ordovicion		Ellenburger	Shelf dolomite	carbonate aquifer			
Cambrian			Shallow marine (?) sandstone	Lower Paleozoic sandstone aquifer			
~~~	Precombria	<u> </u>	Igneous and metamorphic	Basement aquiclude	Basement aquiclude	0	0
_ <u>_</u>		960), Myers (en communicat	1969) , Cronin (1961)				

- **(b)** Stevens (1980, written communication
- © Geotechnical Engineers (1978)
- This study

permeabilities (for further discussion, see p. 10); therefore all pre-Leonardian formations were grouped together into a single hydrogeologic unit for this preliminary regional analysis.

Middle and Upper Permian strata consist almost entirely of halite, anhydrite, dolomite, and fine-grained siliciclastic red beds, which grade southward into shallow-marine carbonates in the Midland Basin (Dutton and others, 1979). Together these formations compose the evaporite aquitard (table 1).

Overlying the Permian evaporites and red beds are the fluvial, deltaic, and lacustrine deposits of the Triassic Dockum Group and alluvial deposits of the Tertiary Ogallala Formation (fig. 5; table 1). The Dockum Group records the final stages of filling of the Permian Basin (McGowen and

others, 1979). Hydrogeologic information on Dockum sandstones is limited; wells over the basin tapping these beds have low specific capacities and produce waters that range widely in salinity. In contrast, potable ground water in the overlying Ogallala aquifer has been heavily pumped for agricultural, industrial, and domestic purposes. The Ogallala is an extensive alluvial apron of sand, gravel, and clay that extends eastward from the Rocky Mountains in the form of coalescing alluvial fan lobes (Seni, 1980). The upper part of the Ogallala Formation is cemented with calcium carbonate or "caliche" that forms the resistant "caprock" rim of the Caprock Escarpment along the eastern boundary of the High Plains (fig. 5).

Shallow, fresh ground waters generally move eastward under the influence of the regional

TABLE 1. Generalized stratigraphic column of the Palo Duro Basin (modified from Bassett and Bentley, 1982).

Era	System	Series	Group	Formation	General lithology and depositional setting	Hydrogeologic element	Hydrogeologic unit
Ö	Quaternary			Blackwater Draw	Eolian sand and silt		
Cenozoic	Tertiary	Pliocene- Miocene		Blanco Ogallala	Eolian, fluvial, and lacustrine clastics	Ogallala aquifer	Shallow
Mesozoic	Cretaceous				Nearshore nonmarine and marine clastics, carbonates		aquifer
Mess	Triassic		Dockum		Fluvial-deltaic and lacustrine clastics	Dockum aquifer	
		Ochoan					
		Guadalupian	Artesia			Evaporite aquitard	
		Ouadaidpiair	Pease River	San Andres	Cyclic sequences: shallow marine carbonates.	San Andres unit 4	
				Glorieta	hypersaline shelf,	carbonate aquifer	Evaporite aquitard
	Permian			Cimarron anhydrite	anhydrite, halite, and continental red beds		
		Leonardian	Clear Fork	"Tubb zone"	Continental rea deas	Evaporite aquitard	•
	-			Red Cave			
			Wichita				
		Wolfcampian		"Brown Dolomite"	a 5	Wolfcamp	
ZOİC					Shelf and shelf-margin carbonates,	carbonate aquifer	
Paleozoic	Pennsyl-				basinal shale, and deltaic sandstones	Pennsylvanian carbonate aquifer	
	vanian				Basinal-shale aquitard	Upper Paleozoic granite-wash aquifer	Deep-Basin Brine aquifer
	Mississippian				Terrigenous clastics, shelf carbonates, and chert	Lower and Middle	
	Ordovician		Ellenburger		Shelf dolomite	Paleozoic carbonate aquifer	
	Cambrian			•	Shallow marine (?) sandstone	Lower Paleozoic sandstone aquifer	
Pre- cambrian					Igneous and metamorphic	Basement aquiclude	Basement aquiclude

consequently, observed hydraulic heads may reflect some past paleohydrologic state of the system. Further, reservoir-pressure decline from production of oil and gas fields along the margins of the basin may have caused large-scale transient underpressuring. The role of these various processes and their impact on the hydrodynamics of the basin were addressed in this study.

Scope

A two-dimensional ground-water flow model was constructed along a cross section through the Palo Duro Basin to characterize regional ground-water flow paths as well as to investigate causes of underpressuring below the Evaporite aquitard, to evaluate mechanisms of recharge and discharge to and from the Deep-Basin Brine aquifer, and to examine transient effects of erosion and hydrocarbon production. This study was designed to investigate various factors affecting the overall groundwater flow pattern in the basin and was not necessarily aimed at producing a fully calibrated predictive model.

In the first phase, the model was used to simulate steady-state ground-water flow conditions using data on hydraulic conductivity from various hydrologic units in the section and hydraulic heads and recharge rates along the boundaries of the model. Objectives were to evaluate the effects of hydrostratigraphy and topography on the

		y = 1	ln(k)	Geometric	Number and	Typical
Hydro	geologic unit	Augusta md		source of data	value, md	
Evar	porite strata	-	-	-	-	.00028 (vertical permeability)
	Wolfcamp carbonate	2.19	2.89	8.90	25 - DST data 70 - TWDB core data 6 - Sawyer No. 1 pumping-test data	.07-300*
Deep- Basin Brine	Pennsylvanian carbonate	2.88	3.73	17.90	25 - DST data 118 - TWDB core data	.07 500
aquifer	Shale	-		-	-	.0000108*
at a	Granite wash	1.27 (2.15 with- out Mobeetie data)	6.17 (4.15 without Mobeetie data)	3.55 (8.60 without Mobeetie data)	10 - DST data 10 - Sawyer No. 1 pumping-test data 415 - Mobeetie field core data 11 - TWDB core data	.01-380*
¥	Pre-Pennsylvan- ian rock	1.56	2.87	4.76	11 - DST data 14 - Sawyer No. 1 pumping-test data	-

^{*} From Davis and DeWiest (1966), Freeze and Cherry (1979), and Davis (1980).

Note: (1) 1 md = 0.00115 m/day for saline water having salt concentration of 127,000 mg/L at 115°F.

(2) DST = drill-stem test; TWDB = Texas Water Development Board (Core Laboratories, 1972).

10 DST, 10 pumping tests in a single granite-wash interval in the Sawyer No. 1 test well, and 426 laboratory core analyses. Of the core sample analyses, 415 are from 6 wells in the Mobeetie field in the Anadarko Basin. Pre-Pennsylvanian permeability data are limited and consist of values from 4 DST of the Ellenburger Group, 6 DST of Mississippian carbonates, 1 pumping test of the Ellenburger Group, and 14 pumping tests in a single Mississippian carbonate interval in the Sawyer No. 1 test well. From this data base, Smith (1983) summarized the permeability values of each hydrogeologic unit and computed the geometric mean, arithmetic mean, and variance of the permeability for each type of data. Additional permeability data from five pumping tests in the Pennsylvanian granite wash at the Stone and Webster J. Friemel No. 1 well in Deaf Smith County, Texas, indicate a permeability range of 10 to 400 md and an average of 140 md. Laboratory

tests on a granite-wash core sample from the same well indicate a permeability range of 97 to 267 md.

It should be noted that none of the aforementioned permeability data represent a vertically averaged permeability of the hydrogeologic unit at a given location, which is the desired nodal point value in two-dimensional areal flow simulations. Although pumping tests give permeability values that represent the average fluidconducting property of a larger volume of the medium than do permeability tests for core samples, the tested zone of the medium is only a small part of the entire thickness of the hydrogeologic unit. No attempt was made to compute the vertically averaged permeability at data points having more than one permeability value because of the insufficiency of information and the variety of testing techniques used to obtain the permeability data. Instead, all the permeability data on each hydrogeologic unit (including those of the neighboring basins) were used to compute the unit's geometric mean and variance.

Table 2 summarizes the effective permeability values and variances of each hydrogeologic unit of the Palo Duro Basin. The large variance values indicate a large natural variation in the permeability of each hydrogeologic unit. By including permeability data from neighboring basins, the effective-average permeability value is slightly increased for the Wolfcamp and Pennsylvanian carbonates but slightly decreased for the granite wash. A conservative approach is maintained by using the larger value for each hydrogeologic unit.

The vertical permeability of 2.8×10^{-4} md for the Evaporite aquitard was derived from the harmonic means of permeabilities of its substrata in two typical cross sections through the evaporite strata. Typical or measured permeability values were used for each substratum: 0.0001 md for redbed shale (Davis and DeWiest, 1966), 0.0073 to 0.012 md for salt and anhydrite (Davis and DeWiest, 1966; Peterson and others, 1981), and 0.24 md for dolomite (Dutton and Orr, 1985). Table 2 includes the typical permeability values of

carbonates, shale, and granite wash taken from the literature.

Porosity

No direct porosity measurements are available for the Deep-Basin Brine aguifer of the Palo Duro Basin. An indirect method using neutron-density logs yielded quantitative porosity determinations of the Wolfcamp and Pennsylvanian strata (Conti and Wirojanagud, 1984). From two neutron-density logs, which penetrate the Pennsylvanian strata at the Stone and Webster Sawyer No. 1 test well in Donley County and the Stone and Webster Mansfield No. 1 test well in Oldham County, porosity values of the Wolfcamp and Pennsylvanian carbonates and granite wash were estimated at 50-ft (15-m) intervals according to the procedure described by Schlumberger (1979). Results of the analyses and some typical porosity values are given in table 3. Using 20 neutron-density logs in the Palo Duro Basin, Conti (1984) made preliminary determinations of Wolfcamp carbonate porosity distributions (fig. 7).

Table 3. Porosity of hydrogeologic units, Palo Duro Basin.

	Porosity from neutron-density log analysis Hydrogeologic unit Porosity from neutron-density log analysis Standard deviation Number of data		ron-density log analysis		
Hydro					Typical value*
Evap	orite strata		•	•	Less than .01
š	Wolfcamp carbonate	.08 (.064)**	.055	53 data points from a 50-ft interval at Sawyer No. 1	.06312
Deep-	Pennsylvanian carbonate	.08	.055	and Mansfield No. 1 wells	
Basin Brine	Shale	-	•		.0525
aquifer	Granite wash	.23	.12	18 data points from a 50-ft interval	.1127
	Pre-Pennsylvanian rock	-	•	·	•

^{*}From Davis and DeWiest (1966) and Davis (1980).

^{**}Average value for Wolfcampian strata (Conti and others, 1985).

Index to TAB 7: Maps related to ground water conditions

Berkstresser and Mourant, 1966, Plate 1, part

Dutton, 1987, Figure 16

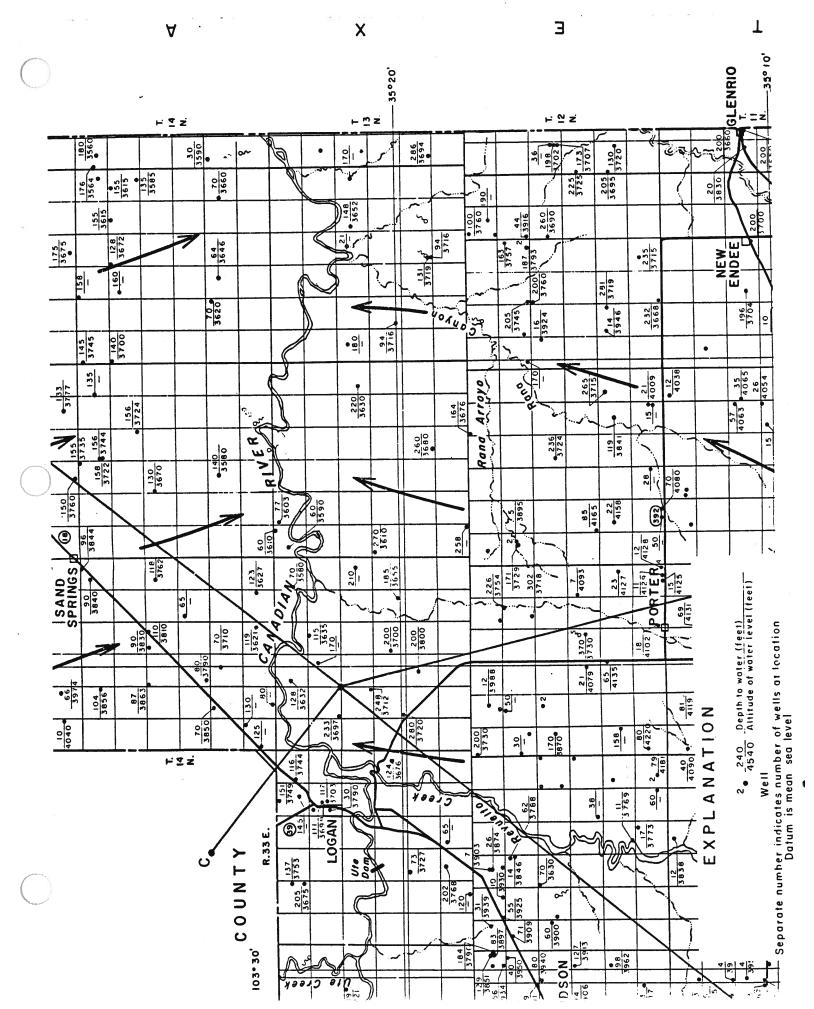
Dutton and Orr, 1986, Figure 5

Dutton and Orr, 1986, Figure B1

HGC, 1984a, Figure 22

Orr, et al., 1985, Figure 1

Orr, et al., 1985, Figure 3



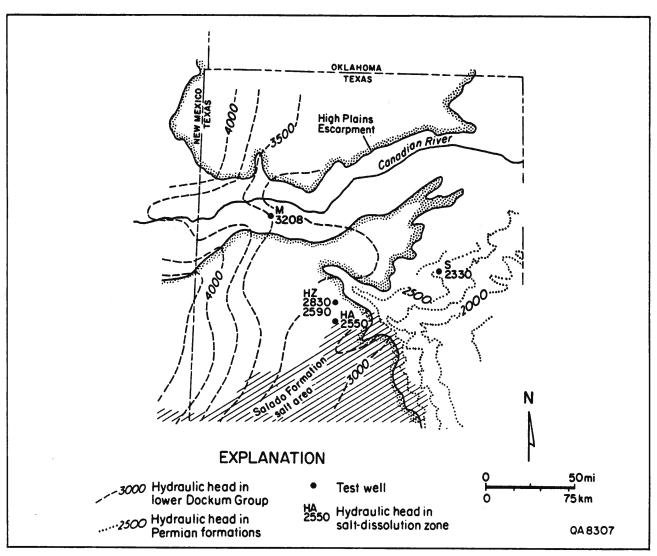


Figure 16. Potentiometric surface of ground water in aquifers overlying salt-dissolution zones and point values of hydraulic heads in salt-dissolution zones. (Modified from Dutton and Simpkins, 1986.)

the Upper Permian section. Reasons for the continuous downward decrease in hydraulic head include (1) ground-water flow through Triassic and Upper Permian mudstones of low permeability results in a loss of hydraulic head; (2) recharge areas for salt-dissolution zones occur at lower elevations around the Southern High Plains than do recharge areas for the upper aquifers, and this limits the maximum hydraulic head in each hydrologic unit; and (3) the discharge rate from the salt-dissolution zones is greater than the recharge rate to each hydrologic unit and prevents hydraulic head from building up to a hydrostatic level at equilibrium with the overlying column of water.

There are too few wells to map potentiometric surfaces of ground water in salt-dissolution zones in different formations that vary in hydrogeologic properties and hydraulic heads. However, because topography strongly influences potentiometric surfaces of ground water in near-surface aquifers (Toth, 1962, 1963, 1978; Hitchon, 1969), the general shape of potentiometric surfaces of salt-dissolution zones can be inferred from the potentiometric surfaces of ground water in the Dockum Group beneath the Southern High Plains and in Permian formations at shallow depth beneath the Rolling Plains. Potentiometric surfaces of ground water in the overlying aquifers (fig. 16) are clearly influenced by topography. For example, folds in the potentiometric surface of the lower Dockum Group along the northern and western limits of the Southern High Plains mark topographically controlled ground-water-basin divides that separate Dockum Group ground water beneath the Canadian River and Pecos River valleys from ground water beneath the Southern High Plains (Dutton and Simpkins, 1986). Because of the overriding influence of topography, the

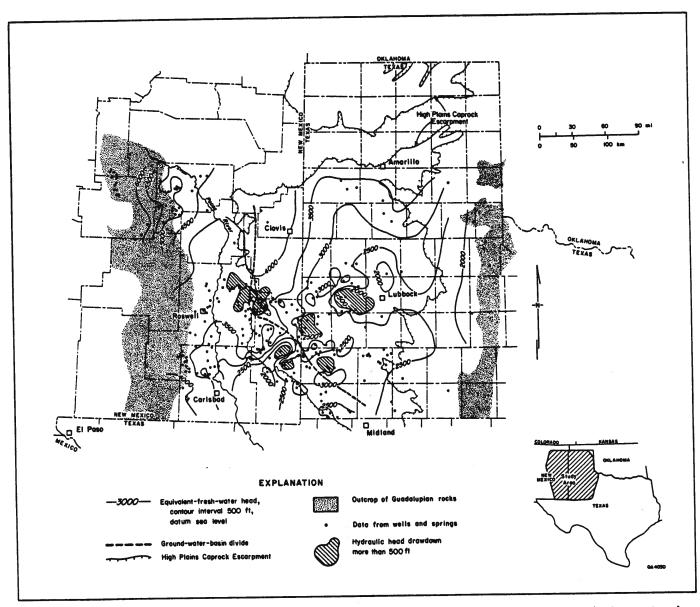


Figure 5. Potentiometric-surface map of the San Andres Formation prepared by hand-contouring hydraulic-head data calculated from shut-in pressures and weight of fresh water.

Chemical Data

We took 50 chemical analyses of ground water in the San Andres Formation below the Pecos Plains from compilations by Griggs and Hendrickson (1951), Hendrickson and Jones (1952), Smith (1957), Nicholson and Clebsch (1961), Berkstresser and Mourant (1966), Hem (1970, p. 145), Mourant and Shomaker (1970), and Dinwiddie and Clebsch (1973). We culled chemical analyses of 161 samples of saline ground water in oil

fields from surveys by Hiss (1975) and by Petroleum Data Service of the University of Oklahoma. The chemical analyses are listed in appendix C. Chemical analyses of San Andres water vary in completeness and in conditions of sample treatment. Samples of brine from oil fields are of two types: those collected during drill-stem tests and those obtained from a wellhead or storage tank at producing wells. Oil field brine samples may be contaminated by drilling mud or by reaction with well hardware or storage tanks (Bassett and

Table B1. Geostatistical parameters used to determine the potentiometric surface of the San Andres Formation by kriging.

Parameter	This study*	Orr and Dutton (1983)*
Class size	7,000 m	5,000 m
Block width	20,000 m	20,000 m
Radius	40,000 m	10,000 m
Range of influence	41,250 m	21,500 m
Nugget	50,000 ft ²	44,309 ft ²
Sill	240,000 ft ²	164,309 ft ²

^{*} Units are those originally used in calculations.

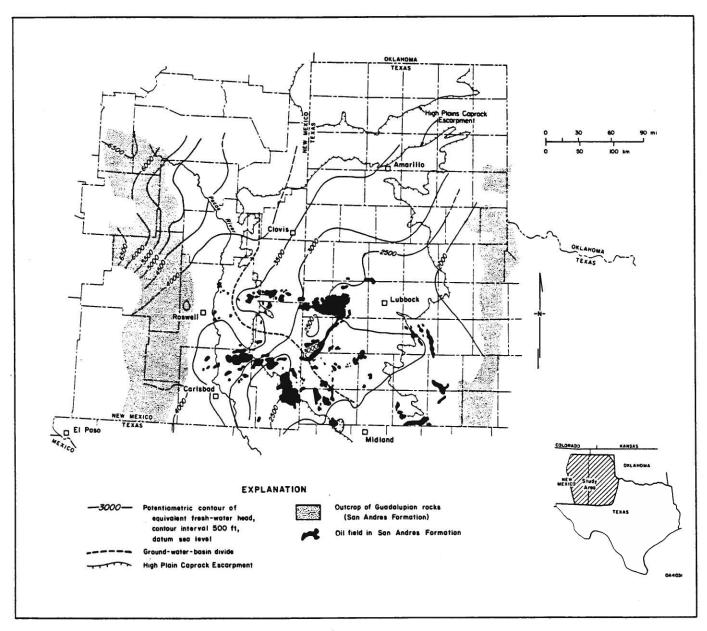
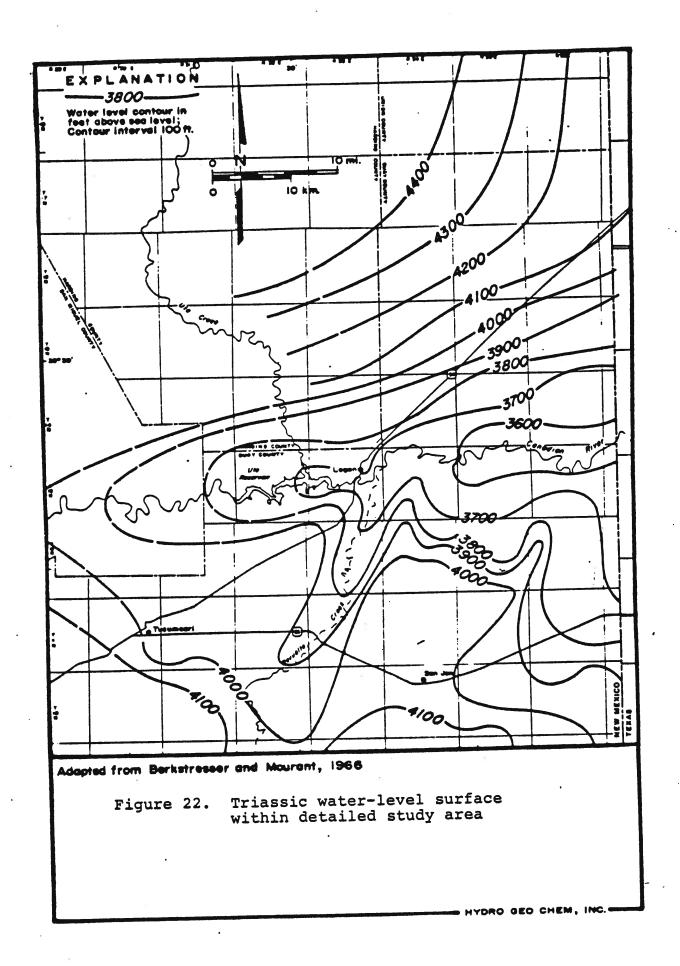


Figure B1. Potentiometric-surface map of the San Andres Formation made by kriging data on equivalent fresh-water head.



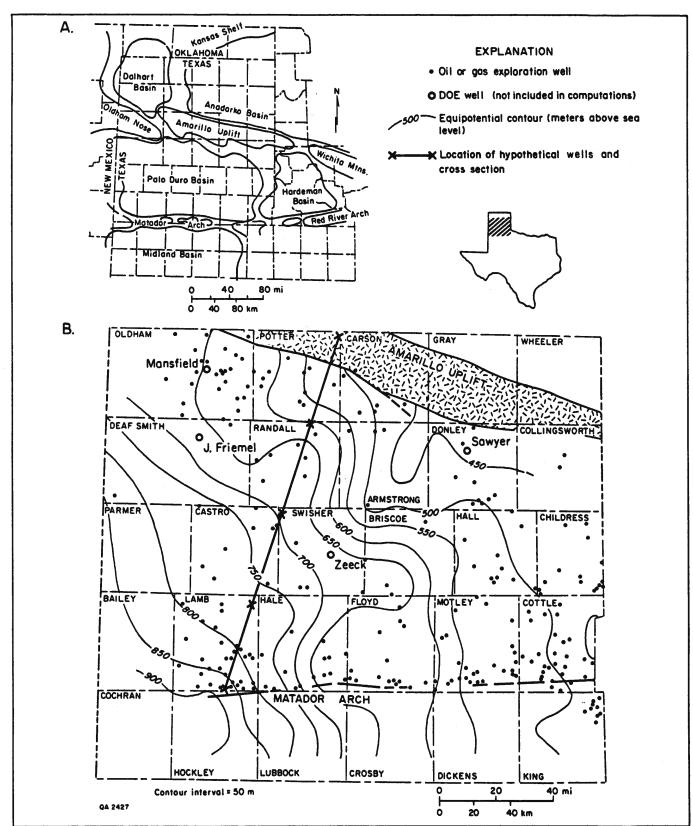


Figure 1. The study area: (A) Structural features of the Texas Panhandle and adjacent areas. (B) Average potentiometric surface of the Deep-Basin Brine aquifer (Wirojanagud and others, 1984) and location of drill-stem tests used in this study. The map indicates that ground water flows from the southwestern part of the Palo Duro Basin northeastward toward the Amarillo Uplift. In the southeastern part of the basin, flow is more easterly. Location of hypothetical cross section depicted in figure 5B is shown.

possible sources of variation, such as data quality and the hydrogeologic setting, (4) document and areally delineate any varying potentials for vertical flow within the Deep-Basin Brine aquifer, and (5) evaluate the significance of any potential components of

vertical flow by estimating vertical flux and flow volumes. By accomplishing these objectives, the present-day potentiometric surfaces, pressure-depth conditions, and vertical hydraulic gradients of the Deep-Basin Brine aquifer can be better understood.

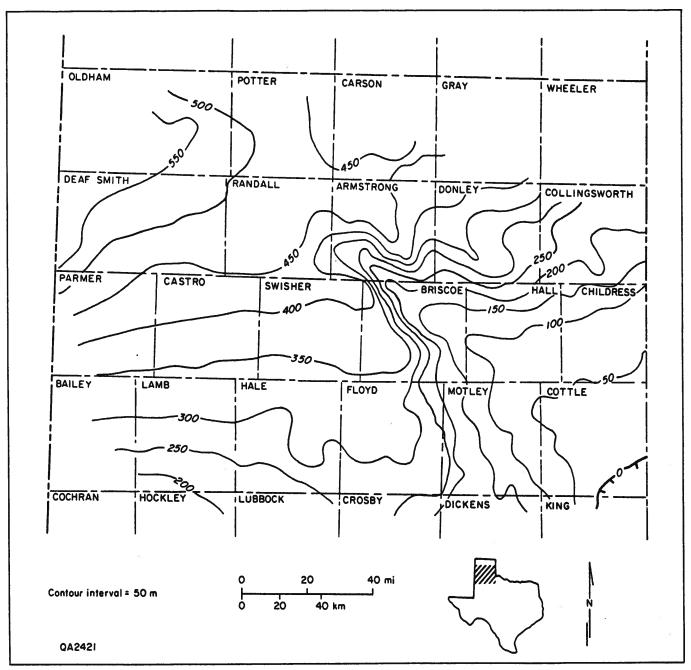
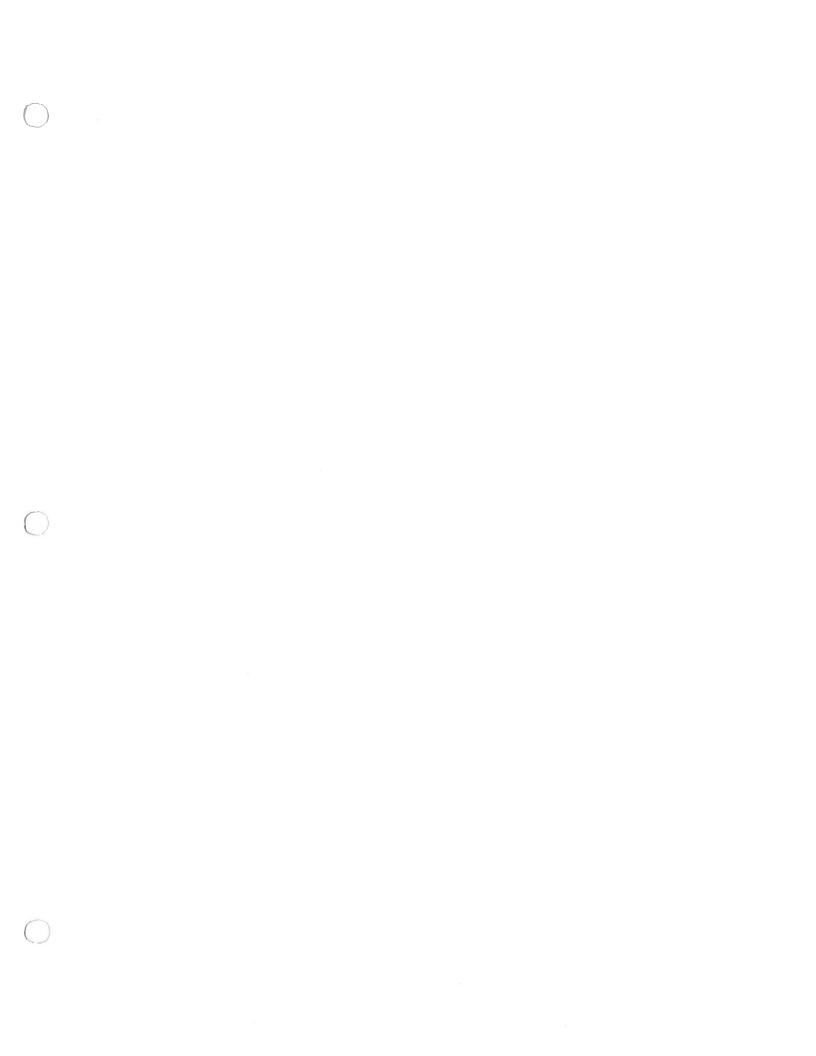


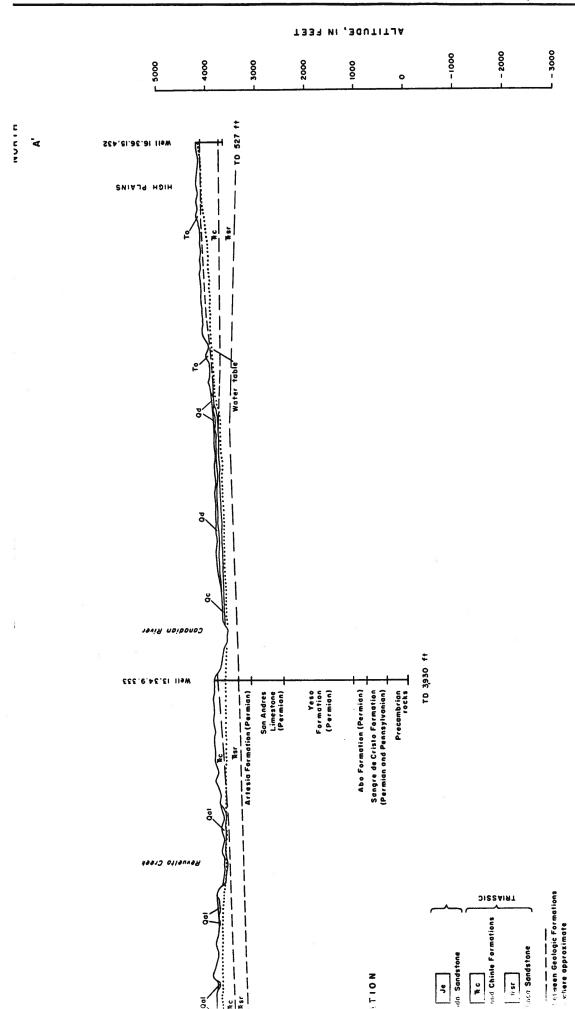
Figure 3. Contour map of differences between potentiometric surfaces of the Ogallala and Dockum aquifers and the Wolfcampian aquifer (Wirojanagud and others, 1984). Potential for downward ground-water flow through the Evaporite aquitard is greatest in the northwestern part of the Palo Duro Basin, where head differences are high, and least in the southeastern part of the Palo Duro Basin, where head differences are low.



Index to TAB 8: Hydrogeologic cross-sections

Berkstresser and Mourant, 1966, Plate 3 (portion)
Berkstresser and Mourant, 1966, Plate 5 (portion)
Dutton, 1987, Figure 8
Gustavson, et al., 1980a, Figure 31

Orr, et al., 1985, Figure 7



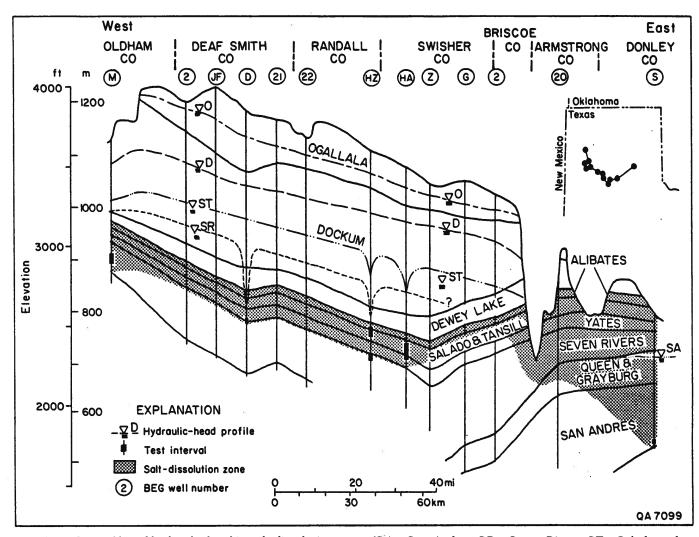


Figure 8. Profiles of hydraulic head in salt-dissolution zones (SA = San Andres, SR = Seven Rivers, ST = Salado and Tansill) and in aquifers in the Dockum Group (D) and Ogallala Formation (O).

RESULTS

Hydrologic Testing

Hydrologic testing indicates that hydraulic conductivity of salt-dissolution zones in Permian formations is low to moderate (2 × 10⁻⁴ to 1.6 ft/d; 6 × 10⁻⁵ to 0.5 m/d). Hydraulic head of ground water in dissolution zones is lower than the hydraulic head in overlying aquifers (fig. 8), confirming the potential for downward movement of ground water from aquifers in the Dockum Group and Ogallala Formation into the Upper Permian section. The coefficient of storage estimated at two wells is about 10⁻⁴, indicating that these salt-dissolution zones are confined. The following sections detail hydrologic properties at each test well.

SWEC Sawyer No. 2 well

The strong influence of wellbore storage on water levels in the SWEC Sawyer No. 2 well during drawdown

and recovery periods is indicated by the linear relation between water-level change and elapsed time on logarithmic plots (fig. 9). Discharge during the first 100 mir of each test came from storage. As shown in figure 9a, the drawdown was 10.0 psi after 5½ min. From equation 5

$$C = \frac{QB\Delta t}{\Delta P} = \frac{(10.8 \text{ gal/min}) (1.0) (5.5 \text{ min})}{(10 \text{ psi}) (7.48 \text{ gal/ft}^3)}$$

$$C = 0.794 \text{ ft}^3/\text{psi} = 0.366 \text{ ft}^3/\text{ft}$$

assuming that the dimensionless formation-volume factor (B) is 1.0 for the pumping well and that specific weight of the brine is 0.461 psi/ft. The actual well capacity of the 8%-inch casing (I.D.=8.097 inches [20.566 cm]) is 0.358 ft³/ft (0.033 m³/m). The close agreement between

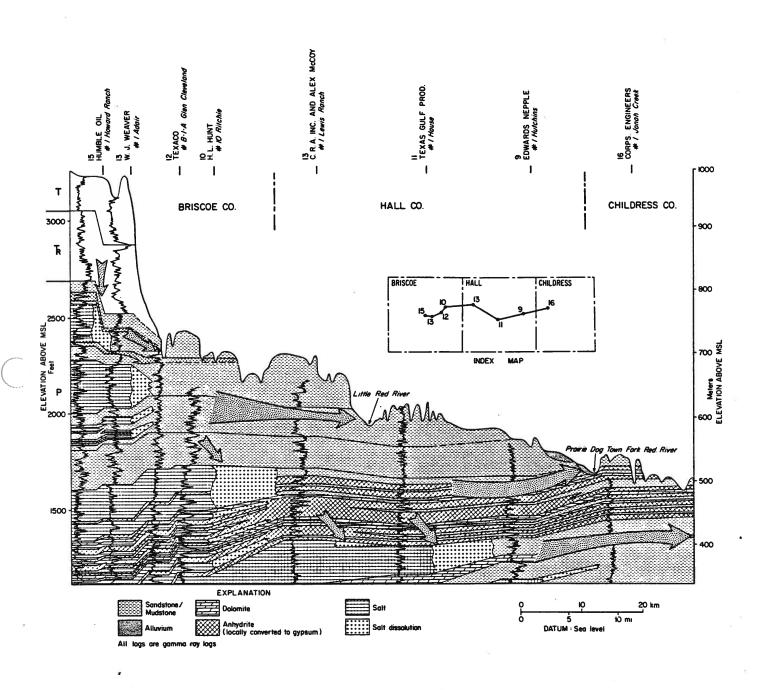
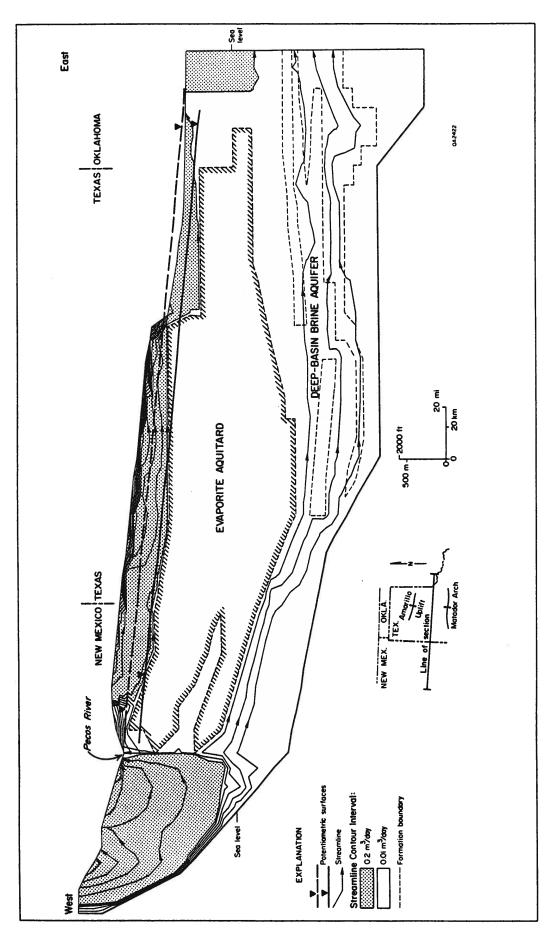
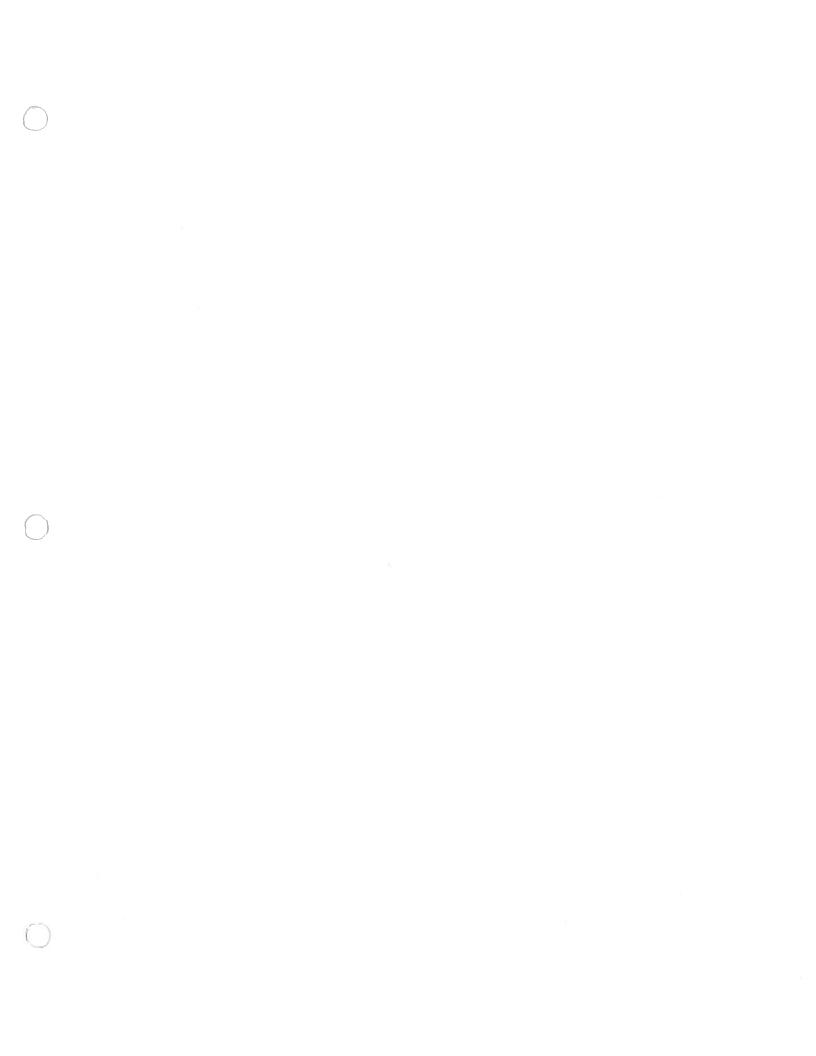


Figure 31. Conceptual model of ground-water movement and its impact on Permian salt units beneath the eastern Caprock Escarpment. Inferred movement paths are indicated by the arrows. Stratigraphic units are identified in figure 6.



approximate level of the actual potentiometric surface. Flow tubes calculated by cross-sectional modeling indicate that much of the potential recharge to the Deep-Basin Brine aquifer is discharged in the Pecos River valley. dashed line indicates the expected potentiometric surface of the Deep-Basin Brine aquifer. Solid line indicates the Figure 7. Simplified cross section through the Palo Duro Basin from eastern New Mexico to western Oklahoma. Heavy



Index to **TAB 9**: Well and spring inventories from the study area

Significant wells in Logan area

Deep wells in Logan area with both water level and water quality data

Shallow wells in Logan area with both water level and water quality data

Berkstresser and Mourant, 1966, Table 1, (parts)

HGC, 1984a, Table 1

Berkstresser and Mourant, 1966, Table 2

Significant Wells in Logan Area

	Comments	Artesian flow of about 30 gpm encountered at 261 feet; river water samples collected during drilling.	Artesian flow of about 3 gpm encountered at 466 feet; river water samples collected during drilling.	Trissic shale caved continuously; core sample 409.3-569.5 feet. Pumping in 7/84 may have dislodged drilling mud or foreign materials, resulting in more representative water level readings.	Pumped for 97 hours 3/78 at 475 gpm; water-quality and completion data suggest both brine aquifer & higher, less-saline aquifer were tested; assuming 425 gpm from brine aquifer T = 2250ft²/d, S = .00013.	Artesian flow encountered at 294 feet; core sample 261-318 feet; soundings indicate broken casing near top of screen is open to aquifer above brine aquifer; drilled with bentonite mud.
Water Quality	Parameters	Cl,MaCl,SO4, total Fe, conductance	Ci,NaCi,SO ₄ ,total Fe, conductance Field: conductance, T; Lab: Na,Ng,Ca,K,Ci,SO ₄ HCO ₃ ,CO ₃ ,TDS,PH	Field: conductance, T; Lab: Na,Mg,Ca,K,Cl,SO ₄ , HCO ₃ ,CO ₃ ,TDS,pH	pH,Na,Cl,SO4, others?	
Water	Number of Analyses	19 ^C in 8	2 42 2 42	į	* - * - * - * - * - * - * - * - * - * -	
	Water Level Measurements	Water levels of packered intervals in hole during drilling	Water level recorder 2 months 1983 ^f	Measured just after completion and monthly 10/83- 7/84; pumped; measured 8 & 9/84	water levels meas- ured during aquifer test; 4 measurements in 8/82; water level recorder 5/83-8/84	Water levels mess- ured during aquifer test
	Log Data	Geologic Log	Geologic Log	Geologic Log Gamma Ray Log	Gamma Ray Log	Geologic Log Gamma Ray Log Electric Log7 ⁸⁸
	Date <u>Drilled</u>	57/9	\$1/7	8 2 9/83	2-3/78	9-10/77
	Total <u>Depth</u>	356		569.5	358	E 80
Land	Surface <u>Elevation</u>	3674.57b	3655.727	3781.0	3674.01	3674.737
	Location	Lat 35º21'13"N Long 103º24'51"E	Lat 35 ⁰ 22'10"N Long 103 ⁰ 22'35"E ^d	Lat 35 ⁰ 21'05"N Long 103 ⁰ 25'40"E	adjacent to DH-1 Lat 35 ⁰ 21 ¹ 12 ¹¹ N Long 103 ⁰ 24 ¹ 50 ¹¹ E	38 feet Su of
	8	USBR DH-19	USBR DH-28	USBR DH-3 ¹	USBR 1-1-1	USBR 1-WO

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		Land					Water	Water Quality	
Neme	Location	Surface <u>Elevation</u>	Total <u>Pepth</u>	Date <u>Prilled</u>	Log Data	Water Level Neasurements	Number of Analyses	Parameters	Comments
USBR OM-28	220 feet SW of TW-1	3676.887	348 848	10/77-1/78	Geologic Log Gamma Ray Log Electric Log? ^m	Water levels measured during aquifer test			Core sample 300-348 feet; drilled with bentonite mud; no water-level response during pump test, so presumed plugged with drilling mud.
USBR 0N-3ª	556 feet SW of TW-1	3678.370	362	1/78	Geologic Log Gamma Ray Log Electric Log? ^m	Water levels messured during aquifer test; also monthly 11/83-9/84 f	4 6 6 7 6 7	Field: conductance, T; Lab: Na,Ng,Ca,K,Cl,SO4, HCO3,CO3,TDS,pH stable 0 and H isotopes carbon-14 tritium Cl,Br,I Field: T, pH, alkalinity, conductance; Lab: Ca,Mg,Na,K, CO3,HCO3,Cl,SO4,NO3, TDS by summation, B	a, K,
USBR ON-4ª	613 feet NW of TW-1	3676.57	382	1/78	Geologic Log Gamma Ray Log Electric Log? ^m	Water levels meas- ured during aquifer test; also 7/19/84	8	Field: conductance, T; Lab: Na,Ng,Ca,K,Cl,SOg, HCO ₃ ,CO ₃ ,TOS,pH	
USBR River Site O ⁱ	Piezometer at toe of Ute Dam Lat 35°20'40"N Long 103°27'36"E	3682.7	22				8	Field: conductance, T; Lab: Ma,Mg,Ca,K,Cl,SO ₄ ; HCO ₃ ,CO ₃ ,TDS,pH	Pumped once - clogged. Lake samples also collected.

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	Name Lo	USBR 1. River Site 1 ⁱ Ut	Piezometer A Piezometer B	USBR 2. River Site 2 ¹ UG	Piezometer A Piezometer B Piezometer C
	Location	1.6 miles below Ute Dam Lat 35°21'12"N Long 103°25'17"E		2.2. miles below Ute Dam Lat 35°2/14"N(?) Long 103°24/52"E	
Land	Elevation		9.899 6.83.8 9.899		3668.7 3668.7 3668.5
Total	Depth		25 94		55 22 22
Date	Drilled		1983		1983 1983 1983
	Log Date		Driller's Log		Driller's Log Driller's Log Driller's Log
Water Level	Neasurements		1 measurement in each piezometer on 8/24/63		1 measurement in each piezometer on 8/24/83
Water Quality Number of	Analyses		19t		19t 7
quality	Parameters		At each piezometer - Field: conductance, T; Lab: Cl, TDS, pH, conductance Besides the parameters listed above, Lab: Na, Ng,Co,KSO4,HCO3,CO3		At each piezometer - Field: conductance, T; Lab: Cl, TDS, pH, conductance Besides the parameters listed above, Lab: Ma, Mg,Ca,K,SO4,HCO3,CO3
	Comments		Piezometer A TD in bedrock; water quality samples obtained by air lifting. River samples also collected.		Steel drill bits left in holes A & B; bedrock at 59.3 feet; water quality samples obtained by air lifting. River samples also collected.

At plezometer A only - Field: T, pH, alkalinity, conductance; Lab: Ca,Mg,Wa, K,CO₃, HCO₃,Cl,SO₄,NO₃, TDS by summation, B Lab: Cl,Br,I

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	Comments		Bedrock at 34 feet; water quality samples obtained by air lifting. River samples also collected.	, No		Bedrock at 18-20 feet; bedrock soft sandstone; water quality samples obtained by air lifting. River samples also collected.
Water Quality	Perameters		At each piezometer - Field: conductance, T; Lab: Cl, TDS, pH, conductance Besides the perameters listed above, Lab: Na, Ng, Ca, K, SO, HCO3, CO3	At piezometer A only - Field: T, pW, alkalinity, conductance; Lab: Ca,Mg,Na K,CO3,KI,SO4,NO3, TS by summation, B Lab: CI.R.1		At each piezometer - Field: conductance, T; Lab: Cl, TDS, pH, conductance Besides the parameters listed above, Lab: Na, Ng,Ca,K,SO4,HCG3,CO3
Vater	Number of Analyses		19t 7	<u>ā</u> ā		18u 7
	Water Level Measurements		1 messurement in each piezometer on 8/24/83			1 measurement in each piezometer on 8/24/83
	Log Data		Driller's Log Driller's Log			Driller's Log Driller's Log
	Date Drilled		1983			1983
	Total <u>Depth</u>		34 20 34			20.5 15
Land	Surface <u>Elevation</u>	e ·	3655.1			3653.7 3653.7
	Location	5.4. miles below Ute Dam Lat 35°22'00"N Long 103°23'30"E			Revuelto Creek 0.2 mile above confluence Lat 35°21'48"N Long 103°22'58"E	
	Name.	USBR River Site 3	Piezometer A Piezometer B		USBR River Site 4 ⁱ	Piezometer A Piezometer B

Because of access problems, piezometers planned for <u>River Site 5</u>, Revuelto Creek about 2.1 miles above confluence, were never installed.

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	Connents		Steel drill bits left in all three holes;	bedrock at 52 feet; Piezometer A failed	early in 1984; water quality samples	obtained by air lifting. River samples	also collected.					Na,		
Mater Quality	Parameters		At each Diezometer -	Field: conductance, T;	Lab: Cl, TDS, pH,	conductance	Besides the parameters	listed above, Lab: Na,	Mg, Ca, K, SO,, HCO3, CO3	At piezometer A only -	Field: T, pH, alkalinity,	conductance; Lab: Ca,Mg,Na,	K, CO3, HCO3, Cl, SO4, NO3,	IDS by summation, B Lab: Cl,Br,I
Water	Number of Analyses			19^							<u> </u>			19
	Water Level Measurements		1 measurement in	each piezometer on	8/24/83									
	Log Data		Driller's Log	Driller's Log	Driller's Log									
	Date Drilled		1983	1983	1983									
	Total <u>Depth</u>		20	31	21									
Land	Surface <u>Elevation</u>		3638.0	3637.9	3637.6									
Sp.	Location	USBR 9.9. miles below River Site 6 ¹ Ute Dam Lat 35°23'30"W Long 103°20'22"E												
	Keme	USBR River Site 6 ¹	Piezometer A	Piezometer B	Piezometer C									179

- a All data from USBR, 1979, unless noted otherwise.
- USBR (1979), Appendix D, lists ground level elevation as 3680 feet; USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3, lists elevation as 3674.5 feet.
- 8 samples to depths of 156 feet collected by air lift, 2 after adding water to hole; 11 samples of artesian flow collected from 261-356 feet, 2 after circulating (USBR, 1979).
- USBR, 1979, Appendix D, lists location as Longitude 103⁰22[,]35" E; USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3, lists location as Longitude 103⁰22^{,32}#E.
- USBR (1979), Appendix D, lists ground level elevation as 3665 feet; USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3, lists elevation as 3655.72 feet.
- f USBR, 1984 (Hydrology/Hydrogeology Appendix).
- Samples of minor artesian flows collected during drilling; note that drilling fluid returns were monitored for electrical conductivity. (USBR, 1979).
- USBR, 1984 (Hydrology/Hydrogeology Appendix): Samples collected from artesian flow at wellhead; CO₂ outgassing and problems with well completion make representativeness of samples suspect. Note that when data were reported in MGC, 1984a, parameters NO₃, Fe and F were included; data source listed was USBR file data.
- i All data from USBR, 1984 (Hydrology/Hydrogeology Appendix), unless noted otherwise.
- Well was airlifted for about 1 hour; conductance had not stabilized when samples were collected. These facts, and the possiblity of CO2 outgassing, make the representativeness of samples suspect (USBR, 1984 - Mydrology/Mydrogeology Appendix).
- USBR, 1979 (p. 19) states that "numerous" samples were taken directly from the pump discharge line during the aquifer test; only the results of two "partial" analyses were reported.
- USBR (1979), Appendix D, lists ground level elevation as 3674.73 feet; USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3, lists elevation as 3675.9 feet (Pipe).
- USBR, 1979,(p. 17) states that "electric logging" was done in the four obervation wellbores; gamma ray logs are presented in Appendix D of the report.
- USBR (1979), Appendix D, lists ground level elevation as 3676.88 feet. USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3, lists elevation as 3682.8 feet (pipe); then notes that 1 foot of pipe has been cut off since elevation was determined.
- USBR (1979), Appendix D, lists ground level elevation as 3672.81 feet; USBR, 1984 (Hydrology/Hydrogeology Appendix) lists elevation as 3673.0 feet (Figure 3) and 3678.3 feet (Table 1).
- P Reported in USBR, 1984 (Hydrology/Hydrogeology Appendix).
- 9 Reported in HGC, 1984a.
- USBR (1979), Appendix D, lists ground level elevation as 3675.51 feet; USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3, lists elevation as 3676.5 feet.
- USBR, 1984 (Hydrology/Hydrogeology Appendix). Data also reported in HGC, 1984a except additional parameter Fe included; data source listed was USBR file data.

USBR, 1984 (Mydrology/Mydrogeology Appendix) does not include precisely 19 analyses for each of these parameters, but the data do break into these two main groups. Note that additional parameters MO₃, Fe and F were included when some of the analyses were reported in MGC, 1984s; data source listed was USBR file data.

USBR, 1984 (Nydrology/Nydrogeology Appendix) does not include precisely 18 analyses for each of these parameters, but the data do break into these two main groups. Note that additional parameters NO₃, Fe and F were included when two of the analyses were reported in HGC, 1984a; data source listed was USBR file data.

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USBR, 1984 (Mydrology/Mydrogeology Appendix) does not include precisely 19 analyses for each of these parameters, but the data do break into these two main groups. Because Piezometer A failed early in 1984, there are fewer samples in both groups for it. Note that additional parameters NO₃, Fe and F were included when some of the analyses were reported in NGC, 1984a; data source listed was USBR file data. >

DEEP WELLS IN LOGAN AREA WITH BOTH WATER LEVEL AND WATER QUALITY DATA

Comments	Samples air-Lifted during drilling of well.	Sample air-lifted during drilling of well.	Samples air-lifted during drilling of well.	Samples air-lifted during drilling of well.	Well flowed at eat. 30 gpm during drilling.				Est. 3 gpm flow encountered at 466 ft. during drilling.	Artesian pressure 0.5 psi.
Range of conductivity values, umhos/cm/	10,200-11,000/(2)	6,100/(1)	6,750/(2)	10,000-10,100/(3)	8,600-37,000/(4)	45,000/(1)	52,000/(1)	49,000-51,000/(4)	16,300/(1)	16,900/(1)
Range of TDS values, mg/l/ (number of analyses)	(0)/-	(0)/-	(0)/-	(0)/-	(0)/-	(0)/-	(0)/-	(0)/-	(0)/-	(0)/-
Water level elevation, ft/(date)	·		•	ı	·	,	ı	,		
Water level, ft below land surface/		8 ft below top of casing, about river level/(6/75)	4 ft below top of casing/(6/75)	3.75 ft in pipe/ (6/75)	Flowing/(6/75)	Flowing/(6/75)	Flowing/(6/75)	Flowing/(6/75)	Flowing/(7/75)	Flowing/(7/75)
Land surface elevation, ft	3674.57b	¥							3655.727 ^C	
<u>Vell/Interval</u>	DN-1/8 51-76 ft	71-96 ft	91-136 ft	131-156 ft	Open hole at 296 ft	Open hole at 316 ft	Open hole at 336 ft	Open hole at 356 ft	DH-2/a Open hole at 516 ft	Open hole at 534 ft

DEEP WELLS IN LOGAN AREA WITH BOTH WATER LEVEL AND WATER QUALITY DATA, CONT'D.

Connents	Water-level recorder installed 2 months in 1983; removed because hydrograph appears to be reflection of stream flows. No hydrograph presented in USBR, 1984 (Hydrology/Hydrogeology Appendix). When re-entered for logging in 1983, could not get tool below 160 ft.	Water levels measured 13 times between 9/83 and 9/84; level fell more than 5 ft after well pumped for sampling by air lift for 1 hour and foreign matter dislodged. See hydrograph. Conductivity had not stabilized at time samples were collected.	Water levels measured during aquifer test in 1978 and 11 times between 11/83 and 9/84. See hydro-graph. Data for aquifer test not reported in USBR (1979).	Water levels measured during aquifer test in 1978 and once 7/19/84. Data for aquifer test not reported in USBR (1979).	Water levels measured during aquifer test in 1978, 4 times in August of 1982 and by recorder 5/83-8/84. Data for aquifer test not reported in USBR (1979). No hydrograph presented in USBR, 1984 (Hydrology/Hydrogeology Appendix). "Numerous" samples of brine collected from flow line during aquifer test, but only 2 partial analyses reported in USBR (1979). "HDS" from these analyses is sum of Na, Cl and SO4. Aquifer test report notes that completion apparently includes part of an upper aquifer.
Range of conductivity values, umhos/cm/	17,500-17,800/(2) [†]	36,000/(2) [†] 9	, 60,000-78,400/(3) ^h	57,000/(1) [†]	(2)/-
Range of IDS values, mg/l/ (number of analyses)	11,985-12,138/(2) ^f	26,434-27,892/(2) ^f	49,072-51,005/(3) ^h	36,406/(1) [‡]	44,900-46,000/(2) ^j
Water level elevation, ft/ (date)	Мах. 3659.07/(9/83) [©] 11,985-12,138/(2) ^f Min. 3658.12/(7/83) [©]	Мах. 3695.86/(9/83) ^e Mín. 3689.10/(9/84) ^e	мах. 3681.26/(5/84) ^e Min. 3680.69/(6/84) ^e	3677.52/(7/84)	Max. 3675.0/(8/82) ^e Min. 3675.4/(8/82) ^e Max. 3674.67/(5/83) ^e Min. 3674.06/(8/84) ^e
Water level, ft below land surface	3.35 aboved 2.40 aboved	85.14d	2.96 aboved 2.39 aboved	1.02 aboved	2.0 above 1.4 above 0.66 above
Land surface Water level, elevation, ft below interval ft land surface	3655.727¢	3781.0	3678.379	3676.57 [†]	3674.01
Well/Interval	DM-2 (cont'd)/ Bottom of 42 ft of casing to 556 ft TD	DH-3/ Screened 368-417.5 ft	<u>04-3</u> / Screened 270-350 ft	<u>ON-4/</u> Slotted 292-376 ft	IN-1/ Not reported

NOTES

a Data from geologic log in USBR (1979), Appendix D, unless otherwise noted.

b USBR (1979), Appendix D lists ground level elevation as 3680 feet. USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3 lists elevation as "3674.5 (bolt)."

USBR (1979), Appendix D lists ground level elevation as 3665 feet. USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3 lists elevation as "3655.72 (top spigot-approximately land surface)."

Calculated

Data from USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 1.

Data from USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21.

USBR (1979), Appendix D lists ground level elevation as 3672.81 feet. USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3 lists elevation as "3673.0 (land surface)", while Table 1 lists surface elevation as 3678.3 feet.

Data from USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21 and HGC, 1984a, Appendix B.

USBR (1979), Appendix D lists ground level elevation as 3675.51 feet. USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 3 lists elevation as "3676.5 (land surface)."

j Data from USBR (1979), p. 19.

SHALLOW WELLS IN LOGAM AREA WITH BOTH WATER LEVEL AND WATER QUALITY DATA

						Field		
Well/Interval (ft)	Land surface elevation, ft	Water level, ft below <u>surface</u> b	Water level elevation, ft (8/24/83) ^c	River elevation, ft (8/24/83) ^c	ros, mg/t range/(no. of analyses)/ [mean] ^d	conductivity, umhos/cm @ 25° C range/(no. <u>of analyses)/[mean]^d</u>	Average Cl:SO ₄ <u>Ratio</u> e	Comments
River Site 1, mile 1.6	9:							Gradient piezomete
Piezometer A Screen 17.5 - 21.5	3668.9	2.2	3666.69	3666.03	11,920-15,737/(19) [13,670]	17,200-25,000/(19) [20,613]	9.3	Bedrock a deepest p
Piezometer B Screen 11.5 - 15.5	3668.8	4.6	3663.40	3666.03	14,457-16,942/(19) [15,585]	19,000-31,000/(19) [23,360]	11.5	31,000 um as quest value is
River Site 2, mile 2.2	2:							Gradient all piezo
Piezometer A Screen 50.5 - 54.5	3668.7	. 6. 6.	3664.85	3664.40	14,172-15,947/(19) [14,950]	20,044-25,000/(19) [22,358]	7.4	Steel bit at 59.3 f
Piezometer B Screen 35.5 - 39.5	3668.7	0.4	366.70	3664.40	14,881-17,224/(19) [15,872]	21,040-36,000/(19) [23,395]	0.0	Steel bit umhos/cm questions is 26,000
Piezometer C Screen 17.5 - 21.5	3668.5	eo m	3664.73	3664.40	13,902-17,224/(19) [15,260]	19,948-26,000/(19) [22,813]	8.7	17,224 mg able by U
River Site 3, mile 5.4	**							Gradient in piezom
Piezometer A Screen 29.5 - 33.5	3655.1	5.6	3652.48	3652.72	8047-26,617/(19) [24,846]	23,500-40,000/(19) [34,624]	0.8	Bedrock a deepest p

it upward in alluvium, but shallow ter water level below river level.

at 22 feet, so may not be in part of channel.^C

umhos/cm conductivity value regarded sationable by USBR; next highest is 26,000.

nt mixed in alluvium; water level in szometers above river level.

it left in bottom of hole. Bedrock ifeet.

oit left in bottom of hole.^C 36,000 mm conductivity value regarded as nnable by USBR; next highest value 000.

mg/l TDS value regarded as question-v USBR; next highest value is 17,124.

it downward in alluvium; water level ometers at or below river level. Bedrock at 34 feet, so may not be in deepest part of channel. © 8047 mg/l TDS value regarded as questionable by USBR, as are values of 10,827 and 15,996; next value is 23,613. 23,500 umhos/cm

SHALLOW WELLS IN LOGAN AREA WITH BOTH WATER LEVEL AND WATER QUALITY DATA, CONT'D.

	Comments		4343 and 6029 mg/l TDS values regarded as questionable by USBR; next value is 10,714.		Bedrock 18-20 feet, believed to be near lowest point of bedrock channel. ^C 14,863, 15,701 and 25,774 mg/l TDS values regarded as questionable by USBR; next value is 14,921.	15,685 and 28,075 mg/l TDS values regarded questionable by USBR; next value is 10,955. 1600 umhos/cm conductivity value regarded as questionable by USBR; next value is 2500.	Gradient downward in alluvium, but shallow piezometer water level above river level.	Bedrock at 52 feet. Steel bit left in hole. ^C Piezometer failed early in 1984. 13,115, 17,796 and 34,558 mg/l TDS values regarded as questionable by USBR; next measured values are 20,077 (low) and 20,846 (high). 22,500 and 25,000 umhos/cm conductivity values regarded as questionable by USBR; next measured value is 28,000.
67 67 60 60 60 60 60 60 60 60 60 60 60 60 60	Cl:SO4 Ratio e		6.9		N. F.	2.5		7.2
Field conductivity umboscem a	25°/C range (no. of analyses) [mean] ^d		16,124-26,000/(19) [19,837]		2800-21,000/(18) [7868]	1600-16,500/(18) [5082]		22,500-32,000/(12) [30,218]
TDS, mg/l range/(no. of	enalysis) (mean)		4343-16,414/(19) [13,229]		1601-25,774/(18) [5168]	1256-28,075/(18) [3688]		13,115-34,558/(12) (20,319)
	elevation, ft (8/24/83) ^c		3652.72			,		3632.19
Water Level	elevation, ft (8/24/83) ^c		3652.73		3653.27	3653.18		3631.99
Water level.	ft below surface b		5.4		4.0	 		0.9
Land surface			3655.1	o Creek)	3653.7	3653.7	o.	3638.0
	Well/Interval (ft)	River Site 3 (cont'd)	Piezometer B Screen 15.5 - 19.5	River Site 4 (Revuelto Greek)	Piezometer A Screen 16-20	Piezometer 8 Screen 10.5 - 14.5	River Site 6, mile 9.9	Piezometer A Screen 45.5 - 49.5

SHALLOW WELLS IN LOGAN AREA WITH BOTH WATER LEVEL AND WATER QUALITY DATA, CONT'D.

Comments		Steel bit left in hole. ^c 10,034 and 22,613 mg/l TDS values regarded as questionable by USBR; next measured values are 12,035 (10w) and 15,048 (high). 16,844 and 16,928 umhos/cm conductivity values regarded as questionable by USBR; next value is 18,500.	Steel bit left in hole. ^C 11,250 mg/l TDS value regarded as questionable by USBR, though three other values are higher; 13,688 umhos/cm conductivity value regarded as questionable by USBR, though six other values are higher.
Average Ct:SO ₄ Ratio e		V.	e. 3
Field conductivity umhos/cm & 25°c range (no. of analyses) [mean] ^d		16,844-24,000/(19) [21,167]	10,800-17,852/(19) [13,583]
TDS, mg/l range/(no. of analysis) [mean] ^d		10,034-22,613/(19) 16,844-24,000/(19) [13,651] [21,167]	6366-11,842/(19) [8816]
TDS, mg/l River range/(no. elevation, ft analysis) (8/24/83) C [mean] ^d	•	3632.19	3632.19
Water level elevation, ft (8/24/83) ^C		3632.20	3632.24
Water level, ft below <u>surface</u> b		5.7	3.
Land surface elevation, ft a	۲.	3637.9	3637.6
Well/Interval (ft)	River Site 6 (cont'd)	Piezometer 8 Screen 26.5 - 30.5	Piezometer C Screen 16.5 - 20.5

NOTES

- a Land surface elevations from USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 1.
- b Water level elevations only are reported in USBR, 1984 (Hydrology/Hydrogeology Appendix). Water levels below land surface have been calculated for this table.
- c Screened intervals, water level elevations, river elevations and some of comments from USBR, 1984 (Hydrology/Hydrogeology Appendix), Figure 2.
- Only values reported in USBR, 1984 (Mydrology/Mydrogeology Appendix), Tables 5 through 20, are compiled here. Additional analyses were made on single samples from the deepest piezometers at Sites 2, 3 and 6, and are reported in MGC, 1984a. Samples collected by USBR were obtained by air lifting. Mean values are those reported in USBR, 1984 (Mydrology/Mydrogeology Appendix); they include USBR corrections to questionable values. v
- Calculated from mean Cl value (mg/l) divided by mean SO₄ value (mg/l). These means are of complete analyses only; some analyses had chloride but not sulfate. USBR, 1984 (Hydrology/Hydrogeology Appendix), Tables 5 through 20. w

GROUND WATER

QUAY COUNTY

TABLE 1. RECORDS OF SELECTED WELLS IN QUAY AND ADJOINING COUNTIES, N. MEX.

Location number: See text for explanation of well numbering system.

'aar completed: Wells designated "old" drilled generally before 1925.

pth: Depths are in feet below land surface. Reported depths are given to nearest foot. Measured depths are given to nearest tenth of a foot. ameter: The diameter of the casing, or the mean diameter of the well if uncased, to nearest inch.

Altitude: Altitude of land surface at well. Altitude interpolated from topographic maps, or aneroid determination to nearest 10 feet.

Water level: Reported depths are given to nearest foot. Measured depths are given to nearest tenth of a foot.

Stratigraphic unit: Qal, younger alluvium; Qc, upland cover of older alluvium; To, Ogallala Formation; Ks, Cretaceous sandstone and siltstone; Jm, Morrison Formation; Je, Entrada Sandstone; Rc, Chinle Formation; Rsr, Santa Rosa Sandstone; Pr, Permian rocks.

Type of pump and power source: E, electric; I, internal combustion; J, jet; N, none; P, plunger or cylinder; S, submersible; T, turbine; W, windmill. Use of Water: D, domestic; I, irrigation; Ind, industrial; O, observations; PS, public supply; RR, railroad; S, stock; N, none.

Remarks: All wells are drilled and cased with steel casing unless otherwise indicated. Ca, chemical analysis in table 3; dd, drawdown; est, estimated; gpm, gallons per minute; log, log in table 6; meas, measurement; perf, perforated, perforations given in feet below land surface; rept, reported, reportedly; T 61°F, temperature in degrees Fahrenheit; USBR, U.S. Bureau of Reclamation; yields are reported unless otherwise indicated.

Location No.	Owner or name	Year com- pleted	Depth (feet)	Diam- eter (inches)	Altitude (feet)	Depth be low land surface (feet))-	Strati- graphic unit	Type of pump and power source	Use of water	Remarks
5.26.22.320	Abercrombie and H Hawkins No. 1— Nappier	1949	7149	9	4518		_	****	-		Oil test; in DeBaca Co., 2½ miles west of Quay Co.
<i>5.27</i> . 1.341	L. W. Barnhill		200	5	4950	121.0	0.00 ==	_			line; log
3.312			77.1	5		131.0	8-23-55	To_	P,W	D,S	_
9.333	L. W. Barnhill		33.3	-	4650	49.6	4-15-55	Qal, "Rc	P,W,I	S	
12.444	L. W. Barnhill	1951		6	4530	28.5	8-25-53	Qal	P,W	S	Ca
15.424	Dick Ballew		182	6	4920	170	1951	To	P,W	S	T 61°F
10.424	DICK DOILEM	1945	64.8	4	4640	43.9	4-15-55	Qal, Rc	P,W	D,S	T 62°F. Ca
17.441	Mrs. N. G. Koll		7.6	48	4510						Pumping water level
25.242	D. O. Bomar	Old	86	40	4510	5.3	4-15-55	Qal, Rc	P,W	S	Dug. T 51°F
29.212	Guy Shipely			2	4890	75	1954	To	P,W	D,S	T 63°F
30.242	Mrs. N. G. Koll		35.1	7	4470	25.7	4-15-55	Qal	P,W	D,S	
00.272	mis. II. G. Koll		13.3	48	4440	12.9	4-15-55	Qal, Rc	P,W	S	Dug. Pumping water level.
31.122	Mrs. N. G. Koll		13.2	30	4420	13.2	4-15-55	Qal	814		Est yield 4 gpm
£00 1111				-			- -10-33	A01	P,W	S	Dug. Est yield 1 gpm.
5.28. 1.111	G. E. Murphy	1950	90	6	4750	70	1950	To	P,W	S	T 58°P John Maddox, driller

1.212	D. C. Wyatt	1943	102	10	4730	48.6 50.9	11-23-43 12-3-47	То	N	N	Yield 100 gpm, dd 40 ft. Rept destroyed; R. F. Davis, driller
1.221	D. C. Wyatt	1946	133.0	16	4720	46.6 54.9	3-29-46 1-11-56	То	T,1	1	R. F. Davis, driller
5.222	W. R. Crawley	Old	97	6	4900	80	_	To	P,W	D,S	Not cased; T 60°F
5.442	B. R. Hood	Old	100	_	4890	80		To	P,W	D,S	_
8.222	B. K. 11000		89.9	5	4880	87.8	8-23-55	To	P,W	D,S	_
18.131	L. W. Barnhill	Old	136.9	5	4910	120.2	8-23-55	То	P,W	D	Water level rising when measured. Weak
19.422	R. R. Adams	1952	110	6	4890	100	1952	То	P,W	D,S	Yield I gpm. John Maddox, driller
22.212	C. A. Morrow	1940	102.8	6	4820	91.0	8-23-55	To	P,W	S	
30.212	R. R. Adams	1955	107.9	6	4870	73.5	8-23-55	То	P,W	S	Perf. 50 to 110 ft. Rept yield 5 gpm. John Mad- dox, driller
31.222	Mr. Baxter	Old	121.0	4	4860	94.6	8-23-55	To	P,W	S	
33.421	-		106.3	5	4820	98.9	8-23-55	To	P,W	S	T 61°F
36.331	State of N. Mex.		95.6	10	4770	92.8	8-23-55	To	P,W	S	Water stains pipe yellow.
5.29. 1.000	Charlie Vance	1940	103	18	4730	-		To	N	N	Not cased. Yield 20 gpm. S. J. Davis, driller
2.131	O. G. Miller	1947	132	14	4710	24.2 25.2	1-19-48 1-11-56	То	P,W	S	Yield 35 gpm. R. F. Davis, driller
4.333	W. Y. Head	1949	128	18	4690	36.5 47.9	6-3-49 1-11-56	То	T,I	1	Not cased, red clay at 125 ft. A. L. Akin, driller
<i>5</i> .111	Wm. Young estate	1948	120	16	4720	_	_	To	T,I	- 1	R. L. Davis, driller
5.211	R. H. Currence	1943	108.0	12	4720	49.9 67.4	11-22-43 1-11-56	То	T,I	ı	R. F. Davis, driller
5.231	R. H. Currence	1943	134	_	4710	42.4 47.9	1-24-44 1-15-49	То	T,I	ı	R. F. Davis, driller
5.312	Willard Carpenter	1940	115	15	4710	41.2 64.5	4-30-41 1-11-56	То	T,I	i	
5.312a	Willard Carpenter	1955	122	18	4710	63.9	1-11-56	To	T,I	- 1	Lee Williams, driller
5.321	Spence Morris	1945	108	16	4710	44.8 64.1	6-13-45 1-11-45	То	T,1	ı	R. F. Davis, driller
5.341	I. D. Linville	1941	140	16	4700	33.6 54.6		То	T,I	!	Yield 1600 gpm. Mr.

Use

of

water

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Type of

pump

and

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P,W

S.E

Strati-

graphic

unit

Qal, Rc

Qal, Rc

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Rc, Rsr

Qal

Qal

Rc, Rsr

Rc, Rsr

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R c(?)

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Qal

R sr(?)

170.6 5-17-56 Rc, Rsr P,W

10-15-54 Rc, Rsr Rc

Altitude

(feet)

4140

4140

3840

3900

3900

3900

3960

4250

4000

3960

4180

4150

4030

3860

Diam-

eter

(inches)

13

13

5

48

48

5

48

42

Year

Depth

(feet)

36

29

167.1

300

258.0

6.6

279.4

200

300

244.2

35.1

65

90

32

21.4

15.2

165

com-

pleted

1953

Old

1949

1939

Old

1919

1950

Owner or name

L. C. Jackson

L. C. Jackson

Griffin, Trew, and Cooper

Frank Warmuth

T. E. Terry

Joe Yarborough

H. E. Norred

E. S. Norred

T. G. Rose

A. C. Ward

W. W. Sweeney Estate

Location

No.

12.34.36.444c

12.35. 2.111

36.444d

6.143

7.234

7.234a

7.234b

15.323

20.333

25.123

26.233

29.143

32.344

32.434

33.422

35.422

35.424

12.36. 2.112

Water level

Date

5-21-53

163.8 5-17-56

4.5 10-15-54

1939

10-28-54

11-6-54

11-5-54

235.9

85

285

118.9

22.1

50

70

28

15

100

20.7

Depth be-

low land

surface

(feet)

12.6

	78
Remarks	
Ca. Yield 40 gpm. Earl Flint, driller. Not cased Not cased. Earl Flint, driller	NEW
Pumping water level. Rept water level 55 ft. Ca —	NEW MEXICO BUREAU OF
Rept depth 380 ft. Rept water level 345 ft. Rept tastes salty	BUREAU
Dug. Rock casing	0
Dug. Rock casing	
Water at 85 ft.; drilled to 200 ft for storage. Weak	MINES &
Cased to 276 ft. Yield 4 gpm. Rept water slightly alkaline	MINERAL
Dug. Rept depth 45 ft. Not cased	
R. J. Thrasher, driller Water rept hard and alkali Dug. Culvert-pipe casing Rept depth 26 ft. Partly caved. Yield 5 gpm. Ca	RESOURCES

2.334	Bill Smithers		300			_		TR sr	N	N	Destroyed. Rept unusable water even for stock
2.334	Diff. Cim.					190	_	TR sr	P,W	D,S	
2.413	Bill Smithers	1938	225		_	170	<u>-</u>	TR sr	P,W	S	
7.334	A. C. Ward		190	5				₹¢ ₹sr	P,W	D,S	
	Louis Lee	_	150	5			_	Rsr Rsr	P,W	D,S	-
8.313	Louis Lee		225		3950	205			P,W	S	
9.334	A. C. Ward		175		3920	163	-	<u>R</u> sr	P,W	Š	
10.224		_	227	5	3980	187		₹ sr		D,S	Not cased, 1954. R. J.
10.434	Bill Smithers	1954	59.5	8	3960	44.2	11-6-54	٦ç	P,W	υ,3	Trasher, driller
11.333	J. L. Liles	1734	37.0	•							Dug. Rept inadequate yield
		2015	28	36	3960	_	_	TR¢	P,W	D,S	R. J. Thrasher, driller
11.333a	J. L. Liles	1945	285.4	6	3950	260.0	10-11-55	TR sr	N	N	R. J. Inrasner, armer
14.311	Ray Adams	1954		_	3960	200		TR sr	P,W	D,S	
16.111	Louis Lee		245		3940	15.9	11-5-54	Qal	P,W	S	Dug. Steel casing
18.242	-		20.2	14	3960	14.0	11-5-54	Qal	P,W	S	Rept depth 26 ft.
29.132	D. S. Gentry		18.6	6	4000	281.4	11-6-54	R sr(?)	P,W	D,S	
29.242	L. O. Gentry	1914	296	4		232	11-0-54	TR sr(?)	P,W	S	
33.222	L. O. Gentry	Old	250	_	3900	232		R sr	N		Yield 10 gpm with 25 ft
34.142	Henry Sasser	1955	540	_	3950	235		K.			drawdown during bail- ing test. Rieddell and Suggs, drillers
							101154	Qal	N	N	Dug
	_	1940	38.1	36	_		10-11-54		N	N	_
12.37.18.424	Ira Johnson		208.8	5	3900	197.6			P,W	S	_
18.4240	Ira Johnson	1954	193.4	6	3880	173.4			P,W	Š	T 63°F
18.442		1946	250	6	3950	225		<u></u> ₹ sr		D.S	Rept poor quality. Ca
19.133	R. L. Martin	1914	225		3900	205	_	TR sr	P,W		Lamb and Hill, drillers
30.133	R. L. Martin	1951	150		3850	130		TR sr	P,W	S	Yield 3.5 gpm. T 62°F. Ca
30.422	R. L. Martin		50		3900	40	_	Qal, Rc	P,W	S	Perf 128 to 140 ft. Weak
13.31. 1.124	R. R. Simms		140	6	4020	126	1918	Έc	P,W	D,S	Pert 128 to 140 tt. Week
25.344	R. S. Bell	1918		5	3950	150		٦ç	P,W	S	Perf 171 to 195 ft. Est yield
26.123	H. E. Osborne	1924	195	3	3730	100		• • • • • • • • • • • • • • • • • • • •	*		2 gpm. T 65°F
)				_	4020	75		Ές	P,W	_	Perf 73 to 85 ft. weak.
26.244	H. E. Osborne	1912	85	5	4020	/3		17.4	- •		T 62°F
20.244	••• •• • • • • •					150	_	٦ç	P,E	D,S	Yield 6 gpm. Pete Knowles,
34.244	H. J. Ellis	1952	175	6	3980	152		κ.	.,-	- / -	driller
34.244	11. 2. 2						_	Qc	J.E	D.S	Pete Knowles, driller
04.444	H. W. Brady	1952	34	10	4030	12	_		P,W	S	Yield 1 gpm
34.444	R. S. Bell	1920	76	6	4010	56		<u>T</u> e c		Š	
36.211		1940	279.3	4	4000	230.4	11-23-53	3 <u>T</u> ₹¢	P,W	Ň	Destroyed. Rept very salty
13.32. 4.311	R. R. Simms	1740	200	_	3940			٦ç	N	14	water
5.131	R. R. Simms						S			٠	WOIGE
			7418		3840		2 11-23-5	3 Tec	D \A'	v	

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	NEW
1	MEXICO
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	MINES
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							rlevel		Type of		
ocation No.	Owner or name	Year com- pleted	Depth (feet)	Diam- eter (inches)	Altitude (feet)	Depth be low land surface (feet)	d	Strati- graphic unit	pump and power source	Use of water	Remarks
13.32.16.211	R. R. Simms	1932	200	6	3910			Τ̈́c	P.W.E	D	Dick Seddon, driller
18.200	R. R. Simms	1940	222.6	5	3880	179.3	11-23-53	Ro	P,W	\$	Seddon and Crouch, drillers
1.311		_	176.1	6	3860	116.0	5-14-54	TR sr	P,W	S	
2.122	Sim McFarland	1951	700	8	3900	151.4	7-22-54	Qc, Rsr	N	Ň	Laughlin and Harris No. 1 McFarland oil test
13.33. 2.413	Mrs. Barker	Olq	160		_	145		R sr(?)	P.W	D,S	merariana on resi
5.244	Bob Rogers		138.0+		3890	137.4	9-17-54	TR c(?)	P,W	S	_
5.442	Arthur Hamby	1927	235	6	3880	205	_	R sr(?)	P.W	D,S	
11.112	Robert McFarland	1952	200	8	_			Qc, Rsr	T,E	ī	Cased to 70 ft. Yield 60 gpm
11.144	Chicago, Rock Island										ab
	and Pacific Railroad	1930	244.4	12	3810	110.9	4-1-54	TR sr	N	N	-
11.312 11.322	Mrs. Ann Bigelow Chicago, Rock Island	1920	50	6	3820	30	_	Qc	J,E	D,I	Used to irrigate lawn
	and Pacific Railroad	_	240.8	16	3820	117.3	8-4-53	TR sr	T,E	P,S	Yield 50 gpm, dd. 87.5 ft. T 64°F; Ca
24.412			150.9	5	3800	123.7	3-2-55	TR sr	P.W	D,S	1 04 1; Ca
28.121	R. L. Stansbury	_	76.8	6	3800	72.8	3-7-55	Rc Rsr	P.W	S	
32.433	R. C. Chance	1947	140	5		120		Rc Rsr	P,W	D,S	
33.124	L. O. White	_	211.2	7	3970	202.1	3-4-55	R sr	P,W,E	D,S	Rept soft water
34.114			81.4	5		65	_	TR sr(?)	P,W	S	
13.34. 1.133	Pyle Ranch	1938	80	_	3650	70		TR sr	P,W	S	_
4.134	J. A. Cox		140	6	3760	128.2	8-5-53	TR sr	P,W	S	T 67°F
8.333	J. A. Cox	1947	250	6	3930	233.4	8-5-53	Rsr	P,W	S	Perf 230 to 250 ft. Yield 6 gpm. Water has poor taste and odor, 69°F.
9.232	R. H. Haddon	1951	185	6	_	170		Ψ			R. J. Thrasher, driller
9.333	Olean No. 1 Woods	1926	3930		3920	170	_	TR sr	P,W	S	_
10.211	R. H. Haddon	1936	150	_	3750		_	_	N	N	Oil test. Log
13.234	Griffin, Trew, and Cooper	1952	231	6	3/30 	115 210	_	Te sr Te sr	P,W P,W	S S	Ca —

17.444	J. A. Cox		320	6	3960	248.1	8-4-53	TR sr	P,W	S	Perf 280 to 300 ft. Dick Seddon, driller	C R
20.333	B. J. Lawrence	1951	287	6	4000	280		TR sr	P,W	D,S	Rept poor quality. Dick Seddon, driller	GROUND
20.333	5. 5. 25				2000	200		TR sr	P,W	S		Z
22.311	W. L. Bloodworth	1907	240		3900 3840		10-20-54	TR sr	P,W	S		9
23.432	C. A. Eiland	1951	222.3	6	4000	200		TR sr	P,W	S	Rept poor taste	WATER
28.144	Tom Ayers		230		3680	76.7	4-8-54	TR sr	Ń	N	Good	3
13.35. 5.113	Pyle Ranch		93.0	6	3650	60.3	4-8-54	TR sr	N	N	Good	긆
6.143	Pyle Ranch	1910	78.1	5	3670	60.5		Rsr	P,W	D		Ħ
6.221	Pyle Ranch	1914	70	6	30/0	00		17.5	• • • •			
13.321	Griffin, Trew, and Cooper	Old	250		3850	220	-	TR sr	P,W	D,S	Rept poor quality; salty taste; disagreeable odor	
	and obspan					0/00	5-17-56	TR sr	P,W	D,S		
19.112	Elmer Wallin	1935	280.0	6	3880	269.8	10-20-54	R sr	P.W	S		
27.343	Tom Ayers		285.9	_	3940	260	10-20-34	K 31	.,	-		
31.444	Griffin, Trew,					050 1	10-14-54	TR sr	P,W	S	Pumping water level; est	QUAY
31.444	and Cooper		280	6	3920	258.1	10-14-54	14. 41	.,	-	yield 1 gpm	É
	and Coope.							TR sr	P,W	S	· · · · · · · · · · · · · · · · · · ·	2
13.36.13.234	A. C. Ward		185	_		170	10-12-55	E sr	P.W	Š		
14.134	A. C. Ward		185.0	5	3800	147.5		Qal, Rs		D,S	Dug. 65°F	COUNTY
15.231	A. C. Ward	Old	22.7	_	_	21.0	11-3-54		P,W	S		č
	A. C. Ward	_	190	_	 .	180		TR sr	P,W	Š	Est yield 2 gpm with 40.2	Z
18.231	A. C. Ward		284.5	5	3810	94.4	2-25-55	'R sr	r, 11	•	ft drawdown	=======================================
20.332	A. C. Wald							~	S,E	D,S	Perf 65-185 ft. Yield 8.5	F C,
27.332	A. C. Ward	1955	167.2	6	3850	131.3	2-25-55	TR sr	3,6	0,0	gpm. 62°F. Ca. Log. Thrasher and Flint, drillers	
				_		94.4	11-3-54	Te sr	S,E	D	Yield 15 gpm	
27.334	A. C. Ward	_	172.5	5	3810	74.4		Qal(?)	P,W	N	Dug. Inadequate supply	
3.37. 7.144	A. C. Ward	1920	20		3980	286.4	11-3-54		P,W	S	Pumping water level. Yield	
30.343	A. C. Ward	Old	300.4	5	3980	200.	11-5-54	Κ	• • •		4.5 gpm. 65°F	
14.33.21.444	Underwood No. 1 Cornett	1938	1370	_	3940		_	-			Oil test in Harding County, 2 miles No. of Quay county line, log	
											•	
	Mr. M. Burnes	1902	100	6	3930	90		Qc, To	P,W	D,S		
14.34. 1.141		1702	196	5	3940	96		Qc, To	N	N	Ca	
1.212			120.0	6	3960	104.	5 6-10-54		P,W	S		20
5.422			130		3880	118		Qc	P,W	S	Est vield 2.5 apm 64°F	
13.141	Pyle Ranch	0.1	100	4	3020	00	2024	್ಷ	D 14/		ec view	

TABLE 2. RECORDS OF SPRINGS IN QUAY COUNTY, N. MEX.

cation number: See explanation in text.

Altitude: Altitude of land surface at spring. Altitude interpolated from topographic maps or aneroid determination to nearest 10 feet.

Stratigraphic unit: Qal, younger alluvium; Qc, upland cover of older alluvium; To, Ogallala Formation; Ks, Cretaceous sandstone and shale; Je, Entrada Sandstone; Rc, Chinle Formation; Rsr, Santa Rosa Sandstone.

Location number	Owner	Name	Topographic situation	Altitude (feet)	Strati- graphic unit	Yield (estimated gpm)	Date	Use of water	Tem- per- ature (°F)	Remarks
7.30.15.432	_	_	Below cliff in gully	4720	То	Seep	8-25-53	None	_	Reported good quality and to have supplied 25 families 1910 to 1930
8.27. 6.430	H. G. Johnson	-	Side of cliff	5100	Je	2	11-2-55	Stock		Perched water, piped to tank
8.31.12.320	_	_	Stream channel	4220	Qal	2	4-21-55		_	
8.32.18.223			Stream channel	4220	Qal	5	4-16-55	Stock		-
35.114	Elder Dennis		Stream channel	4480	Ks	5	4-2-55	None		Spring at fault contact of Cretaceous and Triassic rocks
9.27.36.244	Mr. Hortenstein	Louisiana Spring	Side of cliff	5220	Ks	2	10-27-53	Stock	55	Chemical analysis in Table 3
9.32.24.322	Mrs. Hut Wallace		Stream channel	4200	Qal	1	4-8-55	Stock	58	
33.333	S. S. Hodges		Stream channel	4190	Qal	25	4-16-55	Stock	_	7.
9.33.24.312	Mr. Pierce	Hopkins Spring	Stream channel	4480	Ks	Seep	2-14-55	None		
10.33.14.212	Mr. Stams	Starns Spring	Side of cliff	4080	Qc	•	2-15-55	Stock		
10.35.32.422	Chapman Bros.		Stream channel	4020	Qal	-	12-1-54	Stock	_	Piped to tank
10.36. 8.233	Chapman Bros.	·	Steep slope	3920	٦ç	3	11-29-54	Domestic and stock	_	Piped to tank
18.224	Chapman Bros.		Stream channel	3970	٦ç	1	11-29-54	Stock		Wine.
11.33.29.211	Otto Collins	-	Side of cliff	3920	٦ç	0.5	2-18-55	Domestic and stock	51	Piped to tank
11.36.30.412	Grady Oldham estate		Steep slope	3950	Έc	0.5	11-5-54	Stock		

12.32. 6.213	Jacob Van Sweden	Cow Springs	Stream channel	3920	٦ç	10	3-8-55	Domestic and stock	_	Piped to tank
	Joe Hettinger		Stream channel	3920	Ŧς	10	3-4-55	Stock		
2.33.17.234	-		Gentle slope	4070	٦c	0.5	11-8-54	Stock		х х
2.34.22.241	Homer Koonsman	Blue Hole	Stream channel	_	ΤRc	100	11-6-54	Stock		
2.36. 5.231	A, C. Ward	Bive riole	Stream channel	_	TR sr	0.25	3-9-57	Stock	_	Chemical analysis in Table 3
3.32. 1.434		_	Side of cliff	3820	TR sr	1	7-26-57	Stock	_	Piped to tank
3.36.27.332			Stream channel	_	Te sr	30	3-8-57	Stock	62	Chemical analysis in Table 3
4.35.34.343	Pyle Ranch		•		Tesr	150	3-8-57	Stock	64	Chemical analysis in Table 3
35.311	Pyle Ranch	-	Stream channel			5	3-31-54	Domestic		Chemical analysis in Table 3
4.37.31.211	Ollie Mae Pyle	Coggin Spring	Stream channel	3580	TR sr	3	3-31-34	and stock	•	
5.34.30.310	Gallegos Estate	Sand Springs	Gentle slope	4110	Qc, To	300	6-3-54	Stock and irrigation		_
	E. A. Stringfellow	_	Stream channel	3850	Qal	100	4-7-54	_	-	
5.36.24.230			Stream channel	3840	Qal	50	4-7-54	Stock	_	_
5.37.19.134			•	•	To	1	5-22-53	Stock	_	Chemical analysis in Table 3
6.37.18.421	R. C. Bell		Stream channel	4130	10		J-22-00	J. 10011		

FINAL REPORT

Table 1: Log Availability from Exploration Wells in New Mexico

Map No.	Location	Well Name		New Mexico Nell ID No.	Availabile Logs
					Duillana (no 2051)
		Chapman No 1	C.T. Shook Shell Oil		Drillers (no 2951) Acoustic, Gamma
. 4	10.31.23	N. Pueblo No 1	Stiett off	14717	Caliper
3	10 21 25	N. Pueblo No 2	Shell Oil	14616	
٠,	10.31.23	N. PUEDIO NO 2	PHETT OIL	14010	Gamma, Caliper
Λ	11.36. 7	Endee No 1	L.B. Newby	_	Drillers (no 855)
	12.28.14		Miami Pet. Co.		
		Chapell No 1	Puretex Oil Co		•
					Porosity, Gamma,
					Caliper
7	12.29.18	Hoover R. No 1	Miami Pet. Co.	15850	Gamma, Laterolog
		Chapell No 2	Puretex Oil Co	. 14890	Neutron-porosity,
					Gamma, Caliper
9	12.32.11	Ute Anticline 1	National Oil C	co. 25563	Dual-Laterolog, Gamma,
				·	Acoustic, Neutron
10		Kimes No 1	O.L. Ledgerwood	xd 15851	Neutron, Gamma
11	12.32.11	Ulmer No 1	S.T. Silverste		Drillers (no 6249)
12	12.32.35	Tippen No 1	N.G. Penrose	15483	Gamma, Neutron,
		•	_		Drillers (no 6876)
13		No 1 Ranch	Marland		Drillers,
14	13.31.24	State No 1	Nucorp Energy	26194	Dual-Laterolog,
					Micro-Laterolog, Gamma
	13.31.25		Standard Pet.		Drillers (no 858)
16	13.32.32	Columbine St. 1	_		Gamma, Acoustic
17	13.33.15	USBR DH-3	U.S. Bureau Re		Drillers, Gamma Drillers
	13.34. 9	Olean No 1 Woods			
19	13.35. 2	N.M. Eng. DH-10	New Mex. St. I		Drillers Noutron
20	14.32.16	State No 1			Mud-Log, Gamma, Neutron Drillers
		Underwood No 1	Cornett	1000	
22	15.33.10	Arthur Cain No 1	J.A. Talley	0010	SP, Laterolog, Dual- Induction
22	15 22 17	Rodoral 1-17	Paul Haskins	21644	
		Federal 1-17		21643	Neutron-Porosity,
24	15.33.21	Conley Cain No 1	Contrey Masoc.	21043	Gamma, Caliper, Density
25	15 22 22	Arthur Cain No 2	Fdmonds. Peter	rs 6619	Gamma, Caliper,
23	13.33.22	ALCHUL CALLI NO 2	Damoraby 1000	0013	Interval-Acoustic
26	15 34 28	State No 1	Powers Wire	23774	Gamma-Gamma
20	2007140			·- -	Gamma, Caliper
27	16.33.27	1-X Olympic St.	Astro-Tex	19016	Caliper, Neutron-
~ (10170121				Porosity, Density,
		•			Gamma
28	16.36.36	State "CP" No 1	Humble Oil	13957	Gamma, Gamma-Gamma, SP,
					Laterolog, Dual-
1					Laterolog

HYDRO GEO CHEM, INC.

FINAL REPORT

LAKE MEREDITH SALINITY INVESTIGATION

Table 1: Continued

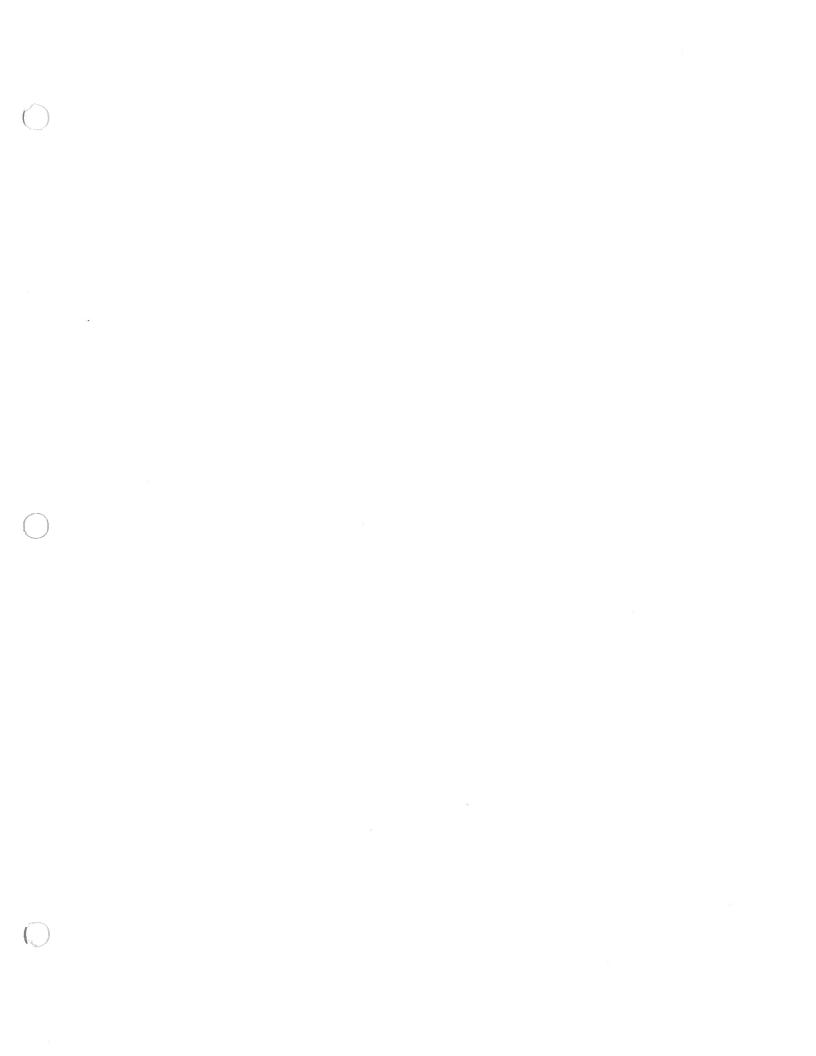
Map No.	Location	Well Name	Operator	New Mexico Well ID No.	Availabile Logs
29 30	17.32.24 17.36.28		Amoco Oil Humble Oil	24178 14899	Gamma, Neutron Gamma, Gamma-Gamma SP, Laterolog, Dual- Laterolog
31	18.32.14	"CM" State No 1	Humble Oil	14900	Gamma, Acoustic SP, Laterolog, Dual- Laterolog
32	18.34.31	"CK" State No 1	Humble Oil	14901	Gamma, Acoustic, Caliper
33 34	18.36.36 19.34.16	BDCDGU1836361K State "EL" 1	Amoco Oil Amoco Oil	23259 24126	Gamma, Caliper, Bulk- Density

iated northwest trending block faulting. Regional isopach maps constructed for the region by Foster and others (1972) and Grstavson and others (1982) indicate that many of the sedimentary systems of the Palo Duro basin can be extended into New Mexico along a synclinal trough defined by the Tucumcari and Cuervo basins.

STRATTGRAPHY

1. Paleozoic Section

The Paleozoic section in Quay and Oldham counties, shown in the stratigraphic column (Figure 4), is comprised of the Sangre de Cristo Formation and the Abo, Yeso, San Andres, and Bernal formations. These units lie unconformably upon the Precambrian surface. The unconformity is marked by arkosic sandstones often referred to as the granite wash, and it reflects tectonic uplift and ero-



Index to TAB 10: Well log data

HGC, 1984b, Figure 4

Spiegel, 1972a, well logs 1-3

USBR, 1984, Hydrology/Hydrogeology Appendix, Figure 1

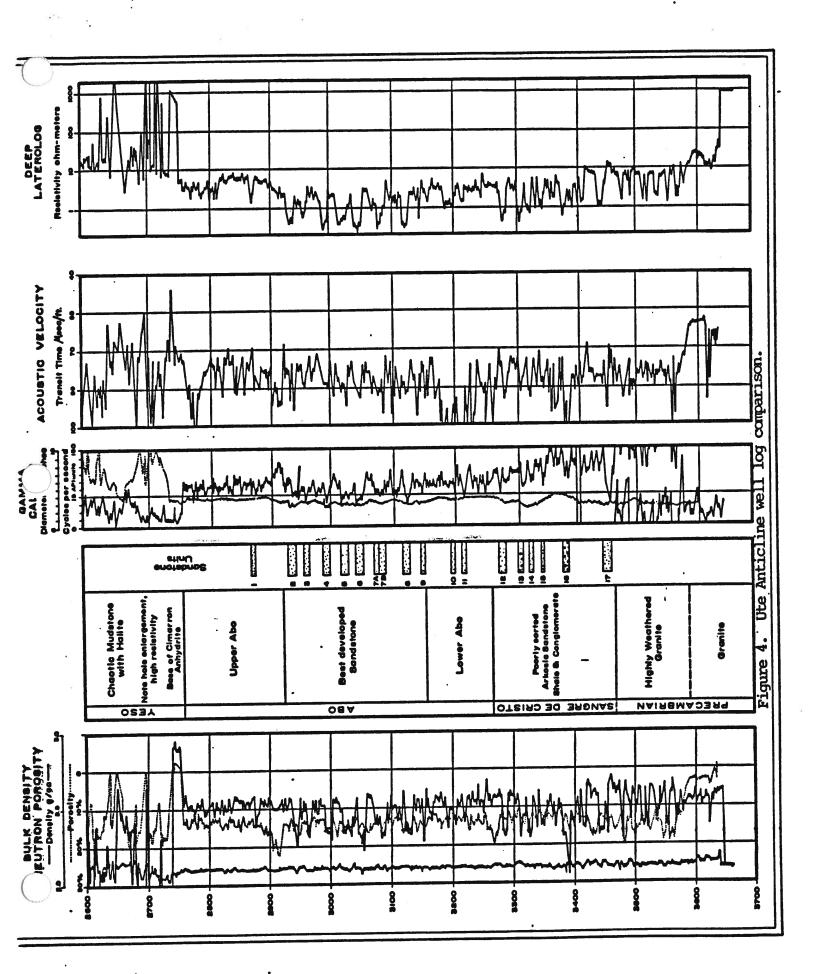
USBR, 1984, Hydrology/Hydrogeology Appendix, Figure 2

USBR, 1984, Hydrology/Hydrogeology Appendix, Figure 3

USBR, 1984, Hydrology/Hydrogeology Appendix, Figure 5

USBR, 1984, Hydrology/Hydrogeology Appendix, Figure 6

USBR, 1979, Appendix D



westward between Logan and bonafide Santa Rosa outcrops in the Pecos Valley.

CONCLUSIONS

The solution, essentially as proposed by Trauger in a memorandum of November 4, 1971, and concurred in by the writer, is to redefine Gould's Trujillo Formation to include only the lower sandstones north of Glenrio (equivalents of the Logan Sandstone) and to assign the upper beds of Gould's Trujillo to the Chinle.

The Santa Rosa in the Pecos Valley probably is equivalent to the combined section of Tecovas and Gould's lower sand-stones of the Trujillo, but the sandstone members of the Santa Rosa may not be physically continuous into the Logan Sand-stone and are definitely not equivalent to the "canyon sand-stone" at Conchas, or the sandstone at Sabinoso dome.

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WELL LOG 1

SAMPLE DESCRIPTION OF DUNES #10 TEST HOLE

Section 2, T. 13 N., R. 35 E., NMPM (Elevation 3585 ft., on east slope of knob on north bank) E. A. Chavez, August 1957

Interval Alluvium:	Description
0 - 5	ss., f.g. to v.g.f., lt. tan to buff (probably wind blown dune sand)
Tecovas:	

- 5 10 ss., v.f.g. It. tan, micaceous, subrounded grains
- 10 25 Cored: Recover 5.4'-All ss., lt. tan to white, clean, soft friable but consolidated, porous, micaceous
- 25 35 Cored: Recovery 5.7'-All ss., as above
- 35-45 Cored: Recovery 4.55'-All ss., as above but with occasional laminae of micaceous gray siltstone.
- 45-55 Cored: Recovery 4'-All ss., It. tan to white, clean, friable but consolidated, porous, with occasional inclusions of calcium carbonate forming incrustations and tiny clacite filled cavities.
- 55-65 Cored: Recovery 3.25'-(Top) 2.55' is ss., lt. tan to white, hard compact, slightly calcareous grading to 0.7 of conglomerate, small pebbled, gray to yellowish gray, vy. calcareous with large fragments of gray dolomite at base.
- 65 70 ss., m.g. to v.g.f., white to lt. gray, clean
- 70 75 ss., v.g.f. tan, argillaceous, slightly micaceous.
- 75 80 ss., f.g. to v.f.g. buff to bwn., slightly micaceous.
- 80 85 ss., v.f.g., bwn., rounded grains
- 85 90 ss., as above but vy. clean, well sorted grains
- 90-100 Cored: Recovery 1' 10"-ss., white, m.g., vy. porous and clean w/1" lens of gray siltstone near bottom.
- 100 105 Cored: Recovery 2'-ss., as above
- 105 110 No sample
- 110-125 Cored: Recovery 1' 9"-Top 18½" ss., as above 2½" Dk. gray soft, sticky shale
- 125 145 ss., f.g. to v.f.g. lt. tan to white, micaceous, rounded grains.
- 145 150 ss., v.f.g., tan, micaceous, subangular grains

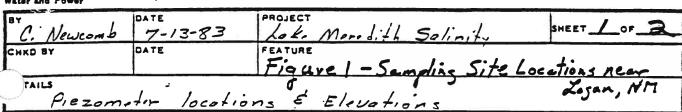
Alibates:

- 150-155 Cored: Recovery 1'5" Lt. gry, crystalline and vy. tight ls. w/occasional intrusions of soft white pyritic gypsum. Vugular in places.
- 155 160 Cored: Recovery 1' 10" Lt. Gry. Is. as above grading to dolomite, sporadically vugular.

Quartermaster:

- 160 161 Shale, gray
- 161 170 Ss., v.f.g., soft, friable lt. bwn., argillaceous, and slightly micaceous
- 170 175 Ss., as above with a lense or red and gray variegated shale at 170'
- 175 180 Ss., as above grading into
- 180 183 Cored: Recovery 2, 1½"-Hard, well consolidated sandy red shale with occasional streaks of blue-gray coloration.
- 183 190 Siltstone, brown, arenaceous
- 190 195 Ss., brown, v.f.g., rounded grains
- 195 209 Ss., as above and argillaceous
- 209 210 sh., red w/blue-gray variegations
- 210 220 Ss., v.f.g., brown, argillaceous and slightly carbonaceous
- 220 230 Siltstone, reddish tan, minutely arenaceous
- 230 235 Siltstone, reddish tan, strongly arenaceous
- 235 240 Shale, red with white variegations

WELL LOG 2		TRUJILLO	O SANDSTOI		ted, this paper):
U.S. Corps of Engineers Core Boring H-4		681	700	19	Hard sand
North dike of Conchas Dam, New Mexico		700	705	5	Shale (blue)
(from U.S.C.E., 1936)		705	724	19	Hard sand
OSS EL. 4173.6 3" diamond bit	0	724	752	28	Sand and green shale
2' Sandstone, gray, griffy, apparently is massive		752	783	31	Hard sand & shale
rock under deep cover This specimen is	F	783	801	18	Hard sand
ols integrated H is rotten rock Shale, reo, mostly sandy, in this layers. Mostly	20				
06 clay between 4150 8 and 41500; also between		TECOVAS	FORMATIC	N	
03' 4/463 and 4/444.	-				combined equal Santa Rosa ss.):
1463 and 41444. Sandstone gray, very hard, firm rock. Shale, pink, hard, very sandy, massive. 14' Sandstone gray, hard, fine grained Some clay in upper half. Lower part slightly porcus.	40	•			
20 /		801	842	41	Sand & shale
33' Sandstone gray, hard, fine grained Some clay	F	842	876	34	Sand & red & green shale
30' In upper half. Lower part slightly porous	60	876	920	44	Sand & red & blue shale
		920	940	20	Sand
	- -	940	955	15	Sand & shale
Limestone.gray, massive, bedded, no porosity	80 2				
Sandistone, gray, gritty, massive, slightly porous	- I	UPPER M	EMBER OF T	HE QUAR	RTERMASTER FM (Bernal fm.):
CANYON SANDSTONE FORMATION EL. 4083.9	- z	955	1022	67	Anhydrite
37' RED SHALE FORMATION	100	1022	1037	15	Red shale
Shale, pink, thin bedded, hard, massive, contains		1037	1050	13	Broken Anhy
c4 very fine sand Firm rock Shale, (clay) red, soft, plastic when wet. A	120 8	1050	1098	48	Green shale, sand breaks
soft formation	120 8	1098	1130	32	Broken shale
RED SHALE FORMATION EL. 4022.3		1130	1131	1	Crevis
06' FINK SHALT SANDSTONE	ő	1131	1135	4	Sand rock
16 Shale, pink, hard, thin hedded, massive. High %	140 E			•	
09 Shale clay), red, suft, plastic when wet	-	ALIBATE	S LENTIL O	F THE QU.	ARTERMASTER FM.
39' Shale, pint, very hard, massive, contains very	9 30	(San Ar	ndres (m.):		
24' 1/PE 5870	1	1135	1145	10	Dolomite
PINK SHALY SANDSTONE EL 4002.0 ARTESIAN SANDSTONE	160	1145			
Sendstone, oray, gritty, posous, massive		1155	1159	4	Hard sand (Dolomite)
1.5 Sandstone, arey, gritty, prous, massive Gray		1133		•	(======================================
07 shale break at 39.72? to 39924.	180	LOWER M	MEMBER OF	THE QUA	RTERMASTER FM. (Glorieta ss.):
ARTESIAN SANDSTONE EL. 3985.2	į.	1159	1179		Sand hard
CLAY SHALE SANDY CLAY	200	1179	1184		Broken sand
	200	1184	1239	55	Hard sand
GRAVEL CONGLOMERATE	1	1104	.233	33	
SAND Z ADOBE CLAY		YESO FM	1:		
LIMESTONE					
		1239	1312	73	Broken sand
)		1312	1351	39	Sand
WELL LOG 3		1351	1364	13	Hard sand
Waggoner and Wharton, Upton #1, 1946		1364	1380	16	Brkn sand
NW Cor. Sec. 25, T. 18 N., R. 26 E., NMPM		1380	1407	27	Sand & lime
(Elev. 4875 ft., reported)		1407	1437		-
Log from files of NMOCC		1437	1451	14	Broken sand & shale
Log nom mes or misses		1451	1494	43	Sand & shale
FROM TO THICKNESS		644005	DE COICEO	544 4440	MACDALENA CROUR UNDIEE
IN FEET		SANGKE	DE CRISTO		MAGDALENA GROUP, UNDIFF:
		1494	1517	23	Red rock
MIDDLE SANDSTONE MEMBER OF THE CHINLE FM (Top et	roded):	1517	1539	22	Red rock & sand
0 20 20 Surface, boulders & hard sa	nd rock	1539	1566	27	***
20 35 15 Hard sand		1566	1632	66	Red rock
35 45 10 Broken shale & sand		1632	1663	31	Red rock sandy
45 70 25 Red sand hard		1663	1713	50	Red rock
		1713	1720	7	Red & blue shale
		1720	1743	23	Anhy & Dolomite brks.
		1743	1768	25	Broken Anhy & shale
90 105 15 Red rock, blue shale		1768	1780		<u> </u>
105 122 17 Broken sandy shale		1780	1820	40	Anhy & shale red
122 158 36 Shale & sand	¥2.	1820	1853	33	Sandy shale, red & blue, with
158 200 42 Red rock and sand					andhy breaks
LOWER SHALE MEMBER OF THE CHINLE FM:		1853	1873	20	Red rock
		1873	1900	27	_
200 290 90 Red shale	le see a less	1900	1941	41	Brkn formation gyp, gravel, shale
290 305 15 Red shale with green shale	DICARS	1300	1771	71	(granite wash)? Show of gas 1937
305 340 35 Green shale breaks	ala a	1941	1994	53	Granite Wash
340 360 20 Red & blue shale & rock bi	1K2.	1941	2022	28	Granite wash lime (gas bubbles
360 385 25 Sandy shale, hd.	المسط	1774	2022	20	on pit 1995-2002)
385 431 46 Green shale, blue shale and	i Sano	2022	2000	58	Granite wash
431 490 59 Grey shale, hard sand		2022	2080		Shale, gyp, silate, granite wash
490 520 30 Blue sandy shale		2080	2100	20	
520 529 9 Hard red rock broken		21 00	2114	14	Granite wash
529 550 21 Shale, sandy shale	X:	2114	2132	18	Granite wash, hard
550 575 25 Bentonite green		2132	2148	16	Granite wash
575 590 15 Bentonite green		2148	2158	10	Hard granite wash
590 650 60 Green shale sand breaks		2158	2171	13	Granite Wash
650 681 31 Sand & shale			2171		T. D.



Piezometa # 0 5:1. 103° 27 36" - 35° 20' 40"

Location: 13.33.21.1224 @ Toe of UTE DAM Elevation Ground: 3682.7 Elevation Top Pipe: 3685.48 W/Cap off

Piezomiter # 1 5% 103° 25 17" - 35° 21' 12"

Location = 13.33.15.4112 A Elevation Ground = 3668.9 Elevation Top PVC = 3672.59 "0" on Staff Gage = 3665.62

Location = 13.33.15.4112 B Elevation Ground = 3668.8 Elevation Top PUC = 3672.75

N pow of Pivil 500 N NO E

Piezomotin = 2 site

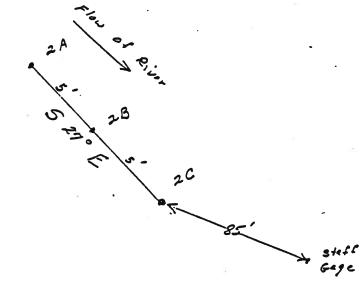
Location = 13.33.14.1332 A Elev. Ground = 3668.7 Elev Top PUC = 3672.55

Location = 13.33.14.1332B Elev. Ground = 3668.7 Elev. TOP PVC = 3670.41

Location: 13.33.14.1332 C Flev. Ground: 3668.5 , Tos PVC: 3672.73

0" on Steff Gage: 3663.32

103° 24' 52" - 35°2" 14"



0" on USGS Staff Gaox = 3665.61

COMPUTATION SHEET

ВУ	DATE	PROJECT	SHEET 20F 3
KD BY	DATE	Figure 1 (continued)	
DETAILS	-		

TW-L

103° 24' 50" - 35° 21' /2"

13-33-14-1333

Location = N 15852525 E 774291.3

Elev. on - 34tor ring = 3674.01

white point on west side

Plezomoter # 3 5:40

103 23 30 - 35 22 00 7 640

Lacation = 13.33.12.3214 A

Elev. on Ground: 3655.1

Elev. Top PUC = 3658.38

Location = 13.33.12.3214 B

Elou. Ground: 3655.1

5/ou, Top PUC = 3658.08

3 A

on Staff Gage = 3651.26

Piezomotor # 4. Site

Location = 13.33.12.4412 A Elev on Ground: 3453.7

Elev Top PUC = 3658.77

Location = 13.33-12.4412 B

Eleu. Ground : 3653.7

Elow Top PVC = 3656.68

1030 221 5811 -

No Stoff Gage

1030201 221- 35023 20 Piezomiter = 6 Site

Sheff Gage = 3631.98

Location = 14.34.33.3324 A Elev. Ground = 3638.0

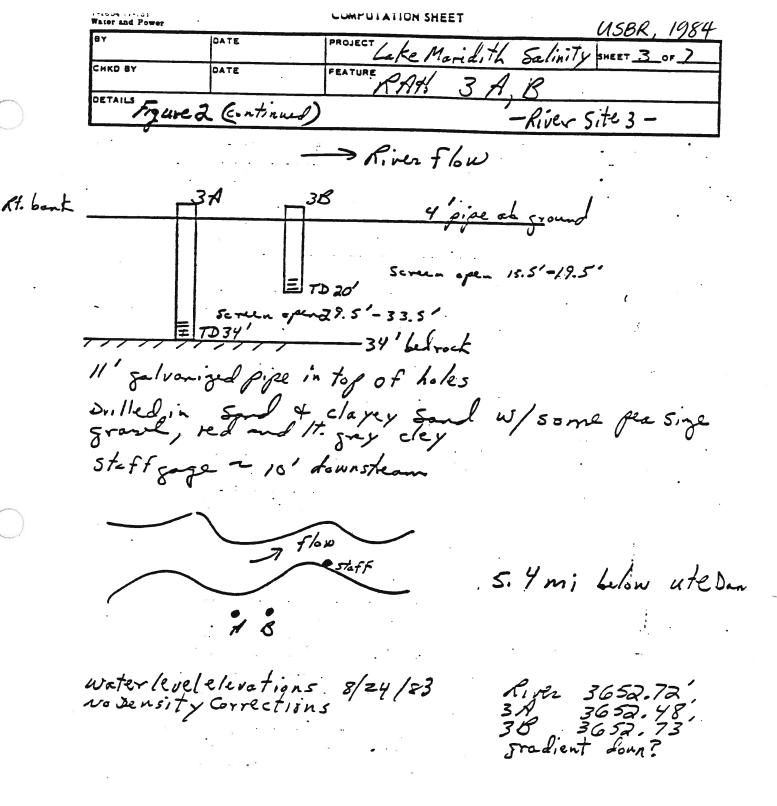
Eleu. Top PUC = 3640.99

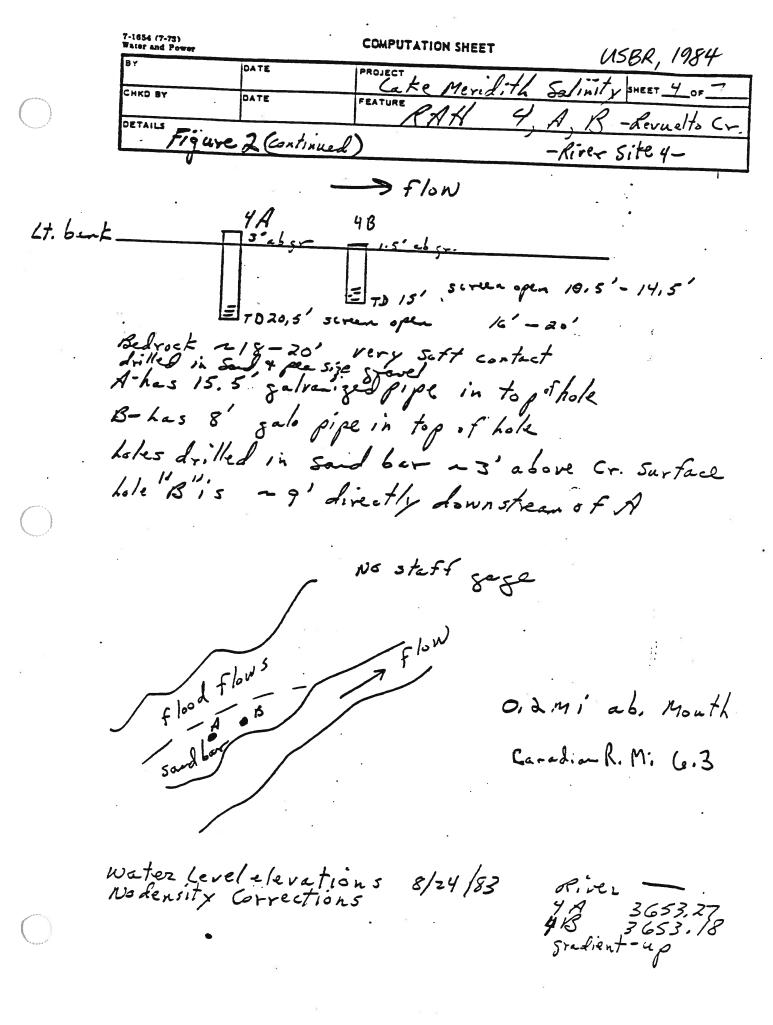
ation = 14.34.33.3324 B when Ground: 3637.9 Elev. Top Puc = 3640.70

Location: 14-34.33-3324 C Eles Ground = 3637.6

1.0

gradient up





Water and Power			USBR. 1984
ВҮ	DATE	Lake Meridith Sel	
CHKD BY	DATE	FEATURE RAH	<- 1 0 11
DETAILS FI	ure 2 6.01	inued) -	River Site 5 -
Not Co	ompleted	as of 6/7/8=	3

Luy bridge

and they Gridge

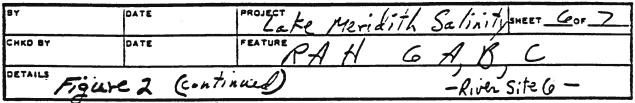
and they Gridge

and May Gridge

and Mouth

uses.

cable way



> flow

68 60 GA Ist bank 国力3211 16.51 - 20.5 & Bit 36,5'-30,5' 45.5'- 47.5' Screen open

Bedrock \$ 52

Bits in bollom of each hole drilled in clayer sand at surface, gravel at 10', Soul w/cley lenses gravel on top of bedrock

Il ft galvanized pipe in top of hale

> f/ow 5 ta 5 f

Approx 6 ab river surface

9.9 mi below ute sam

8/24/83 Water level elevations No Density Corrections

River 3631-99 3632.20 3632.24

gradient?

reasonably good road to w/in 50' of piez.

7-1654 (7-73) Water and Power

BY	DATE	Lake Meredith Salinity Control SHEET_OF_
THEO BY	DATE	Observation well Locations
DETAILS FIZ	ure 3 - 06	servation well Locations Near Locan NIM

well No.	State Plane North	Cordinates : East	Elevation
DH-1 DH-2	1585226.9	774266.3	3674.5 (bolt)
DH-3 FW-1 POW-1 OW-2 OW-3 OW-4	15859 02 15852525 1585178.6 1585081.1 1584830.4	770028 774291.3 774245.7 774153.6 773931.6	3655.72 (Top Spisofin Lend Surface) 3781.0 (Land Surface) 3674.01 (Top outer ring-Lend Surface) 3675.9 (pipe) 36828 (pipe) 76 3673.0 (Land Surface) 3676.5 (Land Surface)

*1 foot of pipe has been cut off since elevation was determined

DH-2 Cocation - 13-34-17-1342 or 103° 22' 32", 35=22'-10"
DH-3 Cocation - 13-33-15-3124 or 103° 25'40", 35-21-05"

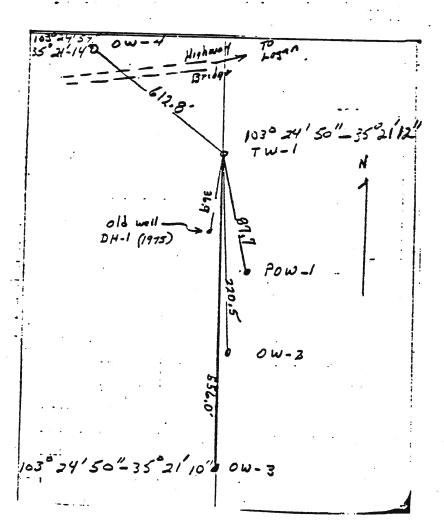


Figure 5 GEOLOGIC LOG OF

					EUL						LE N		JM * 3 94(1 . 1 nr 6
													STATE MEN MEXICO
													. ANGLE FROM HOREZ 98.0 DOM
BEOUN 9/17/83 FIN	ISH€D.	.941; ; 8	? .	DEPTH	10 B	OP JC	K . !!!	 .	.TOTAL	. DEP	TH	569.	5. BEARING
DEPTH TO WATER		9/9/81		LOCGE	n RY	5111	LEY S				PEVI		BY JOE JACKSON
•	-	המוצאנה	1104	100 AHIL	WITH A	WAL .	-	ğ	٠,	CL ATT	FRYALS		
HOLES		PTM ETI	55		113371 30-28364 111-3632610	223	1 5	1000	3772.	<u>u</u> -	EVATIONS 17671	22	CLASSIFICATION AND
7.	FRON		01 aPC 128 1 1 NCM(\$1	867	225	1001	1100 1233 A		H 430	CEPTI	100	7 21	PHTSICAL CONSTITION .
	15. 6.	10	0-		8	3	5 -	22	8-	30	3	<u> </u>	•
DRILLED USING SOM DRILLING RIG, RLAN MUMP (24 OFF MAXIMUM)		1		·							3779.4	1	0-11.0 FT.: QUATERNARY ALLUVIUM
CAPACITY) AND CURE DRILL OPERATOR FROM RHANTLEY	.:							100	- :		3777.0	1	.8-1.6 FT.: GRAVELLY SAND.
PROJECT, NEW MEXICO:	1			:				l			ł	١.	NAMEROUS CALICHÉ FRAGMENTS AND PERMIES6-9.8 FT.: SANDY GRAVEL MITH CORRLES.
USED 3-7/8 INCH ROCK BIT 0.0-16.2 FT.									10 -		3770.0	1	.0-II.0 FT.: SILTY SMO.
USED NO DIAMOND BIT 16.2-3.55.9 FT. USED NO CARRIDE BIT							ŀ		. :			1	.0-519.0 FT.: IMIASSIC DOCKUM GROUP,
335.9 178.2 F1. DRILLED 350.0-378.2	j										3767.0	'	1.0-14.0 FT.: SANDSTONE. SILTY, MICACEOUS, MEDIUM ORAIDED, BROWN,
AND DRILLED 370.2- 569.5 FT. USINO NO DIAMOND RIT. 10P	1										3762.7	1	4.0-18.3 FT.: CLAYSTONE. CLAYEY, MCD TO RED-DHOMM, MITH INTERNEDS OF
OF ROCK DEPTH BASED	1						1		- 20				RED-HRUBH SILISIDNE (1) 1-17.7 FT. AND FINE GRAINED, MICACEOUS CROSSDEDDED SANDSIDME (16.2-16.7 FT.) AND (17.7-17.9 FT.), SIRONG
MD CUITING. MATER LOSS DURING	1											١.	PEACTION WITH HCL. TAM. 8.3-30.5 FT.: SANDSTONE.
DRILLING: INTERVAL IFT.I PERCENT	İ								;		3750.2	'	STITE MICACEGUS. FINE-GRAINED, CROSSRED- DED. LIGHTLY TO MODIFICATELY CEMENTED. CHE
48.0- 50.6 50 122.0-150.0 40							l				3.33.6	1:	MARKER RIOM CRUSHES SMALL PIECE (CONF.) CORE STICK 1.7 FT. STROND REACTION WITH HELE TAN TO BROWN, E. PT YELLOW FROM 29.0-10.5 FT
273.0-273.5 100 273.5-106.0 40 327.0-3-2.5 50									- 39		3750.5		28.8-26.89 FT.; CLAYSTONG. GREENISH "RAY,
10FF CHE CASING TO 362.0 FT.1 350.0-370.2 NO												1.	27.3-27.4 F1.: CLAYSIONC.
430.0-450.0 86										1.		١,	RCD. 0.97.0 ft.: SHALC.
DRILLED WITH CLEAR WATER EXCEPT IN INTERVALS AS FOLLOWS:	1												SANDY, MICACEOUS, SLIGHTLY FIRSTLE TO BLOCKY STROND PEACTION WITH HELL PREDOMINANTLY HED WITH CREENISM GRAY LAYERS AT WILT-MARGET.
E-2 MID DEPTHIFT.1 2 GAL. 61.0 -5 GAL. 319.8	1										.		AMD 93.0-95.0 FF. WITH SOME CREENISH GRAY YELLOH BROWN MOTTLING AND BANDHAD. CONSIDER-
5 DAL. 350.0 5 DAL. 370.0	1											١.	ABLY LESS SAMD IN GRAY COLORED INTERVALS.
50 LDS REVERT 370.2 FT. 5 GAL. 529.9 8 GAL. 568.2				i					:	-	3739.0	'	7.0-53.6 FT.: SANDSTONE, FINE TO MEDIUM GRAINFD, MICACEDUS, CROSSNED- DED, MODERATE TO STHOND REACTION WITH INC.
HOLE REGAN CAVING	1		`	ŀ.					- 50			'	70 STOMES TO VERTICAL FRACTURES WITH HITM AND MANSANSE STAINING FLW THIN BYOS IN THE 50.6 FT. CONTAIN ROURIED TO OBLOWD FHACH HIS
AT 375.0 FT. IN PED MEDSTONE AND DRIEN SHALE FROM	1												OF THOMM AND GRAY CLAYSTONE (-) INCHI. I AM. TO BHOMM.
P97.0-3%0.0 FT. AT 419.3 FT., MOLE CAVED HACK TO APPROX.			l						-		3727.4	۰	3.6-65.4 FT.: SHALE. CLAYEY, SLIGHTLY FISSILE TO BLOCKY, MODERATE-
SGS.0 F1. FACH FING ROOS WERE PUBLICA.				·									ET MELL-CONSOLIDATED. SLIGHT MEACTION HITM HCL. COME STICKS UP TO 1.1 FT. IN LEWIN BROWN TO PCODISM MYSME WITM THIN GREENISM
NTER CASING SET 10 STA. 0 FT., HOLE DEVIATED FROM PRE-									- 60 -			١.	GRAY LAYERS AND HOTILING. 5.4-92.3 FT.: SANDSTONC.
VIOUSLY PRILLED HOLE, CONSOLID- ATED CAVING AND												"	GRAINFD. THIN SANDY SILTSTONE LAYERS
OHMATION ROCK	1									-	3715.8		THROUGHOUT. CAPHONACEOUS MAILPIAL AND MICA ON BEDUING PLANES. 50-70 OFGREE FRACTURES 72.2-77.5 FT., SLIGNILY CEMENTED. HAK 10
HERE PLCOVERED FROM 362.0 TO 403.3 FT. FORMA-			ĺ	-									STRONG PEACTION WITH MCL. GRAY TO DHOWN HITH LIMONITE STAINING AND SPOTS.
TION ROCK MAS CORED FROM WOS.3									70		1		83.2-09.8 FT. SMALE. CLAYEY. FISSILE. HELL CONSOLIDATED.
10 369.5 F1.	l.							i					CORC STICK 8.8 FT. LONG. GRAY.
			1							1	ii i		MEDIUM GRAINED. THIN TO MEDIUM BEDDED. NEAR VERTICAL FRACTURES
ø													THROUGHOUT, BUT STRONGLY TRACTURED 87.5-28.9 FT. SLIGHTLY CEMENTED. YELLOM.
	-0								- **				2.3-118-9 FT.: SMALE. CLAYEY. FISSILE. THIN LAYERS 10.7 FT.
									Ė,		3897.8		THICK! OF CHIENISH GRAY SHALE AT 92.3 FT.,
. •											3895.0		THIN LAYERS OF FINE-GRAINED, WELL CLHENIED CHAY SANDSTONE AL 93.4 FT. AND 97.4 FT. STOOMS REACTION TO HEL DECONING MODERALE
											3696.9		BCLOM 116.0 FT. RCD DROWN. 18.9-199.7 FT.: SANDSTONE.
-											3698.7	Ι'	SILTY. MICATEOUS. CAMPONIZED WOOD LAYERS AND LAMINATIONS OF MICA AND CAPPIONACEDIS
•						183			E :				MATERIAL HITM ASSOCIATED PYRITE AND CHALCO- PYRITE ON DEDOING PLANTS. FINE GHAINLU. PROCRATELY TO SCIGNILY CEMENTO. SCIGNI TO
									ŧ :				NO REACTION WITH HCL. NEAR VERTICAL FRAC- TUPES AT 121.5-125 0 FT., 132.5-133.0 FT.
													AND 145.0-146.0 FT. 45 DEGREE FRACTURES
COMMENTS: SET 14.0 FT. OF 4 INC						,			1	44T 10H		RAL F	ORMALAS USED TO COMPUTE PERMEABILITYS
MA CASINO TO 18.5 FT. FT. MA CASINO TO TC. 9/13/03: OPPUTED HIN	9 FT. 0	N R/31/8	13. 9	/11 THRO	UON				•		G Logg	7	MEN L ORGATER THAN OR COURL 100
I FF. SIMD FRIM 417 FINCH DIAM TER PYC SCE	NEW AND	ff. <u>SC</u> 1 371.0 f	1 49.5	FT. OF	1-1/8				,	k • 20	0 S100-1	<u>.</u>	DACH L LESS THAN 10 AND OREATER THAN OR LOUAL .
1/P INCH DIAMFIER IR A SAND PACK IN HIRE FRO	MR PVC M 417.5	TO \$17.5	51. 571.	PLACED	3.Q F1	37			PC PPC A	81L ! ! Y	VALUES SHO	-	C COMPUTED FROM THESE THEORETICAL FORMARAS AND LOSSES AND OTHER FACTORS THEOREMS IN THE TESTING.
BENIENLE 358.5-161.5 381.5 FT. 10 090/MO E									THERES	ORE.	HESE ARE NO	I IN	E PENERAL! ITIES, MERLLY MELATIVE VALUES.

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HIGHES INCREMENTATION CLASSIFICATION AND CLASSIFICATION AND PHYSICAL-COMMAND PHYSICAL-COMM		TO HATER			·•/#3		DECE	אח ר	SHIB	LET 54	Antx		I Atres	PFV		D. BY JOS JACKSIN
298-1 297-1		NOTE		or P	7 <u>1,510</u> 244	1104	المالك	49 <u>H</u> 74	MIN.	10 mg	ין ככשר. גףל	7 7	-17	(Tyat 5	13 FC4 51 19.6	CLASSIFICATION AND PHYSICAL COMMITTON
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2973.1 30.00 State March State S														-		SAMPY, COPPRISH REDUCTION SPOTS, STORY
100 MIN SET, ON NEW COMPLEX CELLS BY 17-19. KE 16-5																AND SHALE, HARD, STRONG REACTION HITH HEL.
200 200					*							210				S39.5-941.1 FT.: SANTSTONE. GREENISH REDUCTION SPOTS. FINE-GRATHED. HOSTLY ROUNDED GHAINS. HARD. STRONG REAC-
Total Tota															$\ $	SALLY FOR FRACHEMIS LIMESTONE (-3/4").
299 199			:										•			MELL CONSOLIDATED. GREENISH IN SALHON-MED.
### 200 20												229			$\ $	MISH CRAY STALE. FEW CACLIFFIELD HALL.
COPE MTS: WE CANNOT TO A THE WAY OFFICE CASHING DATIFEST. SET 18.0 WE CANNOT TO THE ST. OF STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE ADDITIONAL 23.2 OF THE CASHING TO STREAM. THE STREAM OF THE STREAM. THE ADDITIONAL 23.2 OF THE CASHING THE THE STREAM OF THE STREAM. THE STREAM OF THE STREAM OF THE STREAM. THE STREAM OF THE STREAM OF THE STREAM. THE STREAM OF THE STREAM OF THE STREAM.		•				·					1.			3557.2	$\ $	MELL CONSTRUCTOR ON THE NAME OF THE PARTY OF
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COPENIS: ST 10.8 FT. OF 10 MEM SUPPLET CASIND 8/17/83. SCI 10.0 ST 10.8 FT. OF 10 ME																•
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STT 10.8 FT. OF 0 INCH SUPPLICE CASINO 8/17/83. SET 10.8 MI CASINO TO 18.5 FT. ON 0/18/85. SET ACDITIONAL 303.5 FT. INA CASINO TO 38/2.8 FT. ON 8/31/83. 9/11 INDUCHO 0/13/85: CONVITED INTE FROM T.O. TO 100-10.9 FT. PLACED 1 FT. SAMP FROM NET 9-NERG FT. OF 1-1/8 E BLM R BLM R BLM ORCATER THAN OR COUAL 10/10/10/10/10/10/10/10/10/10/10/10/10/1												F	}	3:02	.0	
SET 10.8 FT. OF 0 INCH SIRPACE CASING 8/17/83. SET 18.8 MI CASING TO 18.5 FT. ON 0/18/83. SET ADDITIONAL 33.5 FT. INA CASING TO 367.8 FT. ON 8/31/83. 9/11 IMBOUGH 9/13/85: OFFICE OF THE TO. TO 0/18/5 FT. PLACED 1 FT. SAND FROM 0/17 5-0/18/5 FT. SET 00.5 FT. OF 1-1/8 R = 0 R =		r MTS+										[XPL	AMATI	OHSI	CENT	RAL FORMULAS USED TO COMPUTE PERMEABILITY
1 FT. GAMD FROM WIT 9-WIR.9 FT. CET W9.9 FT. OF 1-1/2 R = Blab 2 ORCATER THAN OR COURL P		SFT 19.8 FT		F. ON R	/1M/AS.	561	1001110	MAL 393	6.0		(4		x •	O L	•1,	F MCH F LUCATER THAN ON COURT 104
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SE	15: T 14.8 FT. CAT140 10	OF 4 18	CH SINT	ACT CAS	140 8/ ST 1 A	(7/83.	SET 10	.0			ŧ .		0 L40		-91 JAUSS RO MANT ROTADRO J HOM
FI	CA" (NO 10 . NM CASIM 13/81: OR	n 10 96-2	.0 F1.	ON 8/31	/83. '	9 /11 1m	MUK 4UM	-							
1	FT. SAMO F CH BLAMFIF	ROM 617. R PVC SC	9-410.5	FT. SI 0 371.0	ET 49.4	3 FT. 0 CHEDULE	00 1-						OLH SION		MACH E EESE THAN 18- AMG OREATER THAN OH EQUAL -
1/	MD PACK IN						<u>-</u>				1200	-81517	A ANTOE 2	,,	ARC COMPUTED FROM THESE THEORETICAL FORMULAS AND LOSSES AND DIMER FACTORS INNERTHE IN THE TEST

COPPE NIS:

TRITIST

SET 18.0 FT. OF 9 INCH SURFACE CASING 8/17/83. SET 18.0

MA CASING TO 18.5 FT. ON 8/18/83. SET ADDITIONAL 30.5.9

FT. IMA CASING TO 36/2.8 FT. ON 8/31/83. 9/11 IMPOINTM
9/13/83. 9/11 IMPOINTM FROM F.D. TO 0418.9 FT. PLACED

1 FT. SAND FROM 917.5-918.5 FT. SET 90.3 FT. OF 1-1/2

IMCM DIANTIER PVC "CEPTEN AND 371.8 FT. SCHOOLE 80 1
1/2 IMTH DIANTIER REAME PVC TO 917.3 FT. PLACED

SAND PACK IN HOIF FROM 917.5 TO 961.9 FT. PLACED 3.8 FT.

BE MITCHIEL 39/4.5-161.5 FT. AND MEAT CEPTENT GROUT FROM

301.5 FT. TO OMITIBE LEVEL, PLACED 9 FT. OF 2 INCH STEEL

PROTECTIVE CASIND MITH 2-FT. STICKUP.

EXPLANATIONS: OCHERAL FORMULAS USED TO CONTUIT PERMEABILITY

9 SOTH Leg

MEN : ORGATER THAN OR EGIAL 19-

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GACATER THAN 10. AND GALLE

PERMEABILITY VALUES SHOWN AND COMPUTED FROM THESE THEORYTICAL FORMAL AS AND DO NOT CONSIDER SYSTEM HEAD LOSSES AND OTHER FACTORS THAN MY IN THE TESTING. THEORYTICAL VALUES. P . PACKER CS . CASING CH . CEPENT

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USBR, 1984

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					•						y, Study STATENew.Mexico
IOLE NO YTTA	ATION	35	21	13"	E 103°	24' 5	10ROU	ND ELEY	3680	A.M.S	L DIP (ANGLE FROM HORIZ.)
COC DEGUN 6/24/75 FIN	ISHED.	.6/.30	J.7.5	DEPTH	OF OVE	RBURDÉ	N .3Q.	Q'	DE	PTH. 3	>6.º BEARING
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LEVEL AND DATE MEASU	RED.4		An 20.				GED B		,		
NOTES ON WATER	TYPE	CORE			LATION	TESTS		EVA.	DEPTH (FEET)	ا ي	CLASSIFICATION AND PHYSICAL CONDITION
LOSSES AND LEVELS,	SIZE	80	DEF (FEI	TH ET)	ł	RESSUR	LENGTH OF TEST	교로	2E	GRAPHIC LOG	CLASSIFICATION AND PHYSICAL CONDITION
CASING, CEMENTING, CAVING, AND OTHER DRILLING CONDITIONS	OF	REC	TROM		LOSS	S S	M.		1	5	SE PRINCAL CONDITION
DRIEFING CONFILIONS		(%)	PROM (F. Cs. or Cm)	TO	(G.P.M.)	(P.S.L.)	(MIM.)		1		3)
Drill Rig:	7-7/	4									0' - 30'
Failing 1500	Rock		1	l					3.0		Quaternary Alluvium
	Bit 7		1	ĺ	1] -	0.0	
Casing: Set 30.7" of 6"	- 1			ĺ	1				-	0 .0	0' - 3' Sand, poorly sorted. Contains some very fine gravel, some fine to very
casing; upon hole	10-		1 1	l		}				0. 0.	fine saud, mostly medium sand. Strong
completion o" casin	B .1	1	4	1	l				10.00	5.=¢	HCL reaction. Contains some quartzite,
pulled, set 31.5' o		e 1			1	l				-0-0-	and feldspars, all less than 5% of tota
4 ⁿ casing and cemen to ground level and		. [14.0	o.°	Mottled reddish color.
relded cap on to se					1	l			1 :		3' - 10' Sand, coarse and poorly sorted
8	20	1		1	1				20-	. 0.	Contains some very fine gravel, some fi
Drill Fluid Loss:	1	.	1	i	1				:	0.:0	to very fine sand, mostly medium to coa
01 - 1061 0%	1	1		i	1	ļ			:	0	Strong HCL reaction. Contains some qua zite, opal, mica, and chalcopyrite, all
106' - 116' 10% 116' - 256' 0%	4				1				1 -	. 0	less than 5% of total. Mottled reddish
256' - 356' Arcesia	n- 1				l	ł			:	0	color.
No loss	30-			-	l				30.00	, v	
	30-		i	i	1				:		10' - 14' Gravelly Clay, coarse fragmen
Sampling: Sampled cuttings	4-3- Dras	4		l	l				:		to greater than 5 mm indicate gravel interbudded. Contains a few calcareous
approximately at 10				l	1	ŀ			1 7		polites, some opal, mica, and chalcopyr
intervals from dril					l				:		all less than 7% of total. Strong HCL
fluid return ditch.	40-				1			l	40-	: : : : :	reaction. Mottled reddish color.
Water samples taken				1	1				:		
at irregualr inter- vals from casing;	1			1					:		14' - 30' Gravel, poorly sorted; very fine gravel and cobbles; coarse sand wi
river water samples	-	.		15		İ		l	:		minor amount of fine to very fine sand.
also taken.	1		1	ı	1	9			:		Contains some calcareous oolites, mica
	50-				1	1			50-		tlakes, small concretions and chalcopyr
Water Samples:	1	.		l	1	l] :		all less than 5% of total. Staining on
l. Packered hole 51' - 76', blew	1	.			1	l		İ			some fragments, mottled buff color.
hole.	4				İ		, ,		56.0		30' - 196'
Chloride 3,450 mg/l	1		1	ł	ł	l		l	:		Triassic Santa Rosa Sandstone
NaCl 5,692 "	60-				1	l .		1	60.00		
Sultate 700 "	1			1	1	ł		1	:	FH	30' - 56' Sandstone, medium to coarse
fotal Fe 0.12 " Conductance 11,000	- 1	.	į	1		Ì		1			grained, small fraction of very coarse sand and some clay and silt. Good HCL
ounauccunce 11,000	3								1 3	_==	reaction. Calcareous and argillaceous
2. Packered hole]	. 1	1		i			l			cement. Fair induration. Contains som
51' - 76'; after	70-				į	ŀ		1	70-		opal, a few small concretions, and a fe
olow-rest cycle sample cleared.			1					l	72.0		mica ilakes (muscovite). Grades into
Chloride 3.060 mg/l	<u> </u>			1	•			l	1 :		shale. Buff to gray color.
laC1 5,049 "	1		1	ĺ	l			1	76.0		56! - 60! Sandstone, medium to coarse
Sulfate 700 "	1	. 1]					l	1 :	===	grained, with shale layers intertedded.
Total Fe 0.12 "	80-		1	ĺ	j				80-		Weak HCL reaction; argillaceous and cal
Conductance 10,200	1			ĺ	İ				1 :		but to brown color.
. River water	- 4	.		1] -	#==4	Date to order told:
5/25/75.	1]	1				1	1 :		501 - 72! Shale, silty with high argill
hloride 3,150 mg/1		. 1	1				ł	l		1:::::	aceous content. Contains a few coarser
NaCl 5,198 " Sulfate 500 "	90-		1					1	90-	! :::::	sand grains. Weak HCL reaction; fairly well indurated. Variegated dark buff t
Total Fe 0.08 "	1			- 9	i	1		1	- 10		brown color.
Conductance 10,200	-1		ļ .		٥	ł			-		
	1		1			1					
										<u> </u>	
•					ii		ΕX	PLAN	ATIO	<u> </u>	
Псорь											
CORE LOSS											
Type of hole .			b=	Diameni	l. H = Ho	rstollite.	S = She	ot, C = Ch	urn		
Hole sealed .			P =	Packer.	Cm = Ce	nented, C	s = Bo	110m of co 2-3/8".	sing Nx = 3"		
ECOVERY Approx. size o	core (X-serie	e) Ex ies) . Ex s) Ex	* 7/8''.	ĀĒ	1-]/8". 2-]/4". 1-29/32	Bx =	2-3/8", 1-5/8", 2-7/8"	Nx = 2-1 Nx = 3-1	/8" /2"	t
ا مناسمانی مناسمان م											

BEGUN FIN DEPTH AND ELEY. OF WA' LEVEL AND DATE MEASI								1			LOG REVIEWED BY	
LOSSES AND LEVELS, CASING, CEMENTING, CAVING, AND OTHER	TYPE	~ ~	PERCOLATION TESTS					EVA.	DEPTH (FEET)	ں	20	
	SIZE OF	CORE RECOVERY	DEPTH (FEET)		LOSS	RESSURE	LENGTH OF TEST	ELE TIO	DE! (FE	GRAPHIC LOG	CLASSIFICATION AND PHYSICAL CONDITION	
	HOLE	(%)	FROM (P, Ca, or Cm)	то	(G.P.M.)	.M.) (P.S.I.) (MIN.)						
Packered hole 1' - 96' and blew ole; water rose to	4-3/										72' - 76' Sandy Shale, some medium to coarse sand, and silt with high argillaceous content. Fair HCL reaction; fair	
<pre>below top casing, pproximately river</pre>											well indurated. Mottled gray color.	
evel. hloride 2,400	10-			E.				21	10		76' - 146' Sandstone, poorly sorted. \ fine sand to coarse sand, mostly medium	
aCl 3,960 ulfate 200	1										to fime sand. Some shale interbedded. Poor HCL reaction; argillaceous and cal	
otal Fe 0.13											careous cement. Contains some quartzil igneous rock fragments, mica, opal, and	
onductance 6,100	20					ļ			20-		chalcopyrite. Calcareous matter also noted. Fairly well indurated. Gray to	
. Fackered hole l'-116'-interval					1				:		mottled buff color.	
ade no water. ackered hole 91-136	, ,								-		146' - 156' Sandstone, very coarse gra	
ater rose to 4' bel op of casing	bw -								30-	77	Some interbedded shale and conglomerate No noticable HCL reaction, argillaceous	
hloride 2,200 mg/l aCl. 3,663 "	"] :		cement; fairly well indurated. Chalco pyrite deposits on some fragments.	
ulfate 500 " otal Fe 0.10 "							l		-		Mottled buff to gray color.	
onductance 6,750											156' - 196' Sandatone, medium to coars A little very coarse to fine gravel	
. Same as 5, taker	40-								40-		material. Quite a bit of shale inter-	
0 minutes later. hloride 2,100 mg/l							1		146.0		bedded. Little or no MCL reaction. Fairly well indurated; argillaceous ce	
aCl 3,465 " ulfate 500 "											Contains quite a bit of quartzite and other siliceous material, other than	
otal Fe 0.02 " onductance 6,750	50-								50-	<i>0</i>	quartz. Mottled grayish white to buff color.	
. Packered hole]								156.0	3-	196' - 261'	
31'-156', blew hole ole dry; hole took	•					1			130.0		Permian Bernal Formation	
ater 24 gal/min; gain blew hole.and:	60-								60-		196' - 201' Gravelly Shale, contains m coarse tragments to fine gravel size,	
ampled.						}					some fine to very fine sands. Fair HC	
hloride 3,200 mg/l aCl 5,280 "	:							ł			reaction; fairly well indurated. Limo ite particles and some particles limon	
ulfate 700 H otal Fe 0.13 H	70-					l		1	70-		stained. Grayish to whitish mottled c	
onductance 10,000								١.			261' - 356' Permian San Andres Formation	
. Same as 7, pump n water, blew hole						1					(Glorieta Sandstone)	
nd took sample. hloride 3,300 mg/l				1					80-		261' - 316' Quartzose Sandstone. Sucretextured fine to very fine sands. Ver	
aC1 5,445 "	~								"		well sorted. No HCL reaction. Argill	
otal Fe 0.13 " onductance 10,100	-					1			-		aceous cement; poorly indurated. Cont. some chalcopyrite and a few mics flake	
onductance 10,100											Light grayish to white color.	
	90-					*2			90-		316' - 326' Quartzoge Sandstone. Sucre textured fine to very fine, well sorte	
									196.0		sand. No HCL reaction. Seams of shall interbedded. Contains some mica flake	
										-0.50	Grayish-white color.	
			<u> </u>					CPLAN	4 7 1 0		<u> </u>	

GPO 884 - 401

EFATURE

STATE SHEET . 4. OF .5. HOLE NO: . PHOLE

Bureau of Reclamation					GEUL	.UGIC	LUC	UFD	KILL	חטנו		SHEET, OF
FEATURE					Р	ROJECT						STATE
DH-1 LO			• • • • • •			• • • • •	GROU	ND ELEV.				. DIP (ANGLE FROM HORIZ.)
BEGUN FII	ORDS.		• • • • •		E				TO	TAL		BEARING
			• • • • • •	. DEPIR	OF OVE	KBUKUE		• • • • • • •	5	rın	• • • •	···· obaning
DEPTH AND ELEV. OF WALLEVEL AND DATE MEAS	TER URED.					LOG	GED B	Y		• • • • •		. LOG REVIEWED BY
		≿		PERCO	LATION	TESTS		₹_ €	£F.		۲ ۲	
NOTES ON WATER LOSSES AND LEVELS,	TYPE	CORE	DEF	PTH	Ι	2	Et	ELEVA. TION (FEET)	DEPTH (FEET)	GRAPHIC	SAMPLES FOR TESTING	CI ASSISION THOM AND
CASING, CEMENTING, CAVING, AND OTHER	SIZE	ပည္ယ	(FE	ET)		3	LENCTH OF TEST	<u> </u>	۵=	1 00	St	CLASSIFICATION AND PHYSICAL CONDITION
DRILLING CONDITIONS	HOLE	Œ.	FROM (P. Ca.	то	LOSS (G.P.M.)		7.9	}		ۍ. ق	WF	
		(%) =	(P, Cs, or Cm)		(G.P.M.)	(P.S.1.)	(MIN.)				~	<u> </u>
NaC1 27,142 "	4-3/	4	i		İ		l	1			Ш	
Sulfate 2,000 H	Drag Bit -			1	ĺ			l	. 3		Ш	•
Conductance 45,000	P '' 3		1	1				1]		Ш	
	1 3		!			1	1		:		Ш	
16. At 336! circu-		1						l	10-		П	
lated 30 minutes, to sample.	** :	1		1	1]		:		11	
Chloride 18,500 mg/	4		1	1	1]		11	•
NaC1 30,525 "	1	1			1	[316		11	
Sulfate 1,950 "	11				1	i			20-		11	
Total Fe 0.80 " Conductance 52,000	20-		l	İ	l				20-			
	1 3	-		1			1	l	:			
17. At 356' after	1 -		1				1		320 -		11	
flowing 15 minutes. Chloride 16,100 mg/	. 1		l	1				1	:	7.2	11	
NaCl 26,565 "	30-		l		1				30-		11	
Sulfate 1,900 "	1 4		l						:		П	
Total Fe 0.27 " Conductance 51,000	1 1		i		ł	1				12	1	
ponductance 51,000	1 3		1					l	336		11	
18. At 356 after	1 3		1		i	ł		l			11	
flowing 30 minutes.	40-		1						40-		11	•
Chloride 15,950 mg/1	1 :	1	l		1				:	_==		
Sultate 500 "	1 4		l		1				346		11	
Total Pe 0.10 "	1 4		ı]				" :		11	·
Conductance 50,000	50-		l			l	1				11	
19. At 356' after	~3		l			1		1	-	===	11	
flowing all night.	1 3		l			1		l			11	
Chloride 17,400			ľ	1	ļ	1		l	356		11	
NaCL 28,875 Sulfate missing	1 3	1	ł		1	l		l	"	1	Н	
Total Fe 0.48	60				j		ŀ	l	60-	1	H	5 ·
Conductance 49,000	1 3				1	ł	1	l		1	1.1	
20. Sample of	1 3		ļ		1	1	1	l		1		
dringing water.						1	l	1	:	1	11	
Chloride 245 mg/1	1_1		1	1		1	ŀ		:	1	11	
NaC3 404 "	70-		i			l	1	}	70-	1	11	
Sulfate 60 " Total Fe 0.03 "	1 3		ł		1	1		l	1	1	Ш	
Conductance 500-600	1 -	1	1				1	İ	:	1	Ш	
L	1 :		1		1		1		:	1	Ш	
21. Mix of drinking water 50% and waser			ł		1	l		l	80-	1	Ш	
from well 50%.			1	190	1	1	ł			1	П	
Chloride 8,850 mg/l	1 1		l			ł		l		1	П	
NaCl 14,602 " Bulfate 900 "	1 3		1			1	ł	1] -	}	Ш	
Bulfate 900 " Total Pe 0.10 "	1 3					1	1	1		1	H	
Conductance 26,000	90-]	1		1	90-	1	11	
	1 1			1			1	1		1	{ }	
22. River water 5/30/75.) 4		1				[1	:	1	11	•
Chloride 2,900 mg/1					1	1				1	Ш	
NaCl 4,785 "			<u> </u>				L	<u> </u>		1	Ш	
							E)	PLAN	ATIO	N .		
							-			_		•
CORE												A
Was at hete			n	Diemore	l. H = M=	vetellise	S = SI-	ot. C = Ch	urn			
CORE APPROX. SIZE	-	Y		Packer,	Cm = C-	mented, (a = Bo	tiem of coi	ing Nx = 3"			
RECOVERY Approx. size	of noie (V-20116	-: - E - :	- 7/8"	, 22 -	1.1/3"	Br :	2-3/8", 1-5/8".	Nx _ 2-1	/8"		

CEATIOS

FEATURE												ST.	ATE	••••
ноге но: ĎĤ-1 со	CATIOI ORDS.	N N	• • • • • •	· • • • • •	Ε	 	GROU	ND ELEV.	TO1	TAL	• • • •	. DIP (ANGLE FROM	HORIZ.)	• • • • •
BEGUN FIN	IISHED	• • • • •	• • • • • •	DEPTH	OF OYE	RBURDE	EN							• • • • •
DEPTH AND ELEY. OF WA'LEYEL AND DATE MEASU	TER JRED.		• • • • •			LO	GED B	Y	• • • • • •	• • • • •		. LOG REVIEWED B	Y	••••
NOTES ON WATER	TYPE	ERY			LATION	TESTS		ELEVA. TION (FEET)	DEPTH (FEET)	U	SAMPLES FOR TESTING		al al	
LOSSES AND LEVELS, CASING, CEMENTING,	SIZE	~ >	DE I	TH ET)		RESSURI	MCTH	F. F. F.	9.F.	GRAPHIC LOG	LES	CLA Phy	SSIFICATION AND SICAL CONDITION	
I CAVING, AND OTHER	OF HOLE	-	FROM (P, Ca, or Cm)	то	LOSS		7.2			25	AMP			
		(%)	or Cm		(G.P.M.)	(P.S.I.)	(MIN.)				7	·		
Sulfate 500 mg/l Total Fe 0.03 "]					
Conductance 11,000	=										Ш			
23. At 356' after	1				i				1		Ш			
flowing 24 hours. Chloride 16,250 mg/	10-	11						}	10-		Ш			
NaCl 26,812 "]			- 0										
Sulfate 1,750 " Total Fe 0.29 "	3					j		12	1		$\ \cdot \ $	¥		
Conductance 49,000	20-		1						20					
24. River water	1		1						";					
5/3175. Chloride 3,000 mg/l	4		1						3					
NaCl 4,950 "														
Sulfate 500 " 30- 30- 30- 30- 30- 30- 30- 30- 30- 30-														
Conductance 11,900	3			9		1					11			
Note: Comductance	=		1						-		Ш			
in micromhom /cm a 25° C	3													
	40-								40-		H			
]										Ш			
	3								3					
	50				İ				50-		11			
	-		1	İ					":					
] 3	Н												
		Ш			ł]		П			
8.	60-								60					
]		П			
	-	1							-		Ш			
	3						1				П			
99	70-								70-		П			
	:	Н.		ĺ							П			
				İ]		П			
	80-	Ш		ŀ			1		80-		Ш		20	
				İ							Ш		¥0	
	-								-	}	11	•		
]		$\ $			
	90-		1						90-		Ш			
			1				1		3		$\ \cdot\ $			
	:		4)]		Ш			
	<u> </u>				50] :					
							E)	PLAN	ATION	1				
CORE														
LOSS				Ni	1 H - U-	wetalite.	5 _ EL	w C = CL	urn				•	
CORE Hole seeled Approx. size	f hole	(X-serie	P =	Packer, = 1-1/2*	Cm = Ce	mented, 1-7/8",	C. Bo	ttom of car 2-3/8",	ing Nx = 3"				· ·	
CORE RECOVERY Approx. size of Outside dis. of Inside dis. of	í casin	(X-serie g (X-ser	ies). Ex	= 7/8", = 1-13/1	6", Åx =	1-1/8", 2-1/4",	Bx 4	1-5/8", 2-7/8",	Nx = 2-1/ Nx = 3-1/ Nx = 3''	2"				
Inside dia. of	casing	(A-Serie	#/ · EX	- 1-1/4	, 4= =	1-47/34	, ux	,						

..... PROJECT..... STATE...... SHEET 5. OF 5. HOLE NO .DH+1.....

CORE

CORE

CORE

FEATURESTATESTATESHEET .. 3 . OF .6. HOLE NO. DH-2

7-1337 (9-69) Bureau of Reclamation	GPO 884	-401	GEO	_OGIC	LOC	OF D	RILL	HOLE	<u> </u>	SHEET3 OF
FEATURE			F							STATE
HOLE NO	OCATION				GROU	ND ELEV		<u></u>		DIP (ANGLE FROM HORIZ.)
BEGUN FI							DE	TAL PTH		BEARING
DEPTH AND ELEV. OF WA	ATER			LOG	GED B	Y				LOG REVIEWED BY:
LEVEL AND DATE MEAS	· · · · · · · · · · · · · · · · · · ·	T	DLATION						~	T
HOTES ON WATER	TYPE WW	DEPTH	JEXTION	T ==	I-	ELEVA. TION (FEET)	DEPTH (FEET)	ဋ	SAMPLES FOR	
LOSSES AND LEVELS, CASING, CEMENTING, CAVING, AND OTHER	AND SIZE OUD OF HOLE	(FEET)	1	RESSUR	LENGTH OF TEST	릭트	26	GRAPHIC LQG	STI	CLASSIFICATION AND PHYSICAL CONDITION
DRILLING CONDITIONS		FROM (P. Cs. TO or Cm)	Loss	1 6	(WIW.)			ຮ	AMF	
	(%)	or Cm)	(G.P.M.)	(F.3.1.)	(MIN.)		<u> </u>			
1	4-3/4 Drag								4	Some very fine sand to silt size parti-
	Bit	1					-			les. Argillaceous content high. Good
	11 11								l l	ICL reaction; fairly good induration.
	10-			1			10-		113	Contains chalcopyrite crystals. Brownis
(4			*	ĺ			11.0	5-0	П.	
		1 1							51	lo' - 336' Ferruginous Siltstone, fairly sandy, fine to very fine sand. Argill-
	1131						216.0	<i>D</i>	}	acous content lessening. Fair HCL
			Ì				:		1 2	reaction. Contains some chalcopyrite
							20-		ا ا	and mica flakes. Reddish-brown color.
]			1			:		1	EXPLANATION OF GRAPHIC LOG:
	[]			1			:		П	
	$\begin{bmatrix} 1 \end{bmatrix}$		1				:			
•]	30-			
			1] :		Ш	
	1131	1 1					×		Ш	INDURATED ROCK
	11 11		1				:		Ш	INDURATED ROCK
	40-		1		1	1	40-		П	Sandstone
	1131				İ	1			П	Shale
	1141			1			:			
	1141		1	1		1			П	Sandy shale
	50-	1	1	1			50-		Ш	Sandstone and shale
•		1	1		1	l			Ш	<u></u>
	11 11								11	CONSTITUENT PARTICLES
	1131	1	1			1			П	
						1				33 Clay
	60-	i i	1		1	l	60-		Ш	Pebbles, gravel, cobbles, or
	11 11	1 1	45	1	1				Ш	boulders
	1141	1	1		1	1	-		Ш	· Sand
	11 31			1		l			П	Silt
	10-		ł	1		1	70-		Ш	
		1	ł		ļ	et.			Ш	MISCELLANEOUS SYMBOLS
	1131				1	1	:	FEET		[F] Proise
	11 11	·								D Pyrite
	80-				1		80-		П	v Mica
	1141				1		1	F	Ш	Coal
	-			1	1		-			_
	11 11								П	
	90-		1		1	1	90-	ļ		
	$\{[]\}$				1			}===		
	1141		1		1					
					1		96.0		11	
	<u> </u>				<u> </u>	<u> </u>		1	Ш	
					<u>E)</u>	CPLAN	ATIO	M		
CORE										
LOSS										÷
Type of hole Hole seeled		D = Diamon	d, H = He	ystellite,	S = Sh Ca = Bo	ot, C - Ch	urn sing			
RECOVERY Approx. size	of hole (X-serie		Ax s	1-7/8", 1-1/8",	Bx :	2-3/8", 1-5/8",	Nx = 3"	/8"		
Outside dia. Inside dia. o	of casing (X-se f casing (X-seri	ries). Ex = 1-13/2 es) Ex = 1-1/2	Ax Ax	1-29/32	Bx Bx	23/8",	Mx = 3.1			
										manager of many of manager than of

CORE CORE

FEATURE					Р	ROJECT						STATE
HOLE NO. VO. Z			• • • • • •		-	• • • • • • • • • • • • • • • • • • •	GROU	ND ELEV	то	TAL :	• • • •	. DIP (ANGLE FROM HORIZ.)
1												BEARING
LEVEL AND DATE MEA	SURED.				•••••	LOG	GED B				-	. LOG REVIEWED BY
NOTES ON WATER	TYPE	CORE	<u> </u>		LATION '		1 +	ELEVA. TION (FEET)	DEPTH (FEET)	<u>ن</u> ا	SAMPLES FOR TESTING	
LOSSES AND LEVELS, CASING, CEMENTING, CAVING, AND OTHER	SIZE OF		DEF (FE	ET)	LOSS	RESSURE	LENGTH OF TEST	3.5	- BE	GRAPHIC LOG	ESTU	CLASSIFICATION AND PHYSICAL CONDITION
DRILLING CONDITIONS	HOLE	(%)	FROM (P, Ca, or Cm)	TO	(G.P.M.)		(MIM.)			5	NY S	
	43/9		J. C,				 				1	
	Drag Bit-	11										·
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							<u>E X</u>	PLAN	ATIO	Ħ		
CORE												
Type of hole	• • • • •		<u>p</u> =	Diamond	, H = Ho	ratellite,	S = She	ot, C = Ch	urn			
CORE RECOVERY Approx. size Outside dis.	of hale	(X-serie (X-serie	s) Ex =	1-1/2": 7/8",	; Ax =	1-7/8", 1-1/8",	8x =	2-3/8", 1-5/8",	Nx = 3" Nx = 2-1/	/8" /2"		· ·
Outside die.	of casin	g (X-ser	105). EX 4	1-13/10	, Ax =	1-29/32	". Bx =	2-3/8	Nx = 3''	4		• •

STATE SHEET ... OF ... HOLE NO. .. PH-2.....

SHEET... 1... OF... 2....

FEATURE . Canadian . Ri	ver.				P	ROJECT	Lake	Meredi	h Sali	inity	Study STATE. New Mexico
LOW DOWN	CATIO	и Ве:	lov.Ute	.Dam.			GROU	ND ELEY.	3,674.	.73'	DIP (ANGLE FROM HORIZ.) 90.0°
BEGUN . 9-23-77 FIN	ORDS.	N	_12_77		E		:	26.51	TO	TAL 3	18-0' BEARWE
DEPTH AND ELEY. OF WAT LEVEL AND DATE MEASU	TER IRED.	Ar	tesian.			LOG	GED B	yŞhir	ey Sh	edix.	Log reviewed by J., L., Jackson
		ERY		PERCO	LATION '	TESTS		₹-F	₹£		20 8
NOTES ON WATER LOSSES AND LEVELS.	TYPE	# > □	DEI	тн		2	ΞË	ELEVA. TION (FEET)	DEPTH (FEET)	GRAPHIC LOG	SE CLASSIFICATION AND
CASING, CEMENTING,	SIZE	COR	(FE	ET)	LOSS	RESSUR	HCTH TEST	<u> </u>	85	ŽŎ	PHYSICAL CONDITION
CAVING, AND OTHER DRILLING CONDITIONS	HOLE		FROM (P, Ca, or Cm)	TO			22			5	PHYSICAL CONDITION
		(%)	or Cm)		(G.P.M.)	(P.\$.I.)	(MIN.)				7
_	1 1				1						0.0' - 26.5': QUATERNARY ALLUVIUM.
Stapp-Hamilton Inc.	1			1	1	l					OLO - 2015 . QUALEMARI ALLOVIOM.
Austin, Texas	1 7		1	ŀ			1		11.0	 	0.0' - 11.0': Sand.
Solicitation No.	11			[1				15.0	-	Approximately 80% medium to coarse
7-07-50- S 0970	38	1			1				10		sand, approximately 20% fine gravel,
n 1050 n.::11:	1 1	1		1	1	ļ	1		26.5		hard, subrounded to subangular rock and mineral fragments, buff. SP
Damco 1250 Drilling rig.	1 1			1					28.7		and manoral araginerous, said
rig.	1 7			1					32.0		11.0' - 15.0': Clayey Gravel.
Drill Fluid	1, 1		l		1]		1		}	Approximately 70% fine, hard, sub-
Additives and Drill	42		1	1	1	1			20	1	rounded to subangular rock and mineral
Water Return.	1		1		I]		<u> </u>	-	1	fragments, maximum size 1.2", approxi- mately 30% medium plasticity fines of
(Ft.) (%)	1		1		1				-	}	medium dry strength, medium toughness,
(* ** / (///] 7								-	1	no dilatancy, weak to moderate reaction
0.0- 6.0 0.0	[]		1			1			58.0		with HCl, reddish-gray. GC
(600.0 lbs. revert)			ŀ	ĺ	-				30-	7	15.0' - 26.5': Sand.
6.0-4 6-6 90.0 40.0-120.0 100.0				l		1			63.5	-	Approximately 75% medium to coarse,
120.0-140.0 70.0	1 1		1		1				68.0		some clay, approximately 25% fine,
140.0-142.0 35.0	1 7		l		ł				73.0		hard, subrounded to subangular rock
142.0 0.0	[,,]			i					78:3		and mineral fragments, reddish-gray. S
(600.0 lbs. revert)					l				70 86		26.5" - 318,0': TRIASSIC SANTA ROSA
142.0-210.0 100.0 210.0-220.0 80.0	1 4			l	1	1	1	l	86.0		SANDSTONE. (TRUJILLO AND TECOVAS
220.0-230.0 30.0			1	l		1	l	1	90.0		FORMATIONS OF TEXAS)
230.0-240.0 60.0						1	İ	l	:	1	
(400.0 lbs. revert)	I KXV		1	i	1	1		l	:	Ì	26.5' - 90.0': Sandstone.
240.0-258.0 50.0	. 50-	11			1			1	50-	1	Medium to coarse-grained, silty, micaceous, moderately indurated, cal-
(700.0 lbs. salt mu 258.0 0.0	1 1		1				1		:	1	careous cement, layers and stingers of
(800.0 lbs. salt mu	1.	-5		l		1	1	1		1 .	interbedded shale, buff to grayish-tam
(200.0 lbs. revert			i				.00	1	:	i	
and 50.0 lbs. salt)	120-	!			1				:	1	28.7' - 32.0', 58.0' - 63.5', 68.0'
258.0-266.0 0.0	60	Ш				1	1	1	60-	1	73.0', and 86.0 - 90.0': Shale Argillaceous, sandy, small amount of
(1,500.0 lbs. bentonite)] :	Н	1	1	1	1		1		1	gravel, sticky when wet, calcareous,
266.0-318.0 0.0	1	11	1	1		1		1]	1	red-brown and gray layers.
_] :	11	1			1		1] :	1	70 / 1 70 71 . 6 6 6 1
Artesian flow below 294.0' between	140		1		7	1	!	ł	:	1	78.4' - 78.7': Soft Coal.
periods of drilling		11	İ	1			ŀ	1	70	1	90.0' - 251.0': Sandstone. '
operations.		11	1	1		ŀ		1		1	Medium to very coarse-grained, silty,
	-	11	1	1			1	1	-	}	poorly sorted, calcareous, moderately
Sampled cuttings at		11	1	1	1			1		Ĭ	indurated, conglomeratic from 203.5' 208.0' and from 220.0' - 230.0',
approximately 10.0' intervals from dril	160]	1		1	1		1	80-	1	200.0' and from 220.0' - 230.0',
fluid return ditch	1 00	1	1	1					911-	ł	
from 0.0' - 261.0'.]	1	1			i	1		}	251.0' - 271.0': Shale.
Logged using	-]	1		1	ł		I	-	- 1	Argillaceous, sticky when wet, some
binocular micro-	-	11	1		1			1		1	calcareous cement, with interbedded sandstone, blue-gray.
scope.	180-		1					1	90	1	January January 91 47
Geophysical logging		11	1					1		1	256.0' - 258.0' and 263.0' - 264.0':
on 10-7-77.		11	i				1			1	Sandstone.
	-	11					1	1	-	1	Medium to coarse-grained, silty,
Core samples	1	<u> </u>	1				1	1		1	some calcareous cement, well indurated, very hard, blue-gray.
obtained from 261.0' - 318.0'.	200	Ш_						<u> </u>	200.0	1	
								CPLAN			
Used 6" ro	ck b	it to	261.0	; NX	cure a	nd star	ndard	split-	tube pe	enetr	ation resistance to 318.0'; set 25' of 6"
CORE surface ca	sing	, wit	h top 1	1.5' b	elow gi	round a	surfac	ce; pull	led sur	rface	casing 10-13-77, installed 79.0' PVC
CORE surface ca	sing	, wit	h top 1	1.5' b	elow gi	round a	surfac	ce; pull	led sur	rface	casing 10-13-77, installed 79.0' PVC to 1.0' above surface, gravel packed to

1.25" well screen attached at bottom of 234.0' of 2" steel casing to 1.0' above surface, gravel

top of screen, sand packed 1.0' over gravel

						GLOL	.00.0			////	1102		
HOLE NO.	POW-1 LC	CATIO	Pp	joa hte	PAP.		• • • • •	GROU	ND ELEV	3.67	••73 !		udy State Naw Maxico
BEGUN	9-23-77 FI	NISHED	. 10	-13-77.	DEPTH	1 OF OVE	RBURDE	N . 26	•5!	DE	PŤĦ	.318	8.Q'. BEARING
DEPTH AN	D ELEY. OF WA	TER URED.	Ar	tesian.			LOG	GED B	y.Shirl	ey.She	dix	• • •	LOG REVIEWED BY J. L. Jackson
			`		PERCO	LATION			₹.F	EE.		ĕ.	
LOSSES A CASING, CAVING.	ON WATER HD LEVELS, CEMENTING, AND OTHER	TYPE AND SIZE OF	CORE	DEP (FEI	TH ET)	LOSS (G.P.M.)	ESSURE	LENGTH OF TEST	ELEVA. TION (FEET)	DEPTH (FEET)	GRAPHIC LOG	SAMPLES FOR	CLASSIFICATION AND PHYSICAL CONDITION
DRILLING	CONDITIONS	HOLE	æ (%)	FROM (P, Cs,	TO	(G.P.M.)	(P.S.I.)	(.NIW.)		'	٥	3	7
DRILLING	CONDITIONS	220 220 240 240 280 30 30 30 30 30 30 30 30 30 30 30 30 30		FR.CM (P, Cm)	•	(G.P.M.)	(P.S.I.)	(MIN.)		251.0 256.0 258.0 263.0 264.0 271.0		3	271.0' - 318.0': Sandstone. Fine-grained, well sorted, very lightly indurated to well indurated, very slightly cemented to highly cemented, mica, thin intermittent shale seams with pyrite erystals and limonite staining, blue-gray.
		•						E >	PLAN		4		
CORE	Type of hole Hole sealed		:: •••••	P=	Diamoni Packer,	d, H = Ho Cm = Co	ystellite, mented, (<u>-</u>	-		÷
RECOVERY	Type of hole Hole sealed Approx. size Approx. size Outside dia. of	of hole (of core (of casing casing	X-serie X-serie (X-serie X-serie	e) Ex = s) Ex = ies) . Ex = s) Ex =	1-1/2" 7/8", 1-13/1 1-1/2"	6", Ax	1-7/8", 1-1/8", 2-1/4", 1-29/32	8x 4 8x 4 8x 4	2-3/8", 1-5/8", 2-7/8", 2-3/8",	Mx = 3" Mx = 2-1, Mx = 3-1, Mx = 3"	/8". /2"		

HOLE NO. OW-2	OCATIO	н	elow U	te Dee	·		GROU	ND ELEV	.3,67	6, 88 :		DIP (ANGLE FROM HORIZ.)90.0°
BEGUN 10-27-77	INISHED	1-4-	78	. DEPTH	OF OVE	RBURDE	N?	o•o:	DE	TAL PTH.3	48.0) BEARING
DEPTH AND ELEV. OF W LEVEL AND DATE MEA	ATER SURED.	. Arte	pien			LOC	GED B	YShir		edix.		LOG REVIEWED BY. J. J. Jackson
NOTES ON WATER	TYPE	ERY FR			LATION	TESTS		EVA.	DEPTH (FEET)	U	S.	
LOSSES AND LEVELS, CASING, CEMENTING, CAVING, AND OTHER	AND SIZE OF	ECOR	(FE	ET)	LOSS	ESSUR	LENGTH OF TEST	ELE (FE	96 F	GRAPHIC	TESTIN	CLASSIFICATION AND
DRILLING CONDITIONS	HOLE	(%)	(P, Ca, ar Cm)	TO	(G.P.M.)	(P.S.I.)	10 (MIN.)			<u>.</u>		*
Stapp-Hamilton Inc Austin, Texas									20.0			.0' - 20.0': QUATERNARY ALLUVIUM.
Solicitation No. 7-07-50-Se970	48								10			0.0' - \$0.0': Silty Sand. Approximately 80% fine to coarse, angular to subrounded sand, maximum size
Damco 1250 Drillin	0										1	0.2", approximately 20% low to medium plasticity fines, low toughness, low iry strength, quick dilatency, strong
ed to total depth;	_		,					•]		1 1	to moderate reaction with HCl, buff. SM
Failing Drilling rig reamed to total	1 80								3		20	0.0' - 348.0': SANTA ROSA SANDSTONE.
depth.									91.0		(1	TRUJILLO AND TECOVAS FORMATIONS OF EXAS)
Sampled cuttings a approximately 10' intervals from	120								-			20.0' - 91.0': Sandstone. Medium to coarse-grained, silty, poorly
drill fluid return ditch from 0.0' -	30-							ļ .	30			sorted, calcareous cement, with layers of shale at 28.0' - 30.0' and 78.0' -
300.0'. Logged	1 :	Ш	1					į	:	1		91.0', and small amount of coal within
using binocular	1 -		ļ	1			ļ		-	1	118	30.0' - 90.0' interval, tan.
Geophysical loggin	9 160								40			91.0' - 230.0': Sandstone. Medium to coarse-grained, silty, poorly sorted, calcareous cement, with gray
on 12-16-77.											4	shale layer 145.0' - 153.0' and very thin gray shale layers interbedded in
obtained from 300.0' - 348.0'.	3										2	200.0' - 220.0' interval, blue-gray.
Hole completion included gravel	200								50		5	230.0' - 261.0': Shale. Sandy, blocky, sticky when wet, cuttings are Lean to Fat Clay, medium
pack around 6.0" casing from bottom									232.0		;	to high plasticity, medium toughness, with thin interbedded gray sandstone,
of hole to unknown depth (242.0') according to as-	240 60								60		Ш	gray. 261.0' - 333.0': Sandstone.
built diagram in file). Added 88									261.0		: <u>.</u>	Fine-grained, well-sorted, slightly cemented to highly cemented, with gray
cubic feet grout to G.L. in three	280 70										П	shale layer 288.0'- 291.0', light gray.
stages, last two sacks 3-1-78. Special watertight									70			333JO' - 348.0': Shale. Sticky when wet, gray.
cap placed on 6"	-		4	}					-	1	Ш	
steel casing.	320	55 18	-4			!	1				Ш	<u>g</u> 1
	320 90	51	-1					İ	222.0			
		95 0					j		333.0			9
	360 90								90			,
#												
â	400								400.0			
Used 4-1/		bit	0.0' -	300.0	o'; NX	core b		PLAN with d	ATIO	N	from	n 300.0' - 348.0'. Set 2.0' of 12"

Fors CORE RECOVERY surface casing 10-31-77. Set surface casing to 24.0' with 1.0' above G.L. on 11-2-77. Grouted 12" casi in hole on 2-9-78. Placed below.G.L. 260.0' of 6" casing with 6.0' above G.L. and with 80.0' of 2" PVC

screen attached to bottom. Driller did not measure casing or hole before placing 6"

DEPTH AND ELEY. OF WAT LEVEL AND DATE MEASU	ER RED	År	tesian	·		. LOG	GED B	Shirl		dix.	LOG REVIEWED BYJ. L. Jackson
MOTES ON WATER		RY		PERCO	LATION .	TESTS		EVA.	DEPTH (FEET)	u	CLASSIFICATION AND
LOSSES AND LEVELS, CASING, CEMENTING, CAVING, AND OTHER	TYPE AND SIZE OF HOLE	CORE	FROM (P. Cs. or Cm)	TH ET)	LOSS (G.P.M.)	PRESSURE	LENGTH OF TEST	ELE TIO (FE)	DE!	GRAPHI LOG	CLASSIFICATION AND PHYSICAL CONDITION
		(%)	or Cm)		(G.P.M.)	(P.3.1.1	(#14.7				<u> </u>
Stapp-Hamilton Inc. Austin, Texas	* 1								,		0.0'- 48.1': QUATERNARY ALLUVIUM.
Solicitation No. 7-07-50-S0970	40								10-		Approximately 80% fine to coarse, angular to subrounded sand, approxi- mately 20% none to low plasticity
ailing 1500 Drill- ing rig.	1								48.1 60.0		fines, low toughness, quick dilatancy, trace gravel, strong reaction with HCl buff to tan. SM
Jsed 7-7/8" tricone rock bit to 362.0'. Set 49.6' of 6"	80 :								72.0		48.1' - 362.0': TRIASSIC SANTA ROSA SANDSTONE. (TRUJILLO AND TECOVAS FORMA-
surface casing 1-12-78. Set 270.0° of 2" steel casing	1								1		TION OF TEXAS) 48.1' - 60.0': Sandstone. Fine to coarse-grained, subangular to
with 80.0' of 2" screen below to 350.0'. Gravel pack	20								30		rine to coarse-grained, subangular to subrounded grains, strong reactions with HCl, tan.
(0.7 cubic yards) from bottom of hole to 260.0' depth, cement grout	1										60.0' - 72.0': Gravelly Shale. Sticky when wet, medium to coarse sand and gravel up to 5/8" maximum size,
emplaced to within 3/4" of G.L. Watertight steel	160-								162. 19 170.0		gray. 72.0' - 162.0': Sandstone.
cap placed on steel casing.	200			2		(4)			191.0	_	Medium to soarse-grained, silty, poorly sorted, mica, calcareous cement contains apatite, thin gray shale
Geophysical logging on 1-20-78. Sampled cuttings at	50								212.0		layers interbedded, blue-gray. 162.0' - 170.0': Shale. Sticky when wet, argillaceous, cutting
approximately 10' intervals from drill fluid return	240								40		are Lean to Fat Clay with medium to high plasticity and high toughness, blue-gray.
ditch from 0.0' - 362.0'. Logged using binocular	64										170.0' - 191.0': Sandstone. Fine-grained, well-sorted, light gray.
microscope. Drilled with clear water.	280								70		191.0' - 212.0': Shale. Sticky when wet, argillaceous, cutting are Lean to Fat Clay, medium to high
water.	:								<u>3</u> 02.0		plasticity, medium toughness, blue- gray.
	320 86								326.0	_	212.0' - 302.0': Sandstone. Fine-grained, well-sorted, rounded to subrounded grains, mica, some 1.0 to 3.0' interbedded shale layers 258.0',
·	260								350.0		269.0', and 286.0', light gray.
:	360								62.0.		Sandy, argillaceous, sticky when wet, gray. 326.0' - 350.0': Sandstone.
	±00	3.			,					1	Fine-grained, silty, hed-brown.

CORE

DEPTH AND ELEV. OF WATER LEVEL AND DATE MEASURED											
LOSSES AND LEVELS, CASING, CEMENTING, CAVING, AND OTHER	TYPE AND SIZE OF HOLE	RECOVER	FROM (P, Cs, or Cm)	TH	LOSS (G.P.M.)	RESSURE	E LENGTH	ELEVA. TION (FEET)	DEPTH (FEET)	GRAPHIC LOG	CLASSIFICATION AND PHYSICAL CONDITION
Stapp-Hamilton Inc.	1 1 1								11.0		0.0' - 11.0': QUATERNARY ALLUVIUM
	٠٠ ١٥ .								101		0.0' - 11.0': Sand. Predominantly fine to medium, maximusize 1/8", round to subangular, hard rapid reaction with HCl, trace of fibuff color. SP
ailing 1500 Drill- ing rig. Used 7-7/8" tricone	1			,					62.0		11.0' - 382.0': TRIASSIC SANTA ROSA STONE. (TRUJILLO AND TECOVAS FORMAT
ock bit 0.0-15.0'. Set 15.0' of 6-5/8" surface casing,	3020								75.0]		OF TEXAS) 11.0' - 62.0': Sandstone.
with 1.0' above G.L. Used 4-1/2" tricene rock bit 15.0' to T.D. Set	120								100.0		Fine to medium-grained, subangular of subrounded grains, moderately indur- slightly to highly cemented, calcard cement, hard, tan.
293.0' of 2" steel casing, with 0.95' above G.L. and 84.0 of 2" slotted steel											62.0' - 75.0': Sandstone. Fine to medium-grained, subangular subrounded grains, silty, clayey, he
casing attached to bottom. Gravel packed from bottom of hole to 287.0', sand 287.0'-285.0', and neat cement to	160								40		75.0' - 100.0': Shale. Very sticky when wet, well cuttings could be described as Lean to Fat C with high toughness, calcareous cem
G.L. Left surface casing in hole. Watertight steel cap placed on steel	200 50								204. 18 .		with fine to medium gray sandstone layers interbedded, mostly gray, but some red-brown.
casing. Geophysical logging on 12-13-78. Sampled cuttings at	240 60								60		100.0' - 204.0': Sandstone. Fine to coarse-grained, angular to subangular grains, argillaceous, round mineral fragments, well indurate approximately 1.0" seam of soft coal in upper 10.0' and thin bed of gray
approximately 10' intervals from drill fluid return	280										shale within interval 162.0' - 172.0 gray.
ditch from 0.0'- 382.0'. Logged using binocular microscope.	280								290.0		204.0' - 290.0': Shale. Thin lenses of gray shale, predoming nantly red-brown, sticky when wet. Well cuttings are Lean to Fat Clay, high toughness and medium to high
Drilled with clear water.	32Q								80		plasticity. Thinly interbedded grassandstone. Fine to coarse-grained, some calcareous cement.
	360								355:0		290.0' - 355.0': Sandstone. Fine-grained, rounded to subrounded grains, well-sorted, mica, tan to g
									382.0		355.0' - 382.0': Shale. Sandy, sticky when wet, blue-gray.
	400	Ц	<u></u>	<u> </u>	<u></u>	ļ	E :	CPLAN	A T 1 0 1		



Index to TAB 11: Water-level data

Water level elevations in wells OW-3 and DH-3, September 1983 through September 1984 (data and graph)

Water level elevations in Ute Reservoir and TW-1, September 1984 through April 1991 (data and graph)

NOTE: Water-level data for USBR piezometers are listed in **TAB 10** - USBR, 1984, Figure 2; water-level data for all significant wells in Logan area are summarized in **TAB 9** - "Deep wells in Logan area with both water level and water quality data" and "Shallow wells in Logan area with both water level and water quality data."

Water	laval	elevations.	-64
water	level	elevations.	π.

Date	OW-3	DH-3
09/14/83		3695.86
10/17/83		3694.96
11/17/83	3680.91	3695.26
12/13/83	3680.98	3695.16
01/12/84	3680.95	3695.08
02/15/84	3681.10	3695.16
03/16/84	3681.01	3695.01
04/13/84	3681.03	3695.00
05/10/84	3681.26	3694.96
06/08/84	3680.69	3694.96
07/10/84	3680.88	3694.80
08/17/84	3681.04	3689.67
09/06/84	3681.03	3689.10

NOTES: Surface elevation of OW-3 is 3678.3 ft; surface elevation of DH-3 is 3781 ft.

Air lifted DH-3 on 7/19/84 1 hour. Black fluid (floating, lighter materials?) like drilling mud blown out of hole, then clear water. Hole may not have been properly cleaned after drilling. Piezometer should be pumped and tested before using for testing.

05/10/84 07/10/84 09/06/84 Water levels in two observation wells, September, 1983 - September, 1984 DH-3 03/16/84 + Date 01/12/84 OW-3 09/14/83 11/17/83 3,696 3,689 3,688 3,686 3,685 3,684 3,683 3,682 3,695 3,694 3,693 3,690 3,687 3,680 3,692 3,681 3,691

Water-level elevation, feet

Ute Reservoir Elevation, observation well and river discharge Reservoir and observation well levels from worksheet STATEQUA.WK1 River discharge, 7—day geometric mean, Logan Station daily values file

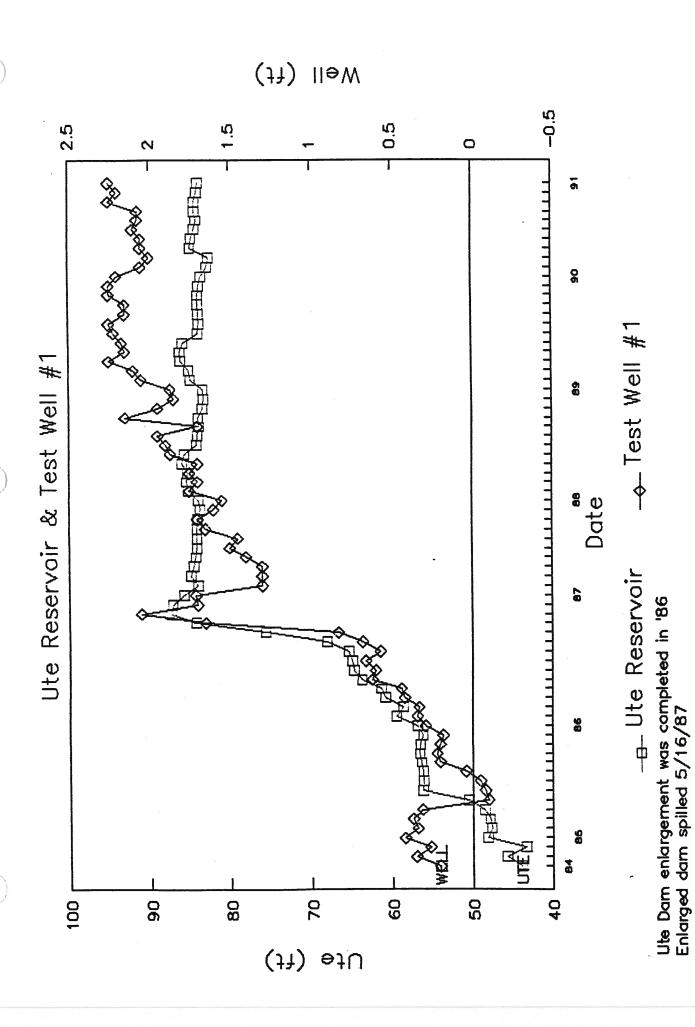
	UTE ¹	WELL ²	RIVER	YEAR	WELL	RIVER
DATE	(ft)	(ft)	(cfs)		(ft X 25)	(cfs X 10)
09/08/84	43.85	0.20	1.61	84	5	16.07
03/21/85	45.69	0.35	1.78	85	8.75	17.83
04/23/85	43.38	0.26	2.00		6.5	19.98
06/21/85	48.12	0.42	3.64		10.5	36.41
07/18/85	47.65	0.34	4.73		8.5	47.25
08/15/85	47.86	0.37	3.55		9.25	35.54
09/13/85	48.50	0.31	2.65		7.75	26.48
10/15/85	50.53	-0.10	2.12		-2.5	21.20
11/13/85	56.15	-0.08	2.10		-2	21.00
12/17/85	56.04	-0.05	1.79		-1.25	17.88
01/17/86	56.14	0.04	1.88	86	1	18.84
02/13/86	56.34	0.20	1.97		5	19.73
03/13/86	56.54	0.22	2.25		5.5	22.52
04/10/86	56.50	0.20	2.50		5	25.05
05/09/86	56.11	0.18	2.69		4.5	26.90
06/06/86	56.80	0.29	2.06		7.25	20.63
07/10/86	59.41	0.34	2.00		8.5	20.00
08/08/86	58.59	0.33	1.96		8.25	19.58
09/05/86	60.80	0.42	1.92		10.5	19.20
10/07/86	61.28	0.44	2.28		11	22.84
11/05/86	63.67	0.62	2.44		15.5	24.36
12/05/86	64.67	0.60	2.33		15	23.28
01/06/87	64.91	0.66	2.61	87	16.5	26.12
02/05/87	65.36	0.57	1.21		14.25	12.13
03/05/87	68.04	0.68	0.84	ř	- 17	8.38
04/03/87	75.59	0.83	0.49		20.75	4.89
05/05/87	84.22	1.65	359.11		41.25	3591.10
05/18/87	87.34	2.05	138.82		51.25	1388.21
06/03/87	87.15	1.70	449.57		42.5	4495.70
07/01/87	85.71	1.71	335.56		42.75	3355.63
08/06/87	83.96	1.30	3.44		32.5	34.40
09/02/87	84.83	1.30	3.31		32.5	33.14
11/04/87	84.47	1.30	4.15		32.5	41.47
12/03/87	84.19	1.40	3.74		35	37.45
01/08/88	84.06	1.50	6.28	88	37.5	62.81
02/05/88	84.05	1.45	3.98		36.25	39.77
03/04/88	84.07	1.65	4.85		41.25	48.55
04/04/88	83.98	1.70	5.03	**	42.5	50.28
05/09/88	83.68	1.60	3.68		40	36.75
06/03/88	83.83	1.55	2.80		38.75	27.96
07/01/88	85.11	1.75	3.92		43.75	39.17
08/05/88	85.36 85.00	1.70	4.12		42.5 43.75	41.16
09/02/88	85.22 85.91	1.75	4.63 5.30		43.75 42.5	46.28 53.05
10/07/88	85.62	1.70 1.87	12.59		42.5 46.75	53.05 125.92
11/04/88	84.08		2.25		40.75 47.5	22.48
12/02/88	84.00	1.90	2.25 3.51	89	47.5 48.75	35.07
01/06/89 02/03/89	83.80	1.95 1.70	4.31	09	42.5	43.13
03/03/89	83.92	2.15	4.46		53.75	44.61
04/07/89	83.33	1.95	4.11		48.75	41.08
05/05/89	83.22	1.85	4.63		46.25	46.28
06/02/89	83.29	1.87	4.11		46.75	41.08
06/29/89	84.87	2.05	5.56		51.25	55.57
07/28/89	85.08	2.10	4.82		52.5	48.17
08/31/89	86.07	2.25	4.26		56.25	42.65
09/29/89	86.15	2.15	2.55		53.75	25.52
10/27/89	85.79	2.17	5.20		54.25	52.03
12/01/89	83.86	2.22	3.70		55.5	37.00
12/21/89	83.71	2.25	3.82		56.25	38.24
			-			

Ute Reservoir Elevation, observation well and river discharge Reservoir and observation well levels from worksheet STATEQUA.WK1 River discharge, 7—day geometric mean, Logan Station daily values file

	UTE ¹	WELL ²	RIVER	YEAR	WELL.	RIVER
DATE	(ft)	(ft)	cfs)		(ft X 25)	(cfs X 10)
01/26/90	83.80	2.15	4.61	90	53.75	46.13
02/23/90	83.88	2.15	5.06		53.75	50.60
03/30/90	83.84	2.25	5.04		56.25	50.39
04/27/90	83.75	2.25	4.22		56.25	42.22
05/28/90	83.43	2.20	3.73		55	37.28
06/29/90	82.73	2.05	3.50		51.25	35.00
07/27/90	82.48	2.00	3.56		50	35.63
08/29/90	84.80	2.05	3.91		51.25	39.14
09/28/90	84.65	2.05	3.70		51.25	36.97
11/30/90	84.24	2.10				
12/21/90	84.03	2.07				
01/25/91	84.24	2.07		91		
02/28/91	84.25	2.25				
03/29/91	83.96	2.20				
04/26/91	83.80	2.25				
05/22/91	83.77					

¹ Ute datum is 3700 above mean sea level (MSL); thus, on September 8, 1984, Ute reservior elevation was 3743.85 above MSL.

² Well TW-1 datum is ground level, 3674.01 above MSL. Positive numbers are water levels above ground level; negative numbers are water levels below ground level.





Index to TAB 12: Aquifer test data

Dutton, 1987, Figure 1

Dutton, 1987, Figure 9

Dutton, 1987, Figure 10

Dutton, 1987, Figure 11

Dutton, 1987, Table 2

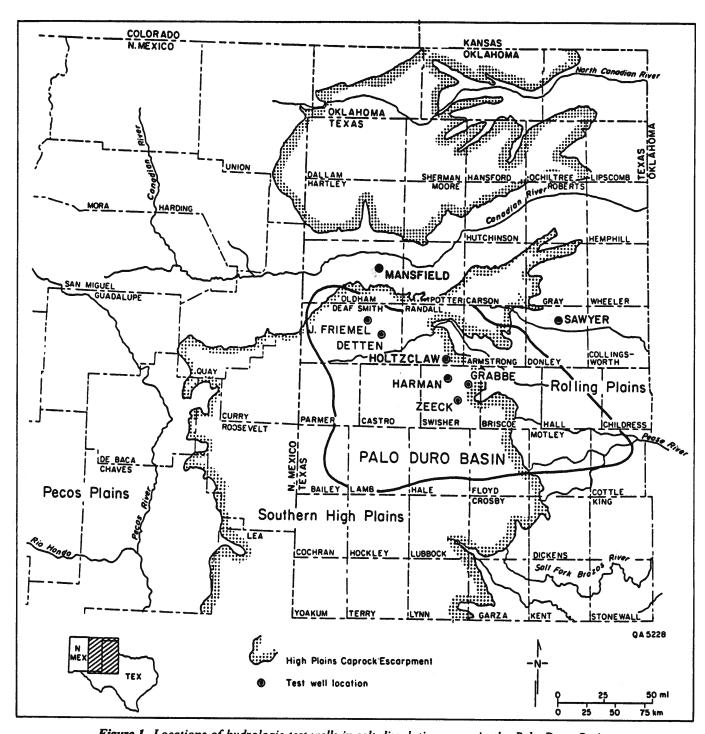


Figure 1. Locations of hydrologic test wells in salt-dissolution zones in the Palo Duro Basin.

completely characterize salt-dissolution zones throughout the Palo Duro Basin.

Four wells were drilled for hydrologic testing and geochemical sampling of salt-dissolution zones in the Texas Panhandle: the Stone and Webster Engineering Corporation (SWEC) Sawyer No. 2, SWEC Mansfield No. 2, SWEC Detten No. 2, and SWEC Harman No. 1 wells (figs. 1 and 2). The objectives of field activities at

these wells were to obtain and chemically analyze uncontaminated, representative samples of ground water from the salt-dissolution zones and to conduct drawdown and recovery tests to determine hydrologic properties. In addition, drill-stem tests in salt-dissolution zones at the SWEC Holtzclaw No. 1 and SWEC J. Friemel No. 1 test wells (figs. 1 and 2) were conducted to measure permeability and hydraulic head. Data from the SWEC

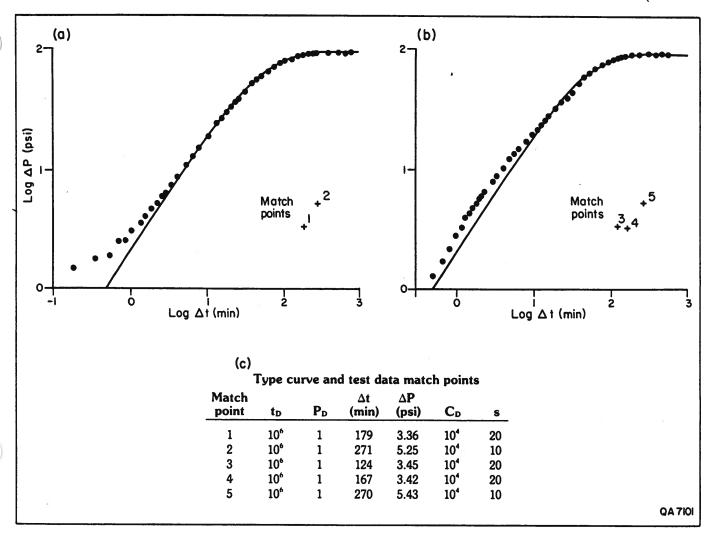


Figure 9. Logarithmic plots of hydrologic test data from the SWEC Sawyer No. 2 well. (a) Water-level drawdown data. (b) Water-level recovery data. (c) Type-curve and test-data match points. Results of calculations shown in table 2.

calculated and actual values of wellbore storage indicates that no open dissolution caverns are hydrologically continuous with the well. From equation 6,

$$C_D = \frac{CE_w}{2\pi nhr^2} = \frac{(0.794) (300,000)}{(2\pi) (0.3) (22) (0.2813)^2}$$

$$C_D = 10^{4.9}$$

The data trace matches type curves with C_D values of 10⁴ to 10⁵.

Hydraulic conductivity was calculated by matching a plot of test results with a type curve for radial flow to a well with wellbore storage and skin effects (Agarwal and others, 1970, fig. 1). Match points are shown in figure 9. The positive skin effect shown by the match possibly reflects partial penetration of the well into the unit 4 carbonate or partial penetration into the salt-dissolution

zone, which includes the unit 4 carbonate and the overlying collapse breccia. Hydraulic conductivity was not estimated by the Jacob semilogarithmic approximation (Kruseman and De Ridder, 1976, p. 59-65) because wellbore storage influenced data throughout test duration.

Wellbore storage factor changed during the early part of the recovery period (fig. 9b) as a result of water draining out of the production tubing above the pump; the pump lacked a check valve. The match of a type curve with data from the recovery period is less certain than the match with drawdown data, but the results of both matches are similar (table 2).

SWEC Mansfield No. 2 well

Test data from drawdown and recovery periods at the SWEC Mansfield No. 2 well match type curves for slight wellbore storage and negative skin effects (fig. 10). Negative skin effect probably reflects the stimulation of fluid inflow that occurred during drilling and well

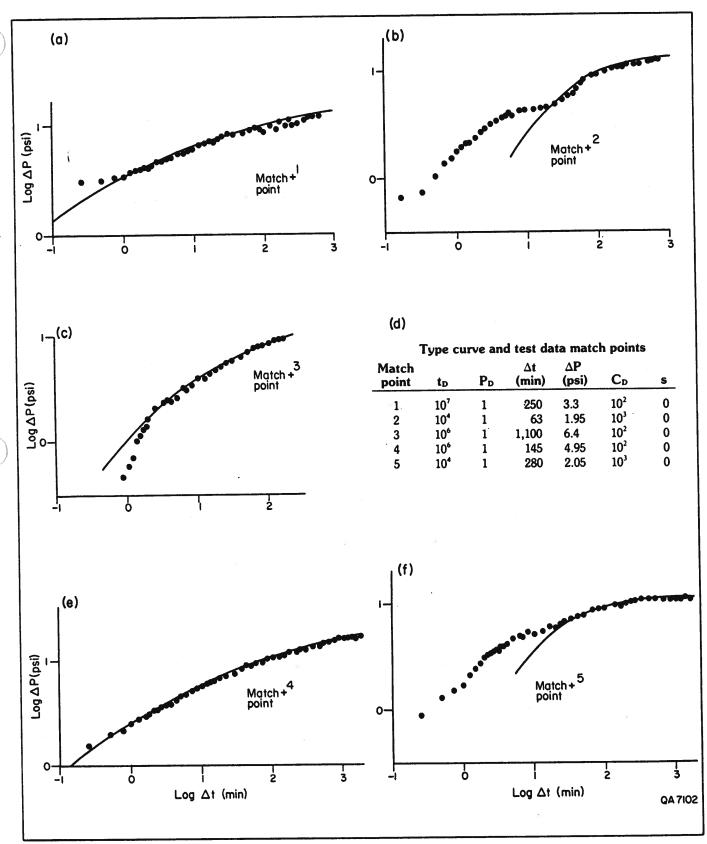


Figure 10. Logarithmic plots of hydrologic test data from the SWEC Mansfield No. 2 well. (a) Water-level drawdown data, test no. 1. (b) Water-level recovery data, test no. 1. (c) Water-level drawdown data, test no. 3. (d) Type-curve and test-data match points. (e) Water-level drawdown data, test no. 4. (f) Water-level recovery data, test no. 4. Results of calculations shown in table 2.

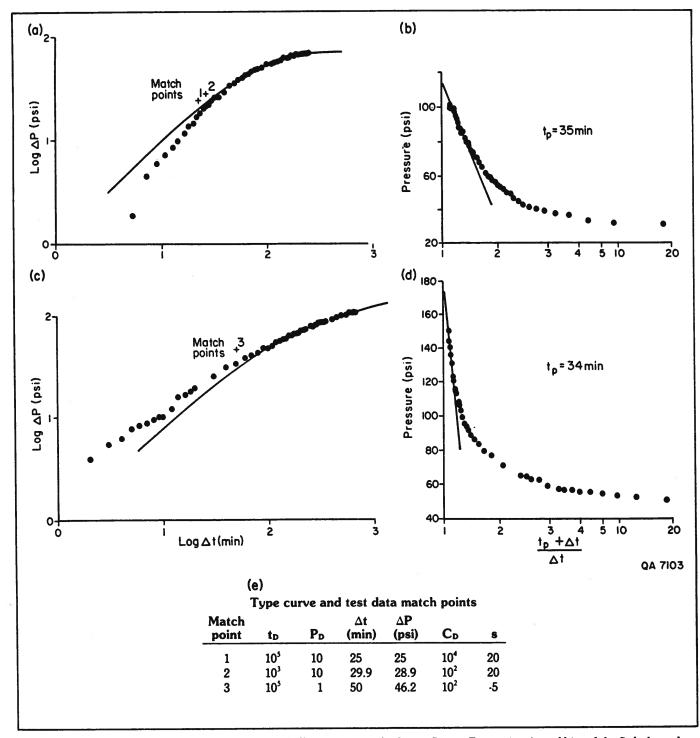


Figure 11. Plots of pressure-buildup data from drill-stem tests in the Seven Rivers Formation (a and b) and the Salado and Tansill Formations (c and d) at the SWEC Holtzclaw No. 1 well. (e) Type-curve and test-data match points.

estimate is only about ± 25 psi (± 0.17 MPa) and accuracy of hydraulic head is only about ± 50 ft (± 15 m).

Data from the drill-stem test in the Salado and Tansill Formations at the SWEC Holtzclaw No. 1 well match a type curve (Agarwal and others, 1970, fig. 1) with a skin effect value of -5 (fig. 11c), indicating that fluid inflow was

probably stimulated by an enlarged wellbore diameter caused by caving of the production interval. Calculated permeability is about $0.04 \text{ md} (4 \times 10^{-17} \text{ m}^2)$, and hydraulic conductivity is about $2 \times 10^{-4} \text{ ft/d} (6 \times 10^{-5} \text{ m/d})$. Formation-fluid pressure was extrapolated to be 175 psi (1.19 MPa) (fig. 11d); final shut-in pressure was 150 psi

$\frac{Q}{(m^3/d)}$	b (m)	PD	ΔP (psi)	s/cycle (m)	$T (m^2/d)$	K (m/d)	Method
			SWEC S	awyer No.	2 well		
58.875	6.71	1.0	5.25 & 3.36	-	2.4 to 3.6	0.4 to 0.6	type-curve match
58.875	6.71	1.0	5.43 & 3.42	-	2.3 to 3.7	0.4 to 0.6	type-curve match
			SWEC Ma	ansfield No	. 2 well		
49.55 49.55	16.76 16.76	1.0	3.3	- 2.83	3.2 3.2	0.19 0.19	type-curve match Jacob approximation
49.55	16.76	1.0	3.7	-	2.8	0.34	type-curve match Theis approximation
53.73	16.76	1.0	4.15	-	2.7	0.11	type-curve approximation Jacob approximation
58.57	16.76	1.0	4.95	-	1.5	0.16	type-curve approximation Jacob approximation
58.57 58.57	16.76 16.76	1.0	2.47	3.102	5.0 3.5	0.4 0.19	type-curve approximation Theis approximation
	58.875 58.875 49.55 49.55 49.55 49.55 53.73 53.73 58.57 58.57	58.875 6.71 58.875 6.71 49.55 16.76 49.55 16.76 49.55 16.76 53.73 16.76 53.73 16.76 58.57 16.76 58.57 16.76 58.57 16.76	\$8.875 6.71 1.0 \$8.875 6.71 1.0 \$49.55 16.76 1.0 \$49.55 16.76 - \$49.55 16.76 1.0 \$49.55 16.76 - \$3.73 16.76 - \$3.73 16.76 - \$58.57 16.76 1.0 \$58.57 16.76 - \$58.57 16.76 1.0	SWEC S 58.875 6.71 1.0 5.25 & 3.36 58.875 6.71 1.0 5.43 & 3.42 SWEC Ma 49.55 16.76 1.0 3.3 49.55 16.76 49.55 16.76 1.0 3.7 49.55 16.76 53.73 16.76 53.73 16.76 1.0 4.15 53.73 16.76 58.57 16.76 1.0 4.95 58.57 16.76 1.0 4.95 58.57 16.76 1.0 2.47	SWEC Sawyer No. 58.875 6.71 1.0 5.25 & - 3.36 58.875 6.71 1.0 5.43 & - 3.42 SWEC Mansfield No. 49.55 16.76 1.0 3.3 - 49.55 16.76 2.83 49.55 16.76 3.32 53.73 16.76 53.73 16.76 1.0 4.15 - 53.73 16.76 53.73 16.76 53.73 16.76 53.73 16.76 53.73 16.76 58.57 16.76 1.0 4.95 - 58.57 16.76 58.57 16.76 1.0 2.47 -	SWEC Sawyer No. 2 well 58.875 6.71 1.0 5.25 & - 2.4 to 3.36 3.6 58.875 6.71 1.0 5.43 & - 2.3 to 3.42 3.7 SWEC Mansfield No. 2 well 49.55 16.76 1.0 3.3 - 3.2 49.55 16.76 2.83 3.2 49.55 16.76 1.0 3.7 - 2.8 49.55 16.76 3.32 2.7 53.73 16.76 3.32 2.7 53.73 16.76 1.0 4.15 - 2.7 53.73 16.76 3.25 3.0 58.57 16.76 1.0 4.95 - 1.5 58.57 16.76 3.38 3.2 58.57 16.76 1.0 2.47 - 5.0	SWEC Sawyer No. 2 well 58.875 6.71 1.0 5.25 & - 2.4 to 0.4 to 3.36 3.6 0.6 58.875 6.71 1.0 5.43 & - 2.3 to 0.4 to 3.42 3.7 0.6 SWEC Mansfield No. 2 well 49.55 16.76 1.0 3.3 - 3.2 0.19 49.55 16.76 2.83 3.2 0.19 49.55 16.76 1.0 3.7 - 2.8 0.34 49.55 16.76 3.32 2.7 0.16 53.73 16.76 3.32 2.7 0.16 53.73 16.76 1.0 4.15 - 2.7 0.11 53.73 16.76 3.25 3.0 0.18 58.57 16.76 1.0 4.95 - 1.5 0.16 58.57 16.76 3.38 3.2 0.19 58.57 16.76 1.0 2.47 - 5.0 0.4

Table 2. Summary of hydrologic tests.

Q = flow rate; b = test-zone thickness; P_D = type-curve pressure match point; ΔP = data match point; s/cycle = straight-line slope on semilogarithmic plot; T = transmissivity; K = hydraulic conductivity

SWEC Holtzclaw No. 1 well: Salado and Tansill Formations

0 1

 8×10^{-4}

0.007

 6×10^{-5}

development, when a large amount of sand was produced, enlarging the test-zone diameter. The unit storage factor estimated from SWEC Mansfield No. 2 well diameter is

0.37

0.62

test 1

test 2

14.04

14.04

10.0

25.0

46.2

$$C = 0.975 \text{ ft}^3/\text{psi} = 0.442 \text{ ft}^3/\text{ft}$$

and the wellbore storage coefficient predicted from equation 6 is

$$C_D = 10^{4.6}$$

However, data traces best match type curves with CD values of 101 to 102. This indicates that wellbore storage did not strongly influence water-level response rate during drawdown periods. Semilogarithmic approximation of the radial flow equation was used for data analysis because wellbore storage had only a small effect during the test. Water emptying from the production tubing into the wellbore early in recovery period also influenced test results at the SWEC Mansfield No. 2 well (fig. 10). The Jacob semilogarithmic method (Kruseman and De Ridder, 1976, p. 59-65) for analysis of long-term data was used to estimate transmissivity from plots of drawdown versus the logarithm of elapsed time; the Theis recovery method (Kruseman and De Ridder, 1976, p. 66-69) was used with data from recovery periods. Approximations from recovery data used a straight line

segment of data from late in each test after the production tubing was drained. Estimates of hydraulic conductivity using data from different tests and different analytic methods are similar (table 2).

type-curve approximation

type-curve approximation

SWEC Holtzclaw No. 1 well

Figure 11a shows the match between a type curve (Agarwal and others, 1970, fig. 1) and pressure-buildup data from the drill-stem test in the Seven Rivers Formation at the SWEC Holtzclaw No. 1 well. The data best match a type curve with a skin effect value of +20. This positive skin value possibly reflects a damaged wellbore face that retards fluid inflow; this is common among drill-stem tests in low-permeability formations. Assuming the test interval equals the transmissive zone in the formation, calculated permeability is about 0.4 md $(4 \times 10^{-16} \text{ m}^2)$ and hydraulic conductivity is about 2×10^{-2} ft/d (7 × 10^{-3} m/d). The semilogarithmic plot (fig. 11b) of pressure data from the shut-in test suggests that undisturbed formation-fluid pressure is about 115 psi (0.78 MPa). The final shut-in pressure was 107 psi (0.74 MPa). Hydraulic head determined from the extrapolated pressure is 2,588 ft (788.8 m) msl, using the specific weight of drilling mud (0.535 psi/ft; 0.012 MPa/m) that flowed into the drill string. Because the drill-stem test data were influenced by wellbore storage (fig. 11a), accuracy of the extrapolated pressure

Index to TAB 13: Ground-water quality data for the study area

A. Tabulations which include more than one aquifer

Berkstresser and Mourant, 1966, Table 3 (parts)

HGC, 1984a, Appendix B, Table B.1 (parts)

HGC, 1984a, Appendix B, Whittemore report, Table 1, p. 172

- B. Tabulations limited to aquifers in single geologic unit
 - 1. Alluvium

USBR, 1984, (Hydrology/Hydrogeology Appendix), Tables 3-20

Analyses of samples collected 10/18/83 from 3 piezometers; much of data also on Table B.1, above; HGC, 1984a p. 162-164

Extracts of soil samples from tributaries, Ashby Lewis 1973 river survey

Analysis of a sample of alluvial sand collected in 1969 for chloride

Analyses of samples from piezometers drilled as part of USBR early work, from CRMWA files

2. Triassic aquifers

HGC, 1984a, Table 6

Analyses of samples collected 9/22/83 from 2 Triassic wells; much of data also on Table B.1, above; HGC, 1984a p. 159, 166

3. Brine aquifer

Summary of brine aquifer water quality data

4. Permian aquifers

Analysis of a sample collected 9/22/83 from Dripping Springs; much of data also on Table B.1, above; HGC, 1984a p. 158

TAB 13 Part A. Tabulations which include more than one aquifer

GROUND WALLS

SPRINGS IN QUAY COUNTY, N. MEX. TABLE 3. CHEMICAL ANALYSIS OF WATER FROMELIS

(Analyses by U.S. Geological Sunggenical constituents are in parts per million)

Location number: See explanation in text. (S) preceding number denotes spring in this table only. Depths and surface. Reported depths are given to nearest foot. Measured depths are given to nearest tenth of a foot. Explanation:

Stratigraphic unit: Qal, younger alluvium; Qc, upland cover of older alluvium; To, Ogallala Formation; R_{(elaceous} sandstone and shale; Jm, Morrison Formation; Je, Entrada Sandstone; Rc, Chinle Formation; R sc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Je, Entrada Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; Rc, Chinle Formation; R_{(elaceous} Sandstone) and Sandstone; R Santa Rosa Sandstone; Pr, Permian rocks.

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THE PRESIDE BUNEAU OF MINES & MINERAL RESUDENCES

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* Composite sample of 7 wells. Depth when sampled.

TYDRO GEO CHEM, INC.

FINAL REPORT

LAKE MEREDITH SALINITY INVESTIGATION

Table B.1: Explanation

a. Type: W = Well

S = Surface Water

Sp = Spring

b. Formation: Pl16 = Piezometer 1, 16 ft depth Pl22 = Piezometer 1, 22 ft depth P222 = Piezometer 2, 22 ft depth P240 = Piezometer 2, 40 ft depth P255 = Piezometer 2, 55 ft depth P320 = Piezometer 3, 20 ft depth P335 = Piezometer 3, 35 ft depth P415 = Piezometer 4, 15 ft depth P421 = Piezometer 4, 21 ft depth P621 = Piezometer 6, 21 ft depth P631 = Piezometer 6, 31 ft depth P650 = Piezometer 6, 50 ft depth Rev = Revuelto Creek CaSL = Canadian River at State Line CaTa = Canadian River at Tascosa CaAm = Canadian River at Amarillo CalM = Lake Meredith Tr = Triassic PSA = San Andres PY = Yeso PW = Wolfcampian

c. Source: 1 = Collected this study

2 = USGS WATSTORE

3 = Water resources data for New Mexico and Texas, various years

4 = Berkstresser and Mourant, 1966
5 = Bureau of Reclamation files
6 = Griggs and Hendrickson, 1951

7 = Bassett and Bentley, 1983

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FINAL REPORT	Canadian B	TDS	E8	15246	15779	16037	14530	15500	14273	15564	22947	24714	11145	12644	7573	13521	19135	20600	20783		290	1060	1540	363	1030	1910	889	1360	1620	1300	040	1530	1100	416	2150	1350	351	1310	632	1260	2 5	1/9	2530	1110	
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HYDRO GEO CHEM, INC	Table B.1: Selected water-quality	Location 1		.33.15.13	12 32 14 210		13.33.14.210	13.33.14.210	33.14	33.12	13.33.12.230	.33	.33.12	.33.12	34	.34	3.34	13.34.05.120	3.34		22 23	13 33 24 13	12 12 71 12	3 6	33	33	33	13.33.21.13	.33	13.33.21.13	E. (13.33.21.13		33	33	.33.21	.33.21	.33.21	.33.21	13.33.21.13	.33.21	3.33.21.1	.33.21.1	13.33.21.13	17.6

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% %	- T	7	7	7	1	· 7	٠ ٦	•	7	7	7	7	• 7	1 .	7	7	7	7	7	• • • • •	1	7			7.8	11.0	1.2	•) t	- 6	7.3	16.0	2.9	7.	. H	7		1.3	7.0	3.6	7	7	7	9		>	17.0
5	250	270	260	280	900	9 6	1 6		300	310	290	410			340	330.	330	310	9 40		2 6	270			‡	143	53	101	101	9 6	2 5	210	116	300	31	238	111	73	42.5	255	51	59	9	· *	;;	4 6	2 6	70
204	250	250	260	280			2 6	007	300	290	280			200	310	270	280	280			200	210			120	581	196		•	671	494	1100	249	1760	555	96	215	9	009	548	170	7	140	7	3 6	700	505	
603	0.0	0.0				9 6	9 6	9 .	0.	0.	0.0		•	2	0.	0.4	0,0	0		9 (9	0.											7.9															
H003	206	212	117	112	1 0	017	117	202	202	219	108		117	204	707	206	210	1 00		2 3	180	160			388	257	727		9	372	204	824	572	824	œ	852	496	229	355	161	314	220	320) ¢	276	710	445	60 60 60
M	9 1-1			7.0	•		•	7.9	6.9	1.9		•	7.1	7.5	7.8	8.2	•				9.9	& 89			7	7	1	•	7	7	7	7	7	7	7	7	7	. S	6.4	7.1	7	' '		•	7 '	7	7	7
N.	250			007	207	287	7.80	780	.2 80	290		207	290	300	300	300	300		200	330	320	250			212	227		10	424	136	518	1200	307	1370	304	970	239	23	6	205	75	1	1 9	h +	7 ;	265	219	49
N.	;	3 4	9 ;	9 5	17	7.7	73	64 68	79	5	9 0	27	73	30	30	9		1 5	17	78	33	21	i		8.7		`		0.	79	2.4	6.3	43	4.5	54	5.0	38	21	9 5	190	32	; ;	4 6	ה ה	31	34	31	34
5	,	g :	T ;	79	20	79	63	55	26	9	7 4	20 1	8	61	80		; ;	2 :	:	6	57	46	:		14	2:	,,,	7	9. 6	25	m	13	29	9.5	82	~	45	61	210	9) C	**	» ·	e i	26	93	30	88
20		242	9	72	30	09	20	70	80	3	3 :	9	8	50	20		2 6	3 ;	9	99	20	90	3		10	1 6	2 :	70	8	531	110	110	1050	000	070	270	808	951	17.5			֓֞֞֜֞֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓	7 ;	511	363	912	973	411
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Pa. b		Cell	Calk	Call	Call	Calin	Cali	Call	1	1	Calk	Cell	CALM	7.10	1			Celk	Calk	Cali	7	2		1	T OF	IL	ī	Tr	Ţ		ř	ŀ											-	Ţ	Ţ	Ţ		H
Date			1 3/13	6/14/73	10/18/73	4/14		. 5	76/66/	C1 /77/1	5/28/75	1 9/75	/21/76	96/8/	767767	0/ /47/9		8/18	2/29/78	2/ 4/80	2/11/81	10/04/	79 /67 /7		Trinssio	12/10/52	7/30/48	8/48	8/88		7/15/48	2/17/1	0/27/48	10/11/48	10/13/48	10/13/48	04/00/01	00/100/1	20/17/	9/27/83	9/22/83	2/25/55	3/26/63	3/26/63	4/10/70	0/15/47	10/15/47	0/15/47
4 9		7 2	1 5/			3/	3					8	-	• •					7	-		•	12	ı	Ţ	W 12	7	₩ 10/	W 12/					, ,			; ;			s '	× '	7	···	~***	•	W 1(; -	i
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8 TIIV		N03		17.0	0.7	0.3	O.3	42.0	4.3	0.1	0.7	0.0		7		7.0		7	7.0		7		2.8	7		7	7	7	7
Lake meredith salinity study	<u>s</u>	ឆ		•	*	•	6	11	•	35	3.5	RJ.		2270		43719		26800	27435		19700		\$760	6580		88154	71600	04000	85566
EREDIT		204		19	1610	130	1130	3 6	102	113	34	37		1840		\$250		2810			7660		1450	1710			2350		
CAKE		6 03	٠			0.0	•					0.0		7		0.0		0.0			0.0		0.0				0.0		
		HCO3		302	152	331	186	294	218	469	185	330		994		904		887	159		830		1280	9201		302	278	173	106
		M		7	7	7	7	7	7	7	7	7		7		9		62.9	73.0		51.7		54.8			7	7	7	7
		N.		14	18	9	43	18	74	247	3.9	16		1510		29000		19640	17 500		17940		4240	7880		60310	40900	26200	47193
		Kg	10	22	135	78	16	12	32	3.4	4.	12		102		610		245			182		102	96		1069	1350	1540	841
ORT		ಶ		89	486	29	351	92	4	7	9	42		890		1360		192	800		624		449	384		1100	3660	1960	6578
FINAL REPORT		TDS		295	2350	475	1720	374	322	645	202	334		7100		80948		49180	49072		36406		11985	21138		164026	120000	172000	142121
_		8.C.		511	2550	723	1990	575	\$33	1040	358	539		10000		70650		\$0000	00069		49600		17500	17800		7	7	7	7
		ьн		7	7	7	7	7	7	7	7	7		ï		0.9		7	4.		7		7	7		æ.	6.	œ.	æ.
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		Fa.b	:	3 3	<u>,</u>		<u>,</u>	<u></u>	<u></u>	<u>,</u>	, pe	F				M/P		/Ir	/Ir		Tr		/Ir	JI.	T a	F	-	F	E
			-	10 I	17 E	17 E	47 E	17 F	17 F	47 F	(7 E	47 E		21 2		83		83 F	83		83 E		83 E	83 E	for	8	8	8	2
NC EN	per	Date	Pormian Volla	San Miguel Count 9/26/47 Psa 6	8/04/47	8/06/47	10/24/47	8/08/47	10/22/47	5/20/47	7/24/47	6/20/47	Voll	12/11/51	Vol1	9/11/83		7/19/83 P/Tr S	10/17/83 P/Tr 4		7/19/83 P/Tr 5		5/06/83 P/Tr 5	7/19/83 P/Tr 5	amp aquifor in Te	0/00/00 M	0/00/00 M	D/00/00 PT	/00/0
EK.	nt in	Typea				_	1	_	1	-				=======================================				-	W 10				_		GRED	_	_	_	_
8	. S	ħ		9 0 1	110	001	001	110	001	001	90	001	TOOL	132	Spri	120		130					130	130	7015				
20 20 20 20 20 20 20 20 20 20 20 20 20 2	B.1.	lon		We I.	04.1	11.2	36.1	22.1	26.1	33.1	16.3	27.1	=	30.2	23	25.1	1-3	15.1	15.1	7-1	15.1	H-2	07.2	07.2	in	0	0	0	9
HYDRO GEO CHEM, INC	Table B.1: Continued	Location		Vator volls in	12,12,04,110	12.12.11.200	12.12.36.100	12.14.22.110	13.12.26.200	15.12.33.100	16.12.16.300	18.15.27.100	Ray No 1 HOOVER	11.28.30.232	Drippi	13.31.25.120 W	Well 0W-3	13.33.15.130	13.33.15.130	Well 0W-4	13.33.15.130	Well DH-2	13.34.07.230	13.34.07.230	Wells in Wolfor	0258000	0635000	0643000	0651000

Table 1. Concentrations and ratios for Canadian River Basin samples.

\bigcirc	Sample No.	Sample description	Date collected	Cl _. mg/L	Br mg/L	I , ug/L	Br/Cl x 10 ⁴	I/C1 x 10 ⁶
	1,A	spring	9-21-83	25.2	0.15	2.3	60	92
	2A	Ute Res.	9-21-83	55.2	0.20	41	36	750
	3A	SALT WELL	9-22-83	1,130	0.25	30.3	2.2	26.7
	4A	REVUELTO CREEK	9-22-83	255.	0.50	32.6	4.4	28.8
		DRIPPING SPRINGS	9-22-83	38,800	9.7	63 .	. 2.5	1.6
	7A	TRIPUSIC SPRING	9-22-83	57.6	0.26	2.5	45	43
	18	UTE DAM	10-17-83	319	0.35	84	11	263
	2B	0W-3	10-17-83	26,600	5.5	70	2.1	2.6
	3B	PIEZ 6	10-18-83	10,600	3.0	49	2.8	4.5
	4B	PIEZ 3	10-18-83	12,800	3.3	51	2.6	4.0
	5B	PIEZ Z	10-18-83	7,340	1.8	19.6	2.5	2.7

Estimated error in concentrations above: Chloride, less than $\pm 2\%$

Bromide, ±0.05 mg/L for concentrations less than 0.5 mg/L ±10% for concentrations greater than 0.5 mg/L

Iodide, from ±5% for higher concentrations up to ±10% for low concentrations

TAB 13 Part B.1 Alluvium

	7-1654 (7-73)	•	COMPUTAT	ION SHEET		-
_ ,,	8Y	9/26/84	PROJECT	+ Merelita	Project	SHEET / OF 4
	CHKD BY	DATE	FEATURE		+ Revuel	
	DETAILS Statis	tical Analy	ses - Re			
	<u> </u>	- All S				+ BX
iter Sangles,	flow/chlor		R squor stl. Erro Intercept	(R)= (R)= (R)= (H)= (H)= (B)=	0.078 22.949 17.061	
For Grouped Wa	flow / Field	·	R Squar	ed = = = = = = = = = = = = = = = = = = =	0.095	
istical Analyses	flow / TO	S	Sample Hi correlati Psquared Stil Error Intercep Slope	(on (pp) = l = of Est. =	80 -0.286 0.082 22.897 17.966 -0.002	
mmary of Stat	9	Field Conducta	Sample Sample Correlet R sque Std Erro Intercap Slope	ion (K) = ~L = ~d Est. = + (K) =	80 0.936 0.877 775.969 -348.662 0.352	
Talle 3- Su	P'- Variance	· closserto I betton line	Somple Cortelat R squar Std. (Erro Interce, Slope a Kelstandir; -	tion (R) = ref Est = of (R) = (B) = indicate inverse;	80 0.869 0.755 1095.103 241.498 0.471	
	std. Error - std.	ies of y values fo	son predicte.	x y values		

Water and Powe

COMPUTATION SHEET

BY	CATE	PROJECT LK Meredith Project SHEET 2 OF 4
CHKD BY	DATE	FEATURE
DETAILS S te	tistical A	alyses - Regression y = A+BX

Field Conductance / TDS Sample une = 80 correlation (P) = 0.877 R squared = 0.768 5th. Error of Est. = 2828.171 Intercept (A) = 2140.874 slope (B) = 1.266

Group II - Alluvial Piezometers Canal. Rab Revuelto Cr.

flow/chlorides

temple time = 1/9 Correlation (R) = 0.238 R squared = 0.057 Stl Error of Est. = 1.454 Intercept (A) = 0.842 5/092 (B) = 0.00017

flow / Field Gaductance

Longle tipe = 1/9 Correlation (R) = 0.209 P squared = 0.044 Std. Error of Est. = 1.464 Intercept (A) = 0.883 Slope (B) = 0.00006

flow/TDS

tomple sine = 119

cortelation (K) = 0.067

R squared = 0.004

Stl. Error of Est. = 1.494

Intercept (A) = 1.935

Slope (B) = 0.00002

Note: Flow data is from the surface Strem near the prigometers.

7-1654 (7-73) Water and Power

COMPUTATION SHEET

BY	DATE	PROJECT LK Meredith Project	SHEET 3 OF W
CHKO BY	DATE	FEATURE	12
DETAILS	Statistical A	nalyses - Regression	Y = A+ 8X

```
chlorides/Field Conductance Sample size = 1/9

Correlation (12) = 0.819

Resquared = 0.670

Std. Error of Est. = 1235.800

Intercept (A) = 845.918

Slope (B) = 0.338
```

Group III Alluvial Prezometers, Revuelto Cr. and Canal. R. Blow Revuelto Cr. Y=A+BX

Flow/chlorides

COMPUTATION SHEET

BY	DATE	PROJECT CK Meredith Pro	iect SHEET 4 OF 4
CHKO BY	DATE	FEATURE	·
DETAILS Stat	istical A	Analysis - Regression	Y=A+BX

flow/Field Conductance Samplesine = 67 Correlation R = 0.030 R Squared = 20009 Sta. lervor of Est. = 35.629 Intercept (A) = 13.441 Slope (B) = 0.0001

flow / TDS

Annale Size = 67

Correlation (R) = 0.019

R squared = 0.0004

Std. error of 5st. = 35.639

Intercept (A) = 14.014

Slope (B) = 0.0001

chlorides/Field Conductance Sample Size = 67

Correlation (R) = 0.974

Regulared = 0.948

5tl. (Error of Est. = 836.877

Intercept (A) = -533.114

Supe (B) = 0.375

Chlorides / TDS

Sample time = 67

Correlation (R) = 0.858

R = 9 and = 0.737

STEP (A) = 85.5//

Slope (B) = 0.47/

Field Conductance/TDS Sample size = 67 Correlation (18) = 0.862 Kesquared = 0.743 STAL Error of Est. = 4851,193 Intercept (A) = 1920.625 Slope (B) = 1.227

Water Quality Analysos, USBR (1984)

SITE 1 River (Mile 1.6): Table 4.

	Flow	(cts)		0.0	9.	6,	5.1	9.	4.	4.4	1 .6	7 .	4.	6 .	7:	2.0	9.	1.7	1 .0	1.2	<u>.</u>	5.0
Field	Water	Temp.			2	සි	8	83	35	5 8	19.0	13	8.9	0	9	17	8	8	27	29.0	30.0	22
Field	onductivity	nmhos	9500	10700	9500	4400	4250	9200	10490	9500	9400	10100	10000	11800	13800	9200	10780	10000	9328	12160	11348	7776
LAB Field	Conduct. C	nmhos			9200	4500	4250	0006	0006	0006	8840	9400	11000	9050	11950	8200	11080	10240	. 10088	10700	11540	9273
	LAB PH	pH Units		8.06	7.84	8.16	8.33	7.85	7.92	7.78	7.97	7.7	7.99	7.97	7.76	7.93	8.13	7.79	7.86	8.03	7.72	8.4
Total Dissolved	Solids	l/gm	5828	5438	5634	2354	2500	5623	5775	5857	5643	6201	6222	5396	8000	5722	6711	5947	5286	6411	8609	5613
	arbonate	₩ J					9.0				0			0			0			0		0.1
	Bicarbonate Carbonate	/gm			396.5		304.8				406.26			431.9			427.0			378.2		390.8
	Sulfate	√gm			830		1120				745			387.6			475			587.5		691
	Chloride	√gm	2320	2880	2750	1000	1128	3000	3350 *	2950	2684	3100	3100	2740	4100	3000	3560	3500	3000	3696	3680	2923
	Potassium	l/gm		10.2	·		6.80	•			12.0			9.24			14.6			12.0		10.8
	Calclum			131.3			8				136	•		128			148	!		122.4		124
	Magnesium			53.4			31.2				84.8			95.2			72.0			5.76		57.1
	Sodium	l/gm	2024	1984			1604				2172	1		2304			2186	}		2281	į	2079
	Date	¥	05-13-83	05-23-83	06-07-83	06-22-83	07-07-83	07-26-83	08-24-83	09-28-83	10-26-83	11-21-83	12-13-83	01-19-84	02-15-84	03-14-84	04-18-84	05-15-84	06-08-84	07-19-84	08-14-84	MEAN

* Value adjusted by USBR.

NOTE: Mean values are calculated from data not reproduced from USBR report.

Water Quality Analyses, USBR (1984)

Field	Water Temp.		,	2	5	16	17	17	17	16	15	15.6	15	14	5	4	16	17	17	17	91
	Conductivity	19000	22500	24500	25000	25500	26000	26000	25000	24000	24000		21500	23400	22000	24880	22080	22400	21640	21084	23360
	Conduct. (24000	24900	25500	23800	22800	22100	24700	22500	22250	23700	21000	21500	24700	23700		21400	22200	23172
	LAB pH pH Units		7.92	7.71	7.70	8.01	7.74	7.51	7.54	7.70	7.54	7.50	8.01	7.60	7.42	7.99	7.64	7.63	7.66	7.51	7.7
Total	Solids Solids	16252	16072	15849	15665	16942	15949	15867	15408	16123	15440	15216	15034	14980	15439	14801	15123	15030	16477	14457	15585
_	Carbonate mg/l		0			0				0			0			0			0)	0.0
	Bicarbonate mg/l		467.2			487.96				529.48			536.8			523.38			453.84		499.8
	Sulfate mg/l		320			352				920	}		R25	}		882			950	}	719
	Chloride S mg/l	7720	6920	8350	0006	8880	0006	9500	8400	8360	8250	8250	8160	8100	9200	8360	8500	8500	8880	8200	8465
	Potassium mg/l		21.1			21.4	; ; }			23.7			910	2		23.6			18.4		21.7
	Calctum mg/l		357.2			308	}			326	3		406			380	3		325.6	0533	338.5
	Magnesium Calcium mg/l mg/l		128.2			134 4				9 acc	200		Š	3		254 4			906	3	195.3
neter: Table 5.	Sodium mg/l	2600	5920			6840				6330	0366		0000	0220		5249	2430		E417	<u> </u>	5795.1
SITE 1 – 16' Plezometer: Table 5.	Date	05-13-83	05-23-83	06-07-83	06-22-83	07-07-83	07 26 83	20 27 00	00-24-03	20-02-60	10-20-03	20-17-11	12-13-03	01 - 13 - 04	02-13-64	00-14B-84	04 1 18 184 05 16 84	+0-01-C0	00-00-04	0/ - 19-04	MEAN

Water Quality Analyses, USBR (1984)

Field	Water Temp.		!	ភ	.	16	16	17	17	16	15	15.6	5	15	15	15	16	15	17	11
Field	Conductivity umhos	22500	24500	22000	23000	21500	25000	21000	19500	20000	17200	18000	17900	20000	20900	20680	20800	18348	20400	18420
	Conduct.			23000	23250	21500	22900	19800	16500	17300	17000	19000	19500	19510	20000	22200	20950	20000	22000	21160
•	Lab pH pH Units		7.82	7.60	7.73	8.0	7.63	7.58	7.62	7.65	7.52	7.70	7.92	7.64	7.62	7.95	7.63	7.62	7.84	7.60
Total Dissolved	Solids mg/l	14502	15737	14738	15029	14948	15248	13411	12025	12534	11920	12412	12150.	14131	14923	12897	13439	12723	14135	12824
	Carbonate mg/l		0			0				0			0			0			0	
	Bicarbonate mg/l		634.4			363.56				497.76			523.3			509.96			469.7	
	Sulfate mg/l		830			760				855	**		787.6			795	,		812.5	
	Chloride mg/l	9099	6720	7950	0006	7800	0006	8500	6400	6160	6200	6200	9009	2500	8500	8760	000	2500	0210	8200
	Potassium mg/l	ē.	19.6			18.4				19.6			48.6			910	2	•	16.6	
	Calcium mg/l		341.2			OBC	3			297.6	2		0830	1		352	3		204	3
able 6.	Magnesium Calcium mg/l mg/l		126.3			122.4	1			188.8	2		A C 0 +	106.1		\$ C	2		452 B	0.50
lezometer: T	Sodium mg/l	5440	6160	;		2800	3			0007			200	3		4630	3		7062	900
SITE 1 – 22' Piezometer: Table 6.	Date	0613	05-23-83	06-07-83	06-22-83	07-07-83	20 00 00	00 -20 -00	00-24-00	03-50-60	10-20-03	42 42 62	12-13-03	01 = 13 = 04	02-13-04	031110	101010	101 i CO	00-100-104	08-14-84

Water Quality Analyses, USBR (1984)

SITE 2 River (Mile 2.2): Table 7.

Flow (cfs)		0.45	2.0	න ල	6.1	-:	1.6	1.8	1.9	2.1	2.0	1.9	. .	4.1	1.5	1.7	0.	6.0	1 .3
Fleid Water Temp.			24	ଛ	30.5	8	35	24	18.0	4	3.3	2.0	9	17.5	8	16	58	33.0	8
Field onductivity umhos	15000	0066	10000	4200	4500	10000	10880	10500	9200	10200	12500	0096	14500	0096	10896	10672	96/6	12484	11820
LAB Field Conduct. Conductivity umhos umhos			9250	4400	4500	9200	0006	9100	9180	10000	10000	8850	13700	9400	11510	10368	8656	11300	11980
LAB PH PH Units		8.37	7.85	8.24	8.33	7.94	8.0	7.88	8.0	7.75	8.06	8.01	7.72	7.90	8.03	7.83	7.91	7.96	7.81
Total Dissolved Solids mg/l	5683	2699	5840	2422	2275	5798	6110	6354	5848	6400	6365	5175	9176	5987	7275	6395	6200	6648	6332
Carbons mg/l		1.8			<u>-</u> 5		0#3		0			0			0			0	
Bicarbonate mg/l		381.86			291.6				413.58			435.54			422.12			373.32	
Sulfate mg/l		375			278				450			450			450			200	
Chloride mg/l	2516	2990	2850	1000	1004	2500	3500	3250	2616	3200	3250	2700	4800	3000	3420	4000	2500	3400	
Potassium mg/l		10.6			6.62				12.5			10.5			14.8			12.3	
Calcium mg/l		134.5			89	}		•	136			128			148			123.2	
Magnesium mg/l		53.4			33.6				105.6			100.8			79.2	8		187.7	
Sodium	4. 8.00 8.00	2120			1440	2			2290			2276	i		2242	!		2308	
Date	05-13-83	05-23-83	06-07-83	06-22-83	07-07-83	07-26-83	08-24-83	09-28-83	10-26-83	11-21-83	12-13-83	01-19-84	02-15-84	03-14-84	04-18-84	05-16-84	06-08-84	07-19-84	08-14-84

Site 2 - 22' Piezometer: Table 8.

Pield	Water Temp.			15	15	15	5	15	15	15	15	14.4	4	15	17	14.75	17	17	12	16	
Field	Conductivity, umhos	21000	24000	24900	24500	24000	26000	25000	24500	24000	23500	22100	21250	21500	21000	23400	21000	19948	20760	21080	
	Conduct.			23000	24000	24000	23000	22800	21800	24200	21500	22100	33000	20100	20200	24000	21720	22440	22500	21480	
	LAB PH PH Units	•	7.83	7.65	99.2	16.7	7.64	7.53	7.64	7.94	7.63	7.63	8.12	7.64	7.50	7.90	7.56	7.49	7.79	7.48	
Total	Solide mg/l	13902	15779	15677	15629	16573	15640	15706	15855	17124	15357	14969	14892	14786	14778	15300	14532	14619	14669	14156	
	Carbonate mg/l		0			0				0	•		0	•		0	,		0	•	
	Bicarbonate mg/l		559.6			352.6		•		428.22			391.6	2		470.92			425.78		
	Sulfate mg/l		925			790				1025			200			200	3		787.5	2	
		5720	0099	8300	9500	8480	0006	000	8550	689	8250		222	7600		7800		000	9240		3
	Chloride mg/l		106			4.02	į			780	- -		9	<u> </u>		30 8	3		130	2	
	Potassium mg/l	\$	365.2	1		338	3			350	300		927 G	9.70		368	3		976 B	200	
	alcium mg/l		1477	•		165 G	2			900	6000		7 000	K00.4		*	30		462.5	7.70	
Table 8.	Magnesium Calcium mg/l mg/l	0802	2600	3		CBOS	3			7000	8	3	9	0220		2009	3053		4050	1004	
вхотевет	Sodium mg/l																				
Site 2 – 22' Plezometer: Table 8.	Date	0 0 0	05-13-00	09-23-63	81 62	07.07.83	07 -07	00 07 00	00 00 00	03-62	10-26-83	11-21-83	12-13-83	01-19-84	02-13-64	03-14-84	04-18-84	00 - 10 - 04	00-00-04	10-51-70	08-14-84

SITE 2 - 40' Piezometer: Table 9.

Field Water Temp.	ट	ត ត ត	21	25. 8.35.	16 17 71	2	9 9 9
Field Conductivity umhos	23000 26000 25500	25000 24500	26000 24500	24000 23500 23000	22200 22500 22000	24400 21760 21040	21440 22160
LAB Conduct. umhos	24500	24000 24500 23800	23600	24700 22100 23100	22000 21600 21400	25200 22600 22120	21400 26040
LAB PH PH Units	7.89	7.84 7.88 7.74	7.51	8.00 7.59 7.32	8.06 7.58 7.46	8.0 7.49 7.39	8.04
Total Dissolved Solids mg/l	15409 16838 16369	16159 17133 16096	16279 16181	16391 15903 15728	15920 15719 15701	15596 15219 15036 ·	15011 14881
Carbonate mg/l	, , , , , , , , , , , , , , , , , , ,	c	. 8	0	0	0	0
Bicarbonate mg/l	461.1	336.7		448.96	307.44	463.6	416.02
Sulfate mg/l	980	910		1005	1025	1045	1112.5
Chloride mg/l	6960 7360 8600	10000 8160	9500 8150	7200 8400 8500	8680 8350 9000	7960 9500 7500	9120 8500
Potassium mg/l	23.2	20.7		19.5	21.2	22.6	19.1
Calcium mg/l	418.1	342		377.6	377.6	400	337.6
Magnesium Calcium mg/l mg/l	144.8	183.6		267.2	256	141.6	150.7
Sodium mg/l	5480 5920	5920		6400	0899	5277	5223
Date	05-13-83 05-23-83 06-07-83	06-22-83 07-07-83	07 -26 -83 08 -24 -83 09 -28 -83	10-26-83 11-21-83	01-19-84 02-15-84 03-14-84	04-18-84 05-16-84 06-08-84	07-19-84 08-14-84

SITE 2 - 55' Plezometer: Table 10.

									Total				
					8				Dissolved		Field	Field	
Q.		Magnasium Calcium	Calchim	Potessium	Chloride	Sulfate	Bicarbonate	Carbonate	Solids	LAB PH	Conduct.	Conduct. Conductivity	
	mg/l	/Bw	l∕gm	l/gm	l/gm	l/gm	/bm	√gm	√gm	pH Units	eoumn	eoqun	
0513R3	5040				5920				14296			21000	
05-23-83	5200	164 24	352 4	193	0099	1045	457.5	0	15416	7.96		24000	
06-07-83				2	8000				14838	7.53	22100	23000	
06-22-83	¥()				9200				15273	7.84	23000	23500	
07-07-83	5480	176.4	302	19.7	8160	940	339.2	0	15832	7.85	22500	22500	
07-26-83	3	:	}		0006				15074	7.69	21900	22000	
08-24-83					0006				14946	7.54	22200	24000	
00-28-83					8300				15458	7.48	21000	24000	
10-26-83	6160	259.2	361.6	17.4	7180	962.5	411.14	0	15947	8.05	23900	23500	
11-21-83	3		}	:	8200				15053	7.48	22000	23000	
12-13-83					8150				14992	7.28	22000	21200	
12-13-03	7600	264	344	20.55	2960	865	412.36	0	15123	7.95	21700	20900	
07 - 15 - 84	3	5	}		2002				14857	7.47	20200	23000	
02-13-04					8000				14843	7.38	20000	21100	
00 - 10 - 10 - 10 - 10 - 10 - 10 - 10 -	4872	144	372	23.8	7440	066	445,3	0	14857	7.78	24100	23000	
0F 16 BA	1	-			8500				14419	7.39	21680	21200	
00-10-04					8500				14398	7.62	22200	20044	
07-19-84	4949	158.4	315.2	18.9	8880	1450	419.68	0	14172	7.98	22600	20560	
DB-14-84					7700				14249	7.59	21420	20296	
5													
MEAN	5614	194	341	8	7642	1042	414	0	14200	7.2	20769	21290	

Water Quality Analyses, USBR (1984)

SITE 3 River (Mile 5.4): Table 11.

Flow (cfs)	8.8	3.1	9.9	5. 8.	2.4	6 .	7 .	, S	2.7	2.5	4:2	2.2	2.3	25	2.8	2.4	2.0	4. 3.
Field Water Temp.		ಜ	8	52	8	58	£	15	=	5.5	0	ເດ	14	16	17	2	52	8
Field anductivity umhos	7500	17000	7100	6500	16250	17000	19000	18800	20000	19500	18000	22900	17500	19526	18520	16168	22760	21640
LAB Field Conduct. Conductivity umhos umhos		16150	8250	6500	15800	16500	17000	18300	18500	18000	15000	20900	16100	20400	18784	15836	20500	21340
LAB pH pH Units	8.09	7.77	8.19	8.20	7.90	7.98	7.81	8.14	7.65	7.82	8.25	7.65	7.82	7.98	7.82	7.91	7.94	7.74
Total Dissolved Solids mg/f	5674	10420	1630	4109	10400 *	10740	11570	10600 *	12665	12324	9299	14652	11336	12111	11895	10385	14294	13275
D Carbons mg/l	0	•		0				0				0		0			0	
Bicarbonate mg/l	397.7			305.0				462.38			373.3			483.92			434.32	
Sulfate mg/l	610	;		388				2150			290			895			1150	
Chloride mg/l	4600	5550	2500	1976	0009	0009	0009	5320	6750	6550	4860	7800	6500	6560	6500	5500	8280	7200
Potassium mg/l	281	!		8.90				18.6			14.6			22.4			20.8	
Calcium mg/l	1714			88				203.2			206.4			236	1		225.6	
Magnesium mg/l	S.	i i		48				203.2			147.2	!		98.4			123.8	
Sodiumdium Mg/l	4240 3808			1752				5040	}		4052			4229			5213	
Date	05-13-83	06-07-83	06-22-83	07-07-83	07-26-83	08-24-83	09-28-83	10-26-83	11-21-83	12-13-83	01-19-84	02-15-84	03-14-84	04-18-84	05-16-84	06-08-84	07-19-84	08-14-84

Site 3 - 20' Piezometer: Table 12.

Field Water Temp.	. 4	<u>4 4 t</u>	र र क	18.9	តិ សិ សិ ជុំ	5 & 4 #	15.2 15.2	
Field Conductivity umhos	20000 26000 24200	24000	23100 18500	17000 18800 16400	17250 17500 17900	19306 19124 19280	17152 17152 19837	
LAB Conduct. umhos	25000	25500	21500 15200	16100	18700 16600 17100	18368 18388 18388	18798	
LAB PH PH Units	8.46	80.08	7.73 7.86	8.18 7.76 7.73	8.29 7.83 7.73	7.71 7.71	7.89 7.81 7.93	
Total Disŝolved Solids mg/l	13415 16348 16191	16250 * 16414	14377 11183	10714 11889 12198	11937 11610 11773	11926 12050 11992	12/64 12029 13229	2
Carbonate mg/l	7.98		0	0	0	o	o 6.	
Bicarbonate mg/l	422.61	429.4		412.36	307.44	444.1 F. 8	464.82 413	
Sulfate mg/l	1375	1325		800	715	098	987.3	
Chloride mg/l	4600 6840 8200	9500 8480	9000 7500 5900	4920 6250 6500	5940 6100 6200	8360 7500 6500	7040 6640 * 6946	
Potassium mg/l	23.6	23.1		16.7	20.4	24.7	17.1	
	237	208		112	188.8	%	200	
Sodium Magnesium Calcium mg/l mg/l mg/l	172	165.6		121.6	166.4	200.6	60 85 85	
Sodium A mg/l	4080	0009		4600	2660	4112	4213 4941	
Date	05-13-83 05-23-83 06-07-83	06-22-83	07-26-83 08-24-83 09-28-83	10-26-83 11-21-83 12-13-83	01-19-84 02-13-84 03-14-84	04-18-84 05-16-84 06-08-84	07-19-84 08-14-84 MEAN	, !

Fleid	Water Temp.		*	4	4	4	5	5	15	15	- - 	15.6	16	16	16	: 1	15	14	14	15	14.9
Field	Conductivity umhos		39500	39000	38000	37000	40000	34000	36000	32500	33000	34100	33800	32000	32000	37096	33280	29560	30520	31880	34624
	Conduct.	884		38000	39000	37000	32800	36000	33000	35900	31600	34000	32900	32550	33000	38200	93060	34080	36000	29400	34676
	LAB pH pH Units		8.23	7.53	7.74	7.89	7.71	7.39	7.58	7.68	7.62	7.17	7.83	7.60	7.45	7.58	7.55	7.52	7.73	7.61	7.63
Total Dissolved	Solids mg/l		26106	26319	23000 *	26617	24000 *	25460	25218	25077	25009	24652	25231	24919	24426	25004	23709	23613	24305	24569	24846
	Carbonate mg/l	2	0			0				0			0			0			0		0
	Bicarbonate mg/l		280.7			451.4				585.6			594.14			553.88			485.56		542
	Sulfate mg/l		1720			1540				1125			1525		53.	1435			1875		1537
	Chloride mg/l		10720	13550	14500	14240	15000	15500	13250	11880	13250	13150	13120	13250	12800	11880	12100 *	13500	12320	11800	13101
•	Potassium mg/l		31.0			30.8				35.8			36.6			32.0			56		32
	Calcium mg/l		438.9			392				436.8			448			428			9		424
aole I.S.	Sodium Magnesium mg/l mg/l		202			252				388			382.4			232.4	i i		220		280
	Sodium Img/I	5720	8360		Ü.	0886				. 6560	1		10280			8333			8550		8240
Sile 3 - 34 Flezometer. Table 13.	Date	05-13-83	05-23-83	06-07-83	06-22-83	07-07-83	07-26-83	08-24-83	09-28-83	10-26-83	11-21-83	12-13-83	01-19-84	02-13-84	03-14-84	04-18-84	05-16-84	06-08-84	07-19-84	08-14-84	MEAN

9.13 26.0 6.7 0.01 Flow (cfs) 3.3 7.8 14.3 5.6 0.4 0.1 4.3 ß 24 35.5 17.7 Field Water Temp. 24 LAB Field .. Conduct. Conduct. 2572 2000 1460 1900 1500 1400 1350 1500 2300 4300 3200 4000 8464 568 umhos 3350 2800 4000 1493 7954 628 2275 1750 1800 1300 1300 1180 1550 umhos LAB pH pH Units 8.34 8.26 8.19 8.28 8.17 8.17 8.13 8.39 8.39 7.79 8.56 8.14 8.25 8.2 2294 1492 2294 1911 2618 1397 1051 5339 **a/** 5339 **a/** 1276 1260 1300 * 1508 Solids /gu Total Dissolved Carbona mg/ <u>1</u> <u>5</u> 3.0 0 0 Ξ Bicarbonate 224.0 219.8 256 √gш 240.3 386.74 209.84 Sulfate 175.0 l⁄gm 88 345 **662** 220 476 Chloride /gm 20 68.5 100 240 240 656 925 100 2150 200 165 123 95 100 327 Potassium 6.5 √gш 6.85 5.33 2.4 9.2 8.7 Calcium mg/ 8.09 89.6 78.4 ೪ 72 8 Magnesium mg/l 24.3 46.4 97.6 50.4 5.28 **8** 8 2 8 8 163.2 894 315 333 ľβш 127 Sodiumdium 05-13-83 05-23-83 06-07-83 06-22-83 07-07-83 07-26-83 08-24-83 09-28-83 10-26-83 02-15-84 03-14-84 04-18-84 11-21-83 12-13-83 01 - 19 - 8405-16-84 06-08-84 07-19-84 08-14-84 Date MEAN

a. Note on table says "concentration in pool?"

				•	8														
Field Water Temp.		57	4 :	4	15	9	17	17	15	14.4	12	=======================================	Ξ	=	_	13	*		13.6
Field Conductivity umhos	16500		9200	5250	2000	2750	2750	2550	2500	2800	2750	3000	2750	2964	3132	2800	2892		5082
LAB Conduct. C		11700 *	10000	5250	4300	2600	2600	2310	2200	2200	2560	2700	2700	3250	2874	3220	3080		3972
LAB pH pH Units	7.87	8.04	8.46	8.16	8.09	8.03	8.23	7.91	7.87	8.2	7.85	7.75	8.08	7.78	7.78	8.37			8.0
Total Dissolved Solids mg/l	10955	7973	7150 *	3454	4000	1532	1580	1482	1256	1276	1504	1717	1735	1795	3766	1999	2754		3688
Carbonate mg/l	, .	•		1.8				0			0			0			9.0		0.4
Bicarbonate mg/l		- - -		202				319.64			306.22			284.26			303.78		373
Sulfate mg/l	5	9		515				475			485			320			512.5		220
Chloride ma/l	4320	3850 *	3000	1204	200	320	360	306.8	215	240	398.8	435	425	511	200	650	280		1279
Potassium ma/l		15.5		4.8				<u>ග</u>	<u>!</u>		4.08			5.6			5.2		6.5
Calcium ma/l		302.7		8.4				27.2	!		512	!		72			89.6		85
Magnesium Calcium ma/l ma/l		147.7		3.6				17.6	!		36.8			31.2			40.3		46
Sodium	3640	98		1676				369.6			692	3		516) 		555		1569
Date	05-13-83	05-23-83 06-07-83	06-22-83	07-07-83	07-26-83	08-24-83	09-28-83	10-26-83	11-21-83	12-13-83	01-19-84	02-13-84	03-14-84	04-18-84	05-16-84	06-08-84	07-19-84	08-14-84	MEAN

Water Quality Analyses, USBR (1984)

Field	Water Temp.			5	4	14	4	5	5	5	15	16.7	14	13	12	12	=	13	4		13.7
Field	Conductivity umhos	21000	21000	20600	13000	8000	2000	4400	3200	3000	2800	3300	2800	3800	3700	5532	6452	7464	4280		7868
	Conduct.			20250	13200	8000	2000	4000	3100	2720	2650	3300	2610	3550	2800	5940	5982	8694	4510		6144
	LAB pH pH Units		8.03	7.68	8.14	8.41	8.18	7.96	8.06	8.30	8.04	7.99	8.35	7.95	7.89	8.01	7.79	7.69	8.30		8.0
Total Dissolved	Solids mg/l	13378	13787	14921	* 0006	5291	4400 *	2483	1964	1736	1601	2029	1736	2208	2040	3349	3753	6250 *	3101		5168
	Carbonate mg/l		0			1.2				0			6.0			0			0		4.0
•	Bicarbonate mg/l		564.86			623.9				374.54			305.0			283.04			322.08		417
	Sulfate mg/l		1200) -		615				462.5			492	!		375	,		612.5		929
	Chloride mg/l	5480	e 300 *	6550	4000	2156	1000	200	510	592 *	517 *	520	645	2002	009	1150	1550	2550	1274		2046.9
	Potassium mg/l		19.9			6.89				£.	;		3 52	1		7.8	2		6.1		7.9
	Calcium mg/l		203.4			5	!	81		5	į		97.0	2		88	3		112		75.5
Table 16.	Magnesium mg/l		129.2			4 32				7.0	į		40	2		98	3		42		86.
Piezometer:	Sodium P	4760	4980			2364				4044	5		902	83/		1075	2		972	<u> </u>	2274
Sito 4 - 20.5' Piezometer: Table 16.	Date	05-13-83	05-23-83	06-07-83	06-22-83	07-07-83	07-26-83	00 - 24 - 83	00 - 28 - 83	20 - 62 - 64 60 - 62 - 64	10-20-63	11-21-05	12-13-03	01-13-04	02-13-64	70 - 41 - 60 70 - 41 - 60	06-16-84	05-10-04	07-19-84	08-14-84	MEAN

Water Quality Analyses, USBR (1984)

SITE 6 River (Mile 9.9): Table 17.

20																						
	Flow (cfs)		35.0	=	S S	4. 3.	0.	9	9	17.5	9	2.0	0	က	2.5	ဖ	13.0	7.5	8	201	21.6	
Field	Water Temp.			18.5	27	ឧ	24	52	17	=	₽	3.3	0	 4.0	14	5	16	8	ຮ	2 2	15.9	
Field	onductivity umhos	7500	4235	7250	10000	9100	17000	7200	4750	4600	7800	13800	22000	16800	18000	11555	5228		1758	806	9416	
FAB	Conduct. Conductivity umhos			2000	10000	9100	17000	6200	4250	3650	0006	12175	18880	15200	15000	10820	5536	17020	1586	921	9645	
	LAB PH PH Units		8.30	7.95	8.06	8.15	7.87	8.12	8.05	8.37	7.98	8.0	7.80	7.70	7.72	8.11	8.05	7.64	8.32	7.57	8.0	
Total Dissolved	Solids mg/l	4459	2383	2038	5970	2696	11025	3882	2685	2549	4475	8243	13209	10711	11889	6321	3431	11630	1121	460	5904	
	Carbone mg/l		0			0				2.4			0			0			0		4.0	•
	Bicarbonate mg/l		280.6			305				273.3			662.46			364.78			231.8		353	
	Sulfate mg/)		395			514				387.5			875			220	1		150		479	
	Chloride mg/l	1536	006	2050	2500	3244	6500	1800	1150	966	2050	4150	6400	5550	0009	2900	1500	6500	297.6	250	2962	
	Potassium mg/l		8.6			11.6				8.35			6.61			16.3	<u>}</u>		3.8	}	11.4	
	Calcium mg/l		69	1		116	•			76.8	}		310.4			148	2		32	}	129	
	Magnesium mg/l		37	•		62.4				8	3		8000			818			9		8	
	Sodium mg/l	1370	756	3		2196	3			1100	3		5480	3		2034			296	3	1890	•
	Date	05-13-83	05-23-83	06-07-83	06-22-83	07-07-83	07-26-83	08-24-83	00-24-00	10-26-83	11-21-83	12-13-83	01-19-84	02-15-84	03-14-84	04-14 18-84	05.16.84	00-01-00 00-00-00	07-19-84	08-14-84	MEAN	

Site 6 - 21' Piezometer: Table 18.

Field Water Temp.	4		1 5	<u>হ</u>	16	5 !	17.7	<u>+</u>	16	4 :	4 ;	<u>र</u>	ត្ !	2	14.9
Field Conductivity umhos	13000 12800 13000	12000	13000	11000 4000	10800	12000	11900	15000	15100	15200 *	17660	17852	15260	16200	13214
LAB Conduct. umhos	13000	11250	11200	9250 9700	9370	11200	12000	14000	13900	18500	16782	17048	18000	18916	13672
LAB pH pH Units	8.16 7.65	8.06	7.86	7.7	8.10	7.74	7.49	7.74	7.56	7.78	7.58	7.42	7.9	7.53	7.8
Total Dissolved Solids mg/l	7752 7694	7197	7116	6366 6483	6999	7470	8296	9807	10454	10876	10706	11842	11397	11598	8816
Carbonate mg/l	0	c	•		0		(9		0			0		0
Bicarbonate mg/l	666.1	R 808			484.34			513.62		472.14			447.74	*	519
Sulfate mg/l	260	29	}		687.5			018		635			887.5		691
Chloride	4000 * 3720	3500	3500	4000	3400	3700	4200	4840	5500	5200	2200	2200	2300 *		4308
Potassium mg/l	13.8	6	<u>?</u>		13.8			19.5		25.1		-	19.4		17.5
Calcium mg/l	91.3	Ş	3		97.6	:		244.8		312			337.6		197
Sodium Magnesium Calcium mg/l mg/l mg/l	49.5	\$	₽		101.6			376		230			142.1		158
Sodium N mg/l	2792 2792	- 6	2070		3244	3		4120		3505			3784		3295
Date	05-13-83	06-07-83	07-26-83	08-24-83	10-26-83	11-21-83	12-13-83	01-19-84	03-14-84	04-18-84	05-16-84	06-08-84	07-19-84	08-14-84	MEAN

Water Quality Analyses, USBR (1984)

Fleld	Water Temp.		4	15	15	15	15	15	15	9	0	1 5	14	16	15	16	15	15	5	14.4
Field	Conductivity	22000	24000	22000	21500	22700	21800	21500	19500	22900	21500	19000	18500	20000	20584	20100	21004	19420 *		21167
	Conduct.		22500	22250	21500	20100	20050	19200	20400	19500	19950	22200	18300	18000	22400	19204	19376	19350	20800	20299
	LAB pH pH Units	7	7.58	7.73	7.92	7.62	7.74	7.64	7.87	7.51	7.70	8.01	7.65	7.74	7.69	7.58	7.63	7.76	7.63	7.72
Total Dissolved	Solids mg/l	14160	15200	14719	14356	14020	14056	13770	13545	13511	13618	13740	13226	13035	13201	12659	13300 *	12173	12035	13651
	Carbonate mg/l	€ (>		0				0			0			0			0		0
	Bicarbonate mg/l		288.0		230.7				679.54			707.6			684.42			583.16		629
	Sulfate mg/l		0/6		935				1280			1170			765			1100		1037
	Chloride mg/l	6040	5920 7750	8500	7280	8000	8000	7050	• 0099	0069	6950	6840	999	e 0099	6400	6500	6500	6864	e 200 •	6937
×	Potassium mg/l		28.2		. 24	i			25.6			7.4.7	:		27.1	: i		19,3		24.8
	Calcium mg/l	!	254.7		260				266.4			9 CPC			276	i		187.2	!	256
able 19.	Sodium Magneslum mg/l mg/l		148.6		151.2	!			67.2	!		216	2		2122	!		128.6		154
ezometer: T	Sodium P mg/l	4920	2260		4360				6160	3		5800	3		4519	2		4365		5055
Site 6 – 31' Piezometer: Table 19.	Date	05-13-83	05-23-83	06-07-183	07-07-83	07-26-83	08-24-83	00-28-83	10-26-83	11-21-83	12-13-83	01-19-84	02-13-84	03-14-84	04-18-84	05-16-84	08-08-84	07-19-84	08-14-84	MEAN

SITE 6 - 50' Piezometer: Table 20

Field Field Conduct. Conductivity umhos umhos	30100 32000 31500	30000 31000 30000	30000 31800 28000 28000	30218
Field Conduct.	30000	30000 29200 27500 27200	30300 28500 32000	29230
LAB pH pH Units	7.73 7.24 7.49	7.89 7.47 7.18 7.42	7.76 7.21 7.28 7.89	7.5
Total Dissolved Solids mg/l	20432 20800 * 20224	20846 20426 20275 20418	20825 20590 20077 18600 •	20319
Carbonate mg/l	· •	0	• • _•	0
Bicarbonate mg/l	555.1	570.0	786	189
Sulfate mg/l	1285	1365	1625	1472
Chloride mg/l	10750 * 10150 11300 *	10400 11100 * 11500 10550	10700 10700 10450 10280	10716
Potassium mg/l	43.2	39.2	4.2. 5.5.	42.5
Calcium mg/l	456.5	75.2	426.4	363.5
Magnesium Calcium mg/l mg/l	163.2	43.2	383.2	235.8
Sodium mg/l	5920 7720	8160	8360 8280	7688
Date	05-13-83 05-23-83 06-07-83 06-22-83	07-07-83 07-26-83 08-24-83 09-28-83	10-26-83 11-21-83 12-13-83 01-19-84 02-15-84 03-14-84 04-18-84 05-16-84 06-08-84	MEAN

AGRICULTURE

VEAL ANALYSIS

PETROLEUM



LABORATORIES

HGC, 1984a

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3016 UNION AVE BAKERSFIELD, CALIFORNIA 93305

MAIN OFFICE 4100 PIERCE ROAD, BAKERSFIELD CA 91308 PHONE 327-4911

Hydro Geo Chem, Inc. 744 N. Country Club Tucson, Arizona 85716 Date Reported: 12/2/83 Date Received: 11/7/83 Laboratory No.: 12517

Attention: Mr. John Ward

WATER ANALYSIS

Sample Description:

Constituents	Parts/millio
Calcium	580.
Magnesium -	145.
Sodium	6,900.
Potassium	35.
Carbonate	0.
Bicarbonate	849.
Chloride	10,832.
Sulfate	1,440.
Nitrate	(∠) 0.4
Total Dissolved Solids, By Summation	20,783.
Boron	1.5

(4) refers to "less than"

B C LABORATORIES, INC.

BY J. Egli

Field Parameters

Well: Piezometer site 6:50'

(Channel deposits)

Location: 13.34.5.12
Date: 10-18-83
Time: 1000
Temp: 18.0°C

pH: 6.60

Alkalinity: 803 mg/l

kc

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AMALYSIS

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LABORATORIES

HGC, 1984 a

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3016 UNION AVE BAKERSFIELD, CALIFORNIA 93305

MAIN OFFICE 4100 PIERCE ROAD, BAKERSFIELD CA 93308 PHONE 327-4911

Hydro Geo Chem, Inc. 744 N. Country Club Tucson, Arizona 85716 Date Reported: 12/2/83 Date Received: 11/7/83 Laboratory No.: 12519

Attention: Mr. John Ward

WATER ANALYSIS

Sample Description: #5B

Constituents .	Parts/million
Calcium -	360.
Magnesium	130.
Sodium	4,800.
Potassium ·	19.
Carbonate	0.
Bicarbonate	255.
	7,788.
Chloride	920.
Sulfate	(4) 0.4
Nitrate Ry Summation	
Total Dissolved Solids, By Summation	1.0
Boron	1.0

(4) refers to "less than"

B C LABORATORIES, INC.

BY J. J. Eglin

kc

Field Parameters

Well: Piezometer site 2:55'

(Channel deposits)

Location: 13.33.14.24

Date: 10-18-83

Time: 1400

Temp: 17.0°C

pH: 6.95

Alkalinity: 325 mg/l

Spec. cond. $(25^{\circ}) = 25,000 \mu mhos$

ETROLEUM

MICAL ANALYSIS



HGC, 1984 a

LABORATORIES 🖂

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3016 UNION AVE BAKERSFIELD, CALIFORNIA 93305

MAIN OFFICE 4100 PIERCE ROAD, BAKERSFIELD CA 93308 PHONE 327-4911

Hydro Geo Chem, Inc. 744 N. Country Club Tucson, Arizona 85716 Date Reported: 12/2/83 Date Received: 11/7/83

Laboratory No.: 12518

Attention: Mr. John Ward

WATER ANALYSIS

Sample Description: #4B

Constituents	Parts/million
Calcium	420.
Magnesium •	170.
Sodium	8,700.
Potassium .	28.
Carbonate	0.
Bicarbonate	520.
Chloride	13,275.
Sulfate	1,600.
Nitrate	(4) 0.4
Total Dissolved Solids, By Summation	24,714.
Boron	1.0

(4) refers to "less than"

B C LABORATORIES. INC.

BY Fali

kc

Field Parameters

Well: Piezometer site 3:35'

(Channel deposit)

Location: 13.33.12.23

Date: 10-18-83

Time: 1130 Temp: 18.0°C

pH: 6.90

Alkalinity: 452 mg/l

Spec. cond. $(25^{\circ}) = 44,500 \, \mu \text{mhos}$

Extracts of soil samples from tributaries collected by Ashby Lewis Canadian River Survey, 01/29/73 and 03/01/73

SAMPLE IDENTIFICATION	6 - Z	8-11	S-12	8-13	8-15	5-17	8-8	8 	2-2	N-7	8-e	9 Z	8-5	8 - 8	e N	X 4	8-2
РНОТО	-	၈	4	* 10	•	107	9	~	61	81	ន	233	83	23	22	23	
RIVERMILE	1.65	9.0	11.6	15.6	18.4	24.8	24.8	4	42.5	46.5	\$	8	Z	55.4	55.8	56.5	
NAME	e :			Tusco- coillo Canyon				Horse	Truj- iilo Creek	Minne – osa Creek	Los Arches Creek				Pedriza Creek		
PARAMETERS																	
Chloride, mg/l	8	400	320	6	1025	350	425	2 5	22	22	150	8	220	6	26	12	
Sodium, mg/l	6	275	275	28	919	230	9	2	107	೭	29	119	94	167	8	5	
Sodium chloride, mg/l	82.5	0.099	577.5	247.5	1691.2	577.5	701.2	206.2	123.7	123.7	247.5	165.0	907.5	165.0	82.5	123.7	
Sulfate, mg/l	8	52	172.5	107.5	330	240	515	185	1225	375	87.5	185	332	310	8	3	

NOTES

Sampling points listed in downstream order from Ute Lake. Samples with N prefix collected in a tributary entering river from the north; those with a S prefix collected in a tributary entering river from south.

Some river mile designations are approximate, based on large - scale maps; they may be changed as small - scale maps become available.

There were two S-17 stations marked on photos and only one S-17 analysis. Ashby Lewis indicated Occobber 15, 1991 that 'S-17' marked on photo 9 probably is the location where a surface water exported as S-17 probably was collected at the 'S-17' marked on photo 10. The location of the soil sample extract reported as N-6 was not marked on any photo; Ashby Lewis indicated that it was probably collected from a tributary on photo 23.

SANd

SAMPle

TEXAS WATER QUALITY BOARD DOMESTIC SAMPLE INFORMATION

51,410
District I County New Mexico Bosin CANADIRA Date 11-2.5-6
Pertaining to CANADIAN RIVER Permit Page
Time NCON Weather Clear Flow 1500 GFR
First on other
Material Sampled: Raw, Primary, Partially Treated, Final Effluent, Stream Material Sampled: Raw, Primary, Partially Treated, Final Effluent, Stream PHOTO # PORT AREA Composite Composite
Point of Collection // // / / / / / / / / / / / / / / / /
Observations: NEW ELEXICE, SALT DEPESITS IN SANCE
How rother low due to water next being ilinely
from lite. Flow serpage water from spring
and Cam. FIELD ANALYSIS
Water Temp F pH Dissolved O_2 Turbidity
Chlorine Residual
Date Shipped 11-26-69 Iced Signature Signature Signature
The second secon
Date Received 11-28-69 Lab. No. 132588 Date Completed 222196
pH NH ₃ -N T. S. Solids
Conductivity NC2-N F. S. Solids
0-PO _A V. S. Solids
B.O.D. ₅
And the second s
Chloride - 1.4.mg
Please ANALYZE For Soluble Solts
FORM NO. G-192 (WCH 34)
Phlorides

Analyses of samples from piezometers drilled as part of USBR early work, from CRMWA files

PARAMETERS

	Alkalinity, mg/l Calclum,	Calcium, mg/l	Chloride, mg/l	pH, pH units	Sodium, mg/l	Sodium chloride, mg/l	Sodium chloride, Specific conductance, mg/l umhos/cm	Sulfate, mg/l	TDS, calculated, mg/l	
3-07-74									712	
Canadian River at Highway 54			SAOO			6270.0	12000			
20 ft well			4250			7012.5	14100		¥	
30 ft. well			2650			9322.5	18200	1		
40 ft. well			4700			7755.0	15500			
50 ft. well			4000			0099	14500	(*)		
4-23-74										
Canadian River at Highway 54										
10 ft. well			3750			6187	11700	14		
20 ft. well			7300			12045	2000			
30 ft. well			7850			12852	21000			
40 ft. well			8400			13860	21500			
50 ft. well			7300			12045	21800			
1-21-75										
1 mile +/- below Ute Dam										
20 ft. well	386	825	4150	7.8	2697		11500	438	7813	
1-27-75										
1 mile +/- below Ute Dam										
20 ft. well	904	299	4400	2.7	2860		12000	999	8360	
3-06-75										
Canadian River at Highway 54										
10 ft. well			9800			14520	22800			
20 ft. well			9400			15510	22100			
30 ft. well			8650			14272	21000			
40 ft. well			9050			14832	21000			
50 ft. well			9800			14520	21500			
3-06-75										
Canadian River near TX-NM state line	in									
20 ft. well			1450			2382	4750			

TAB 13 Part B.2 Triassic aquifers

Table 6: Summary of chemical characteristics of 23 samples of Triassic groundwater in the study area

Ion	Range	Avera	ge
	mg/l	mg/l	meg/l
Ca Mg Na K HCO ₃ CO ₃ SO ₄ Cl Br pH	2 - 210 0 - 190 4 - 1370 3.2 - 7.1 88 - 852 0 - 75 56 - 1760 4 - 1130 0.15 - 0.50 6.9 - 9.3 527 - 5520	54 35 366 4.7 449 14 416 152 0.29 7.7	2.7 2.9 15.9 0.12 7.4 0.46 8.7 4.3 0.004

ERICUL TURE

ANALYSIS

ETROLEUM



HGC, 19842

BAKERSFIELD, CALIFORNIA 93305 3016 UNION AVE

MAIN OFFICE 4100 PIERCE ROAD, BAKERSFIELD CA 93308 PHONE 327-4911

Hydro Geo Chem, Inc. 744 North Country Club Road Tucson, Arizona 85716

10/18/83 Date Reported: Date Received: 10/3/83 Laboratory No.: 10730

WATER ANALYSIS

Sample Description: 6. 9/22/83 1600

Constituents

Calcium Magnesium Sodium Potassium Carbonate Bicarbonate Chloride Sulfate Nitrate

Total Dissolved Solids

Boron Silica Hardness as CaCO3

Electrical Conductivity, Micromhos pH:

Parts/million

210. 85. 49. 4.9 0. 355. 42.5 600. 1.3 1,375. 0.08 27.

875. (51.1 gr/gal)

1,610. 7.3

B C LABORATORIES, INC.

ad

Field Parameters

Well: "Logan Cemetary

Windmill" (Triassic)

Location:

13.33.1.43

Date:

9-22-83

Time:

1600

Temp:

18.0°C

6.96

pH:

Alkalinity:

300 mg/l

PETROLEUM

HGC, 1984a

HEMICAL ANALYSIS



·BAKERSFIELD, CALIFORNIA 93305

MAIN OFFICE 4100 PIERCE ROAD. BAKERSFIELD CA 93308 PHONE 327-4911

Hydro Geo Chem, Inc. 744 North Country Club Road Tucson, Arizona 85716

Date Reported: 10/18/83 Date Received: 10/3/83 Laboratory No.: 10728

WATER ANALYSIS

Sample Description: 4. 9/22/83 1030

Constituents

Calcium Magnesium Sodium Potassium Carbonate Bicarbonate Chloride Sulfate Nitrate

Total Dissolved Solids

Boron Silica

Hardness as CaCO₃

Electrical Conductivity, Micromhos

рH

B C LABORATORIES, INC.

ad

Parts/million

140. 190. 205. 7.1 0. 761. 255. 548.

0.4 2,122. 0.10

16. 1,132. (66.1 gr/gal)

2,800. 7.5

Field Parameters

Well:

"Revuelto Creek

Windmill: (Triassic)

Date: Time:

13.33.24 9-22-83

Temp:

1000 18.0°C

pH:

Location:

6.93

Alkalinity:

580 mg/l

TAB 13 Part B.3 Brine aquifer

Summary of brine aquifer water quality data

DH-1

Date	6/75	6/75	6/75	6/75	6/75	6/75	6/75	6/75	6/75	6/75
Parameter										
Chloride, mg/l	11800	11800	12950	2350	16450	18500	16100	15950	17400	16250
iron, total, mg/l	0.80	0.24	0.43	0.08	0.48	0.80	0.27	0.10	0.48	0.29
Sodium chloride, mg/l	19470	19470	21368	3878	27142	30525	26565	26318	28875	26812
Specific conductance, umhos/cm	34000	37000	36000	8600	45000	52000	51000	50000	49000	49000
Sulfate, mg/l	1650	1450	2150	500	2000	1950	1900	500	missing	1750
Notes	1,2	1,3	1,4	1,5	1,6	1,7	1,8	1,9	1,10	1,11

Date	7/75	7/75	5/12/83	7/19/83
Parameter				
Bicarbonate, mg/l			1280.18	1076.57
Calcium, mg/l			448.5	384.0
Carbonate, mg/l	4800	4950	5760	6580
Chloride, mg/l	4800	4950	5760	6580
Fluoride, mg/l			0.6	0.5
iron, total, mg/l	0.24	1.08	1.34	0.19
Magnesium, mg/l			89.0	96.0
NaCl, mg/l	7920	8168		
Nitrate, mg/l			2.8	
pH, lab, pH units			7.84	7.62
Potassium, mg/l			54.8	36.8
Sodium, mg/l			4240	7880
Specific conductance, umhos/cm	16300	16900		
Specific conductance, field, 25°C, umhos/cm			17500	17800
Sulfate, mg/l	1015	1013	1450	1710
Temperature, field, °C				18
Total dissolved solids, mg/l			11985	12138
Notes	1, 12	1, 13	14	14

	Date	7/19/84	7/19/84
Parameter			
Bicarbonate, mg/l		762.5	784.5
Calcium, mg/l		235.2	344.0
Carbonate, mg/l		0	0
Chloride, mg/l		15920	15200
Magnesium, mg/l		200.0	249.6
pH, lab, pH units		7.92	7.90
Potassium, mg/l		37.1	37.3
Sodium, mg/l		9335.0	9125.0
Specific conductance, fie umhos/cm	eld, 25°C,	36000	36000
Sulfate, mg/l		2175.0	2175.0
Temperature, field, °C		18.0	18.0
Total dissolved solids, m	ng/l	27892	26434
Notes		15	15

	-	Date	3/07/78	3/12/78	4/07/78	6/01/87	6/07/87
	Parameter						
	Alkalinity, P., mg/l as CaCO ₃				0		
	Alkalinity, M., mg/l as CaCO ₃				1194		
	Aluminum, mg/l				4.25		
	Bicarbonate, mg/l					88.5	1030
	Calcium, mg/l					128	696
	Calcium, mg/l				2000		
	Carbonate, mg/l					. 0.00	0.00
	Chloride, mg/i		22000	22000	18600	13800	18900
	Copper, mg/l				<0.5		
	Fluoride, mg/l				1.43		
_	Hardness, total, mg/l as CaCO ₃				2300		
-	Iron, mg/l				90		
	Magnesium, mg/l					200	254
	Magnesium, mg/l as CaCO ₃				300		
	Maganese, mg/l				0.030		
	Nitrate, mg/l						
	Nitrate, mg/l as NO ₃				<0.5	0.00	
	pH, pH units		7.10	7.56	7.31	8.20	7.78
	Potassium, mg/l					206	239
	Sodium, mg/l		14300	14300	12900	9050	12100
	Specific Conductance, umhos/cm					40500	43200
	Sulfate, mg/l		9700	8600	2800	1420	2120
	Total dissolved soilds, mg/l				36649		
	Total dissolved solids, 180°C, mg/l					35300	36600
-	Anions + cations, mg/l	#()				24900	35400
	Notes		16	16	17, 18	17	17

_		, ,	
	IVV	_	1

	Date	5/28/87	6/08/87
Parameter	-		
Bicarbonate, mg/l		161	1050
Calcium, mg/l		400	696
Carbonate, mg/l		0.00	0.00
Chloride, mg/l		12100	19400
Magnesium, mg/l		146	224
Nitrate, mg/l		0.00	0.00
pH, pH units		7.80	7.70
Potassium, mg/l		253	231
Sodium, mg/l		8200	12500
Specific Conductane umhos/cm	ce,	38300	43200
Sulfate, mg/l		2820	2180
TDS, 180°C, mg/l		33000	37500
Anions + cations, m	ıg/l	24000	36300
Notes		17	17

	Date	1/24/78	4/25/78	7/19/83	10/17/83	10/17/83	1/19/84	5/28/87	6/01/87	6/03/87	6/08/87	
Parameter	ı											
Alkalinity, field, mg/l					836					679		
Alkalinity, total, mg/l as CaCO ₃										3		
Bicarbonate, mg/l		753		877.83	159		731.53	1210	1080		686	
Boron, mg/l					3.2							
Bromine, mg/l						5.5						
Calcium, mg/l		768		792	800		1576	880	864		832	
Carbonate, mg/l		0.0		0	0		0	0.00	0.0		0.00	
Chloride, mg/l	£3	27000	31500	26800	27435	26600	23760	13500	18100	28000	23700	
del Carbon-13, per mil, PDB					-4.65							
del Oxygen–18, per mil, wrt SMOW					-9.84							
del Deuterium, per mil, wrt SMOW				G.	-71.2 -71.7							
Fluoride, mg/l				0.7								
lodine, ug/l						2						
Iron, mg/l				0.47							۰	
Magnesium, mg/l		273		244.8	88		552.0	380	568	٠	283	
Nitrate, mg/l					4.0>			0.00	0.00		0.00	
Percent modem carbon				ıo	5.72 +/-0.58							
pH, pH units		7.3	6.45					7.40	7.30	7.1	7.58	
pH, field, pH units					6.36							

·	Date	1/24/78	4/25/78	7/19/83	10/17/83	10/17/83	1/19/84	5/28/87	6/01/87	6/03/87	6/08/87	
Parameter												
pH, lab, pH units				7.61			7.78					
Phosphorus, -T, mg/l as P										<0.01		
Phosphorus, -0, mg/l as P										<0.01		
Potassium, mg/l		94.6		62.9	75		65.4	253	250		282	
Sodium, mg/l		17800	20475	19640	17500		22400	9050	12100		15800	
Specific conductance, umhos/cm		67800	20000		00069			23600	50100	53760	53200	
Specific conductance, field, 25°C, umhos/cm				65800	78400		> 60000					
Sulfate, mg/l		2600	2725	2810	2880		2770	3400	2480	2255	2600	
Temperature, field, °C				18.5	19.0		18.5					
Total dissolved solids, mg/l				49180			51005					
Total dissolved solids, 105°C, mg/l	_	49000							8			
Total dissolved solids, 180°C, mg/l	_							48000	48600		49300	
Total dissolved solids, calculated, mg/l			54700									
Total dissolved solids, by summation, mg/l		,			49072							
Total suspended solids, mg/l										47		
Tritium activity, TU					-2.3+/-0.5							
Volatile suspended solids, mg/l										11		
Anions + cations, mg/l		49300						28700	35100		44400	
Notes		17	11	6	ଷ	2	8	17	11	17	17	

¥				
	Date	1/31/78	4/25/78	7/19/83
Parameter	- 4			
Bicarbonate, mg/l		803		1011.95
Calcium, mg/l		576		624
Carbonate, mg/l		0.00		0
Chloride, mg/l		16400	29000	19700
iron, mg/l				0.83
Magnesium, mg/l		322		182.4
pH, pH units		7.30	6.61	
pH, lab, pH units				7.55
Potassium, mg/l		68.4		51.7
Sodium, mg/l		10600	18850	17940
Specific Conductance umhos/cm	9,	46500	48000	
Specific conductance umhos/cm	o, field, 25°C			57000
Sulfate, mg/l		2200	2650	2660
Temperature, field, °C				17.5
Total dissolved solids	s, mg/l			36406
Total dissolved solids	s, 105°C, mg/l	31900		
Total dissolved solids calculated, mg/l	3,		50500	
Anions + cations, mg	y/I	30900		
Notes		17	17	23

NOTES

- 1 From well log in USBR, 1979, Appendix D.
- 2 Sample #11, flowing open hole at 296 ft.
- 3 Sample #12, flowing open hole at 296 ft.
- 4 Sample #13, flowing open hole at 296 ft.
- 5 Sample #14, flowing open hole at 296 ft., after circulating.
- 6 Sample #15, flowing open hole at 316 ft.
- 7 Sample #16, flowing open hole at 336 ft.
- 8 Sample #17, flowing open hole at 356 ft.
- 9 Sample #18, flowing open hole at 356 ft.
- 10 Sample #19, flowing open hole at 356 ft.
- 11 Sample #23, flowing open hole at 356 ft.
- 12 Sample #5, flowing open hole at 516 ft.
- 13 Sample #6, flowing through casing set at about 40 ft.
- 14 Reported in USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21, and HGC, 1984a, Appendix B, Table B.1. Values are from USBR, 1984, except for fluoride, iron and nitrate, which are from HGC, 1984a. Note that samples from DH-2 may not be representative of the brine aquifer, because the well is only cased to about 40 feet (USBR, 1984, Hydrology/Hydrogeology Appendix, p. IV-53 and 56.)
- 15 Reported in USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21. Note that samples from DH-3 were obtained after air—lifting for about 1 hour; the specific conductance had not stabilized at the time of sample collection (USBR, 1984, Hydrology/Hydrogeology Appendix, p. IV 53 and 56.)
- 16 Reported in USBR, 1979, p. 19.
- 17 From CRMWA files.
- 18 Iron in suspension.
- 19 Reported in USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21 and HGC, 1984a, Appendix B, Table B.1. Note that outgassing of CO₂ rapidly changed pH of water samples collected from well OW-3, and probably, other brine aquifer wells (USBR, 1984, Hydrology/Hydrogeology Appendix, p. IV-53). Values are from USBR, 1984, except for fluoride and iron, which are from HGC, 1984a.
- 20 Reported in HGC, 1984a, p. 156,157,161, and Appendix B, Table B.1. Appendix B, Table B.1 includes a different value (69000) for specific conductance than p. 161. Del Carbon-13 and Percent modern carbon (PMC) values reported in USBR, 1984 (Hydrology/Hydrogeology Appendix), Attachment C.
- 21 Reported in HGC, 1984a, p. 172.
- 22 Reported in USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21.
- 23 Reported in USBR, 1984 (Hydrology/Hydrogeology Appendix), Table 21, and HGC, 1984a, Appendix B, Table B.1. Values are from USBR, 1984, except for iron, which is from HGC, 1984a.

TAB 13 Part B.4 Permian aquifers

abricul ture



PETROLEUM



BAKERSFIELD, CALIFORNIA 93305

MAIN OFFICE 4100 PIERCE ROAD, BAKERSFIELD CA 93308 PHONE 327-4911

Hydro Geo Chem, Inc. 744 North Country Club Road Tucson, Arizona 85716

Date Reported: 10/18/83 10/3/83 Date Received: Laboratory No.: 10729

WATER ANALYSIS

Sample Description: 5A 9/22/83 1400

Constituents

Calcium Magnesium Sodium Potassium Carbonate Bicarbonate Chloride Sulfate Nitrate Total Dissolved Solids Boron Silica Hardness as CaCO₃

Electrical Conductivity, Micromhos pН

B C LABORATORIES, INC.

ad

Parts/million

1,360. 610. 29,000. 64. 0. 904. 43,719. 5,250. (<) 0.4 80,948. 3.5 5,913. (345 gr/gal)

Field Parameters

6.6

Well:

"Dripping Springs"

(Permian-Yeso?)

Location:

70,650.

13.31.25.12

Date:

9-22-83

Time:

1400

Temp:

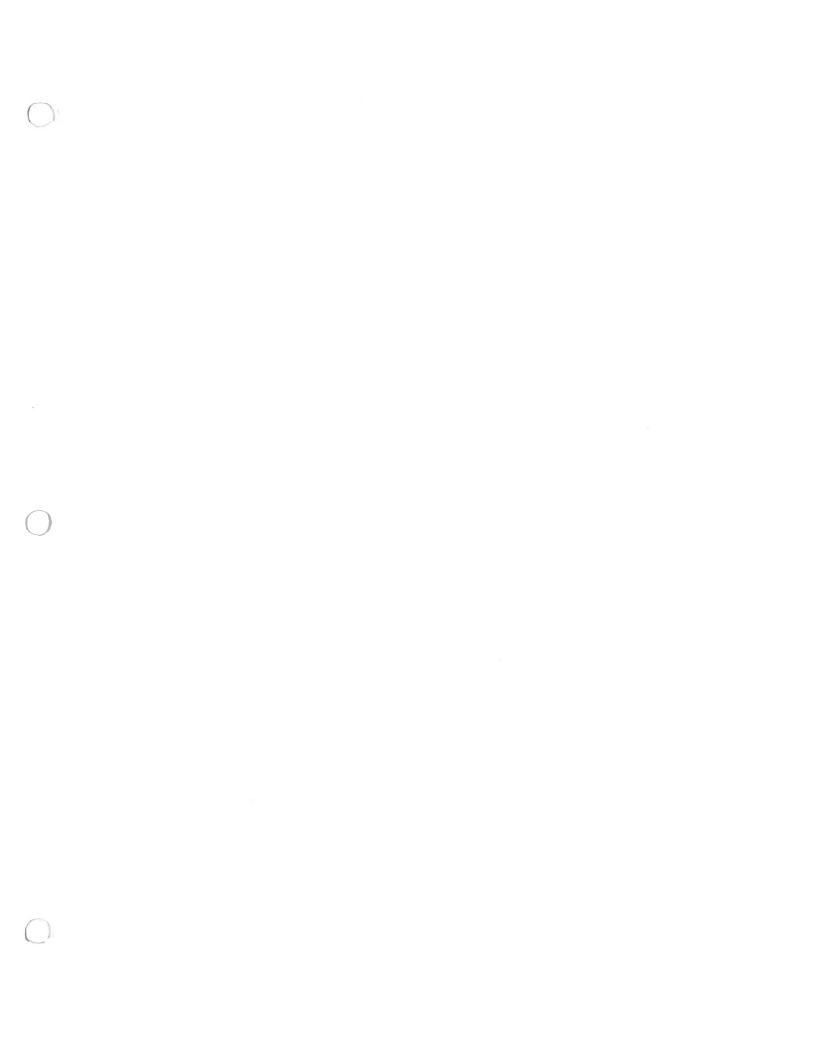
22.8°C

pH:

6.02

Alkalinity:

 $765 \, \text{mg/l}$



Index to TAB 14: Maps, charts and graphs related to water quality

A. Figures related to several aquifers

HGC, 1984a, Figure 43

HGC, 1984a, Figure 44

HGC, 1984a, Figure 45

HGC, 1984c, two unnumbered figures

USBR, 1979, Figure 3

B. Figures related to aquifers in single geologic unit

1. Alluvium

HGC, 1984a, Figure 34

HGC, 1984a, Figure 35

HGC, 1984a, Figure 36, 20 and 21

USBR, 1984, Figures 7-14

USBR, 1984, Figures 25-42

USBR, 1979, Figure 2

2. Triassic aquifers

HGC, 1984a, Figure 32

USBR, 1984, Figure 61

3. Brine aquifer

HGC, 1984a, Figure 33

USBR, 1984, Figure 62

4. Permian aquifers

HGC, 1984a, Figure 30

HGC, 1984a, Figure 31

Figures related to pre-Leonardian aquifer(s)

Bassett and Bentley, 1983, Figure 11

TAB 14 Part A. Figures related to several aquifers

+ 01 × 10/4B

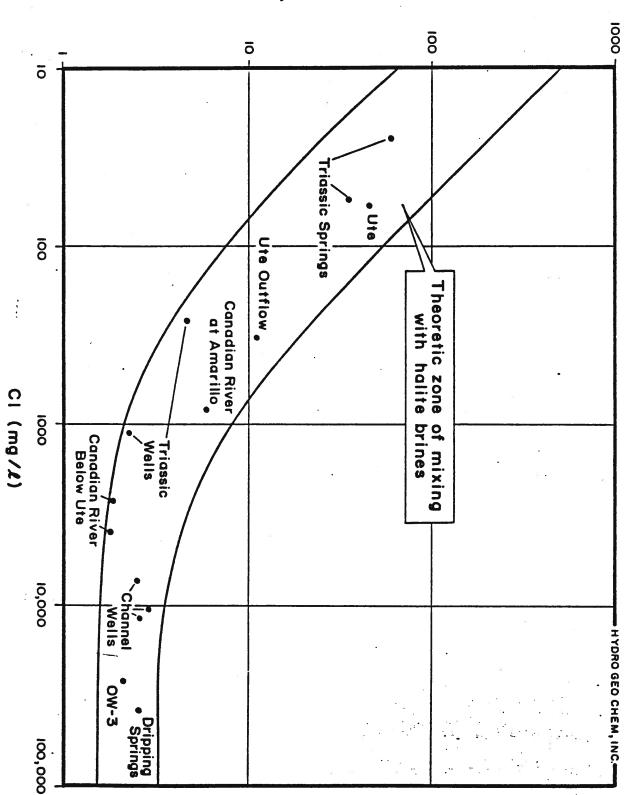


Figure 43. Bromide/chloride ratios of surface and groundwater

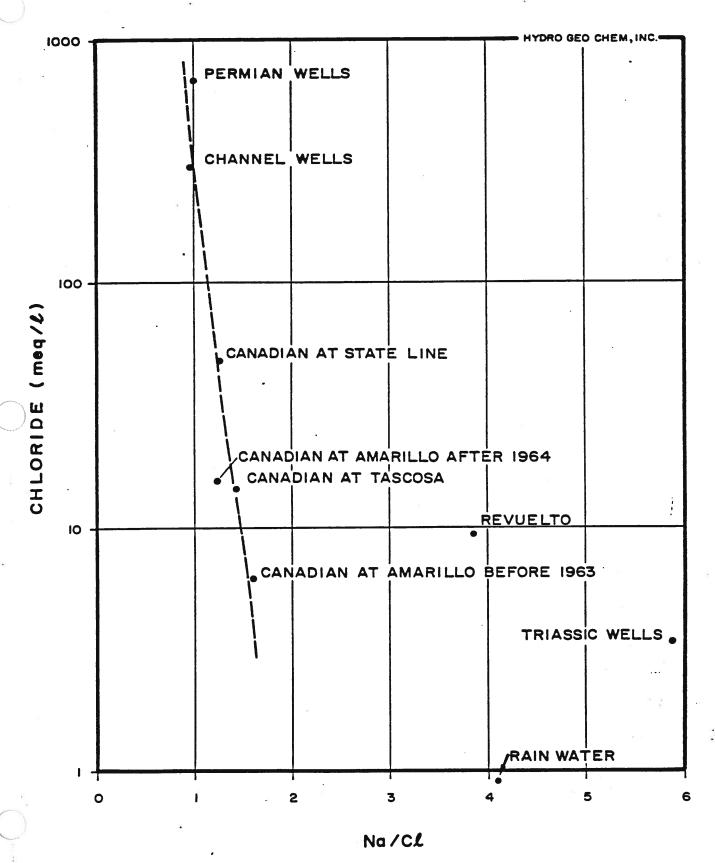


Figure 44. Sodium/chloride ratios of surface and groundwater

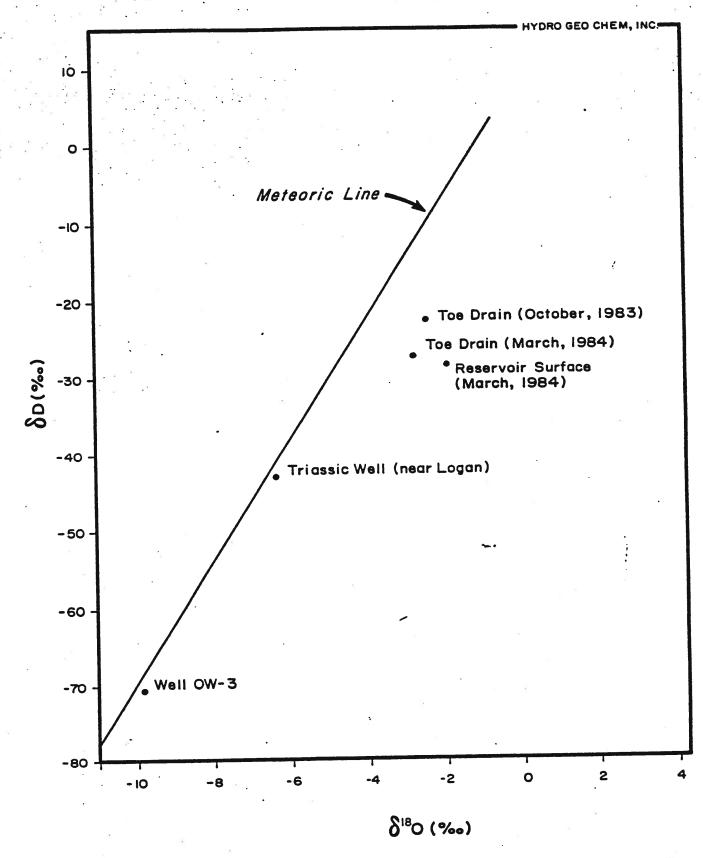
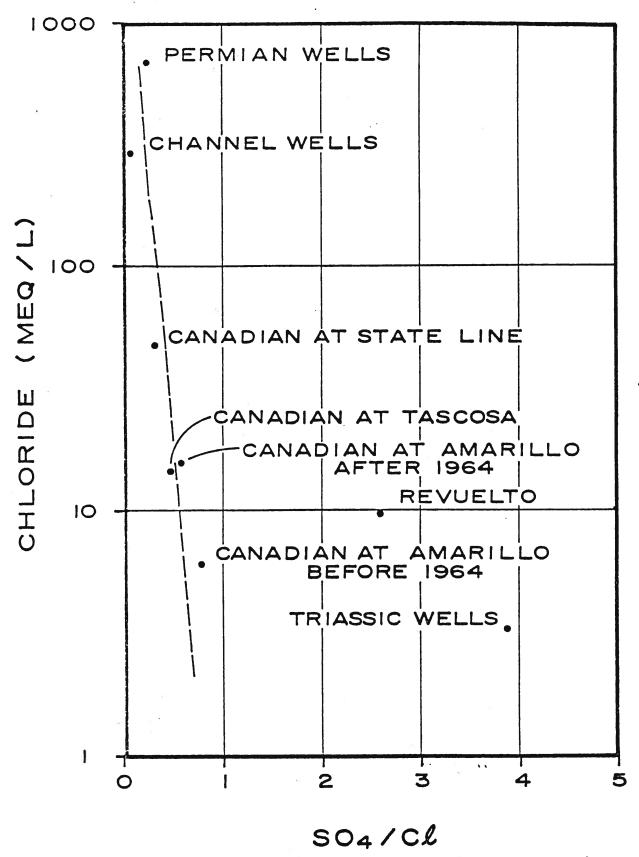
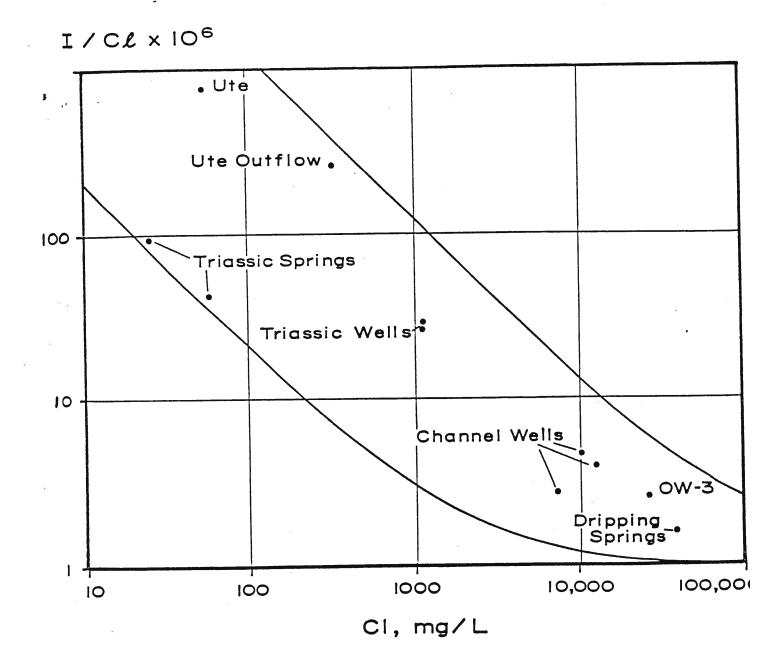


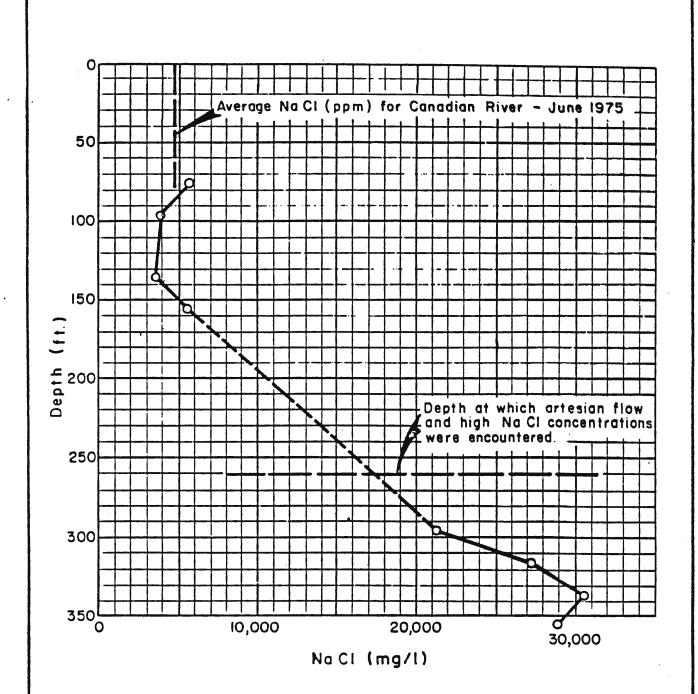
Figure 45. Stable isotopic distributions



SULFATE - CHLORIDE RATIOS OF SURFACE AND GROUNDWATERS



IODIDE - CHLORIDE RATIOS, THIS STUDY



UNITED STATES
DEPARTMENT OF THE INTERIOR
BUREAU OF RECLAMATION
LAKE MEREDITH SALINITY STUDY, TEX.-N.MEX.

Na Cl vs. Depth

DH-I

JUNE 1975

FIGURE 3

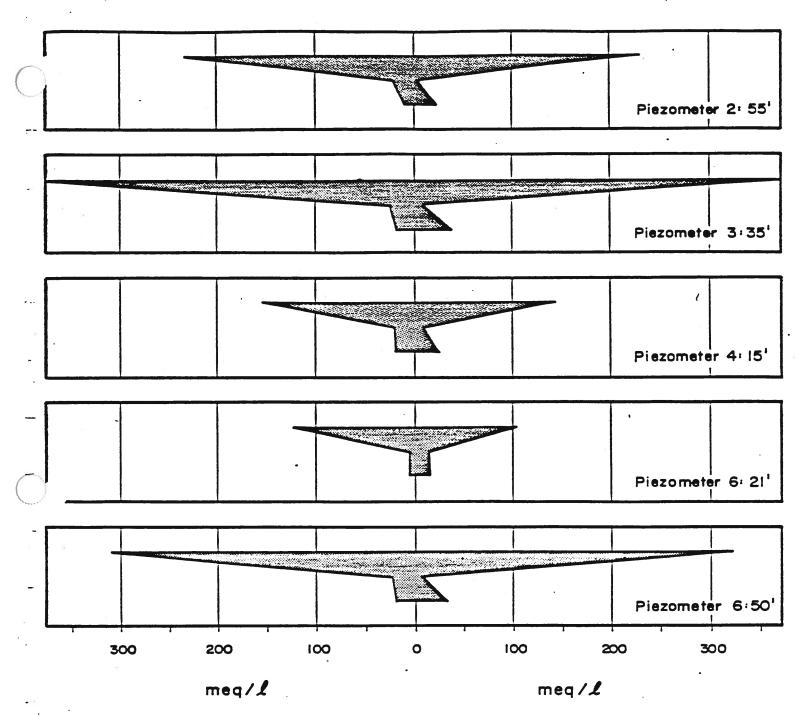
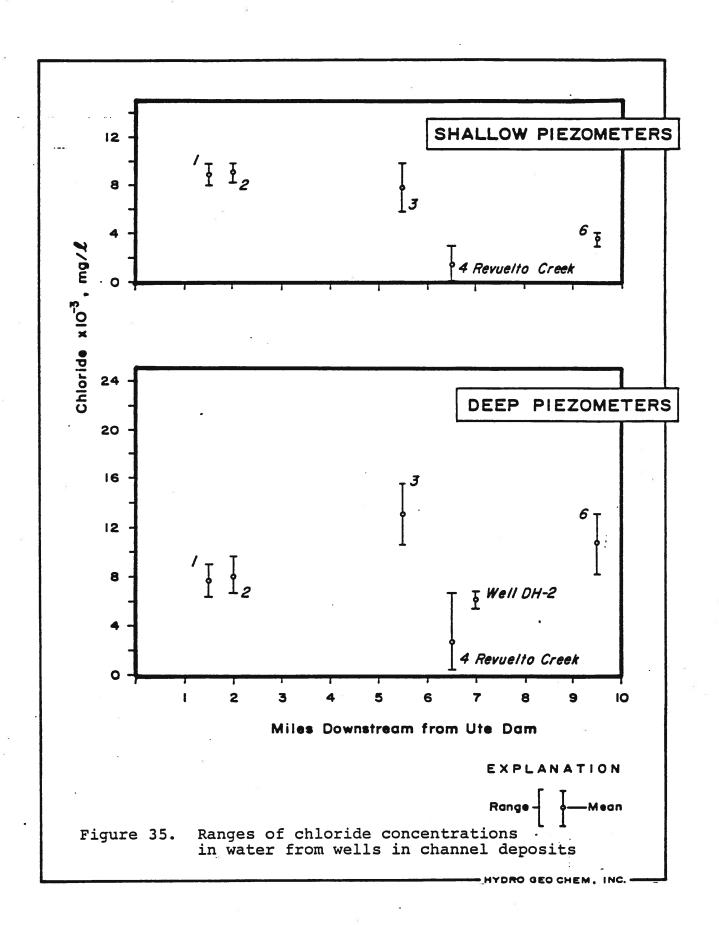


Figure 34. Stiff diagrams of water from wells in channel deposits



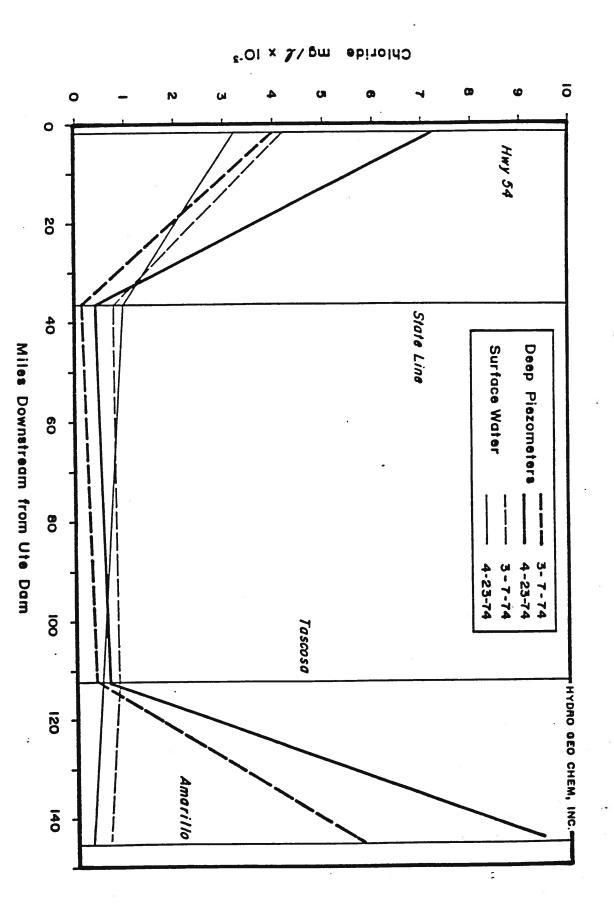
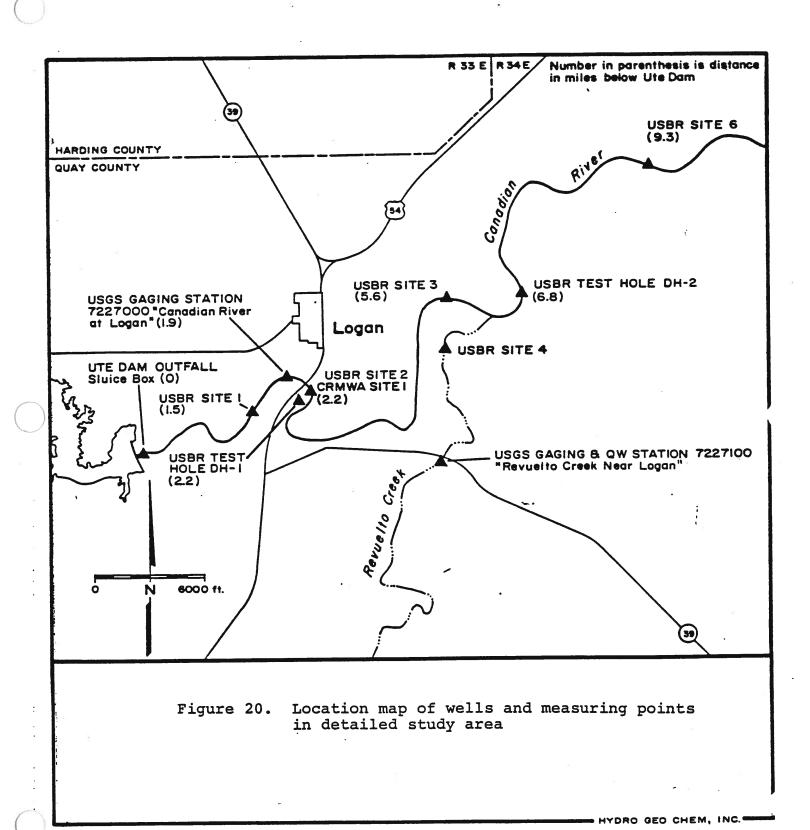


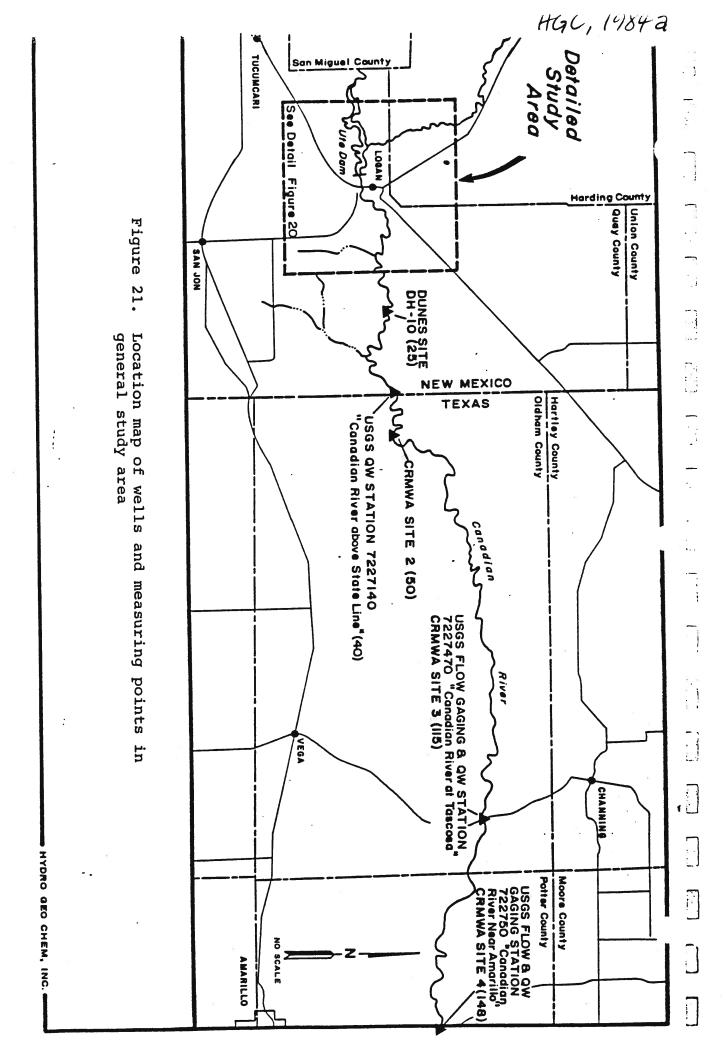
Figure 36. Ranges in chloride concentrations in CRMWA piezometers

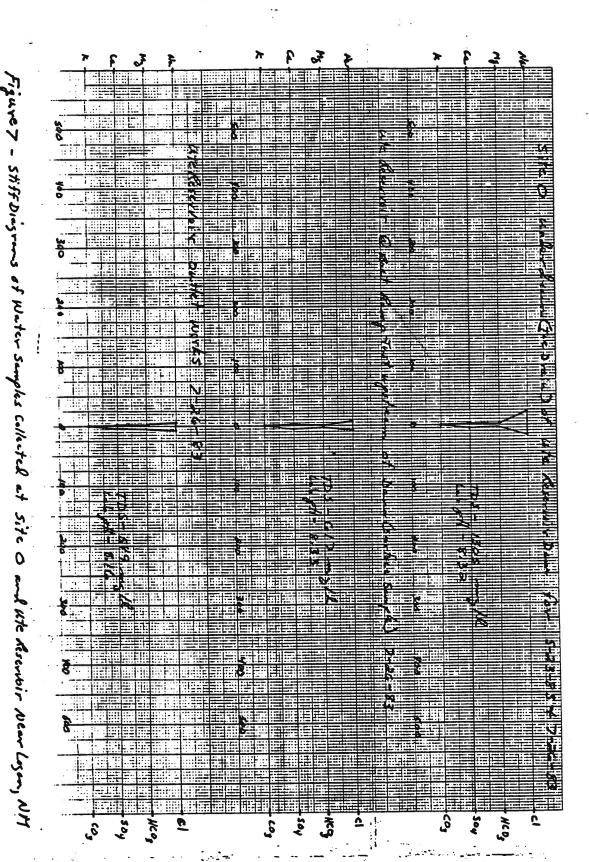
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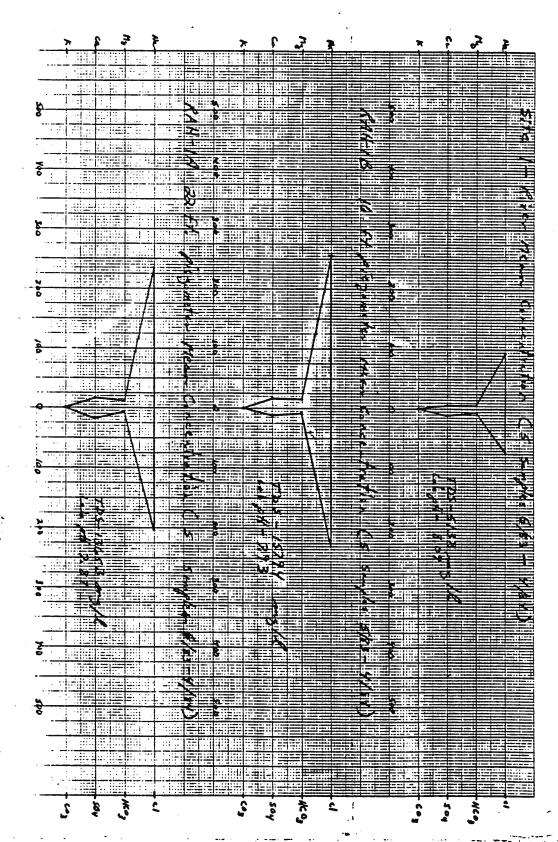
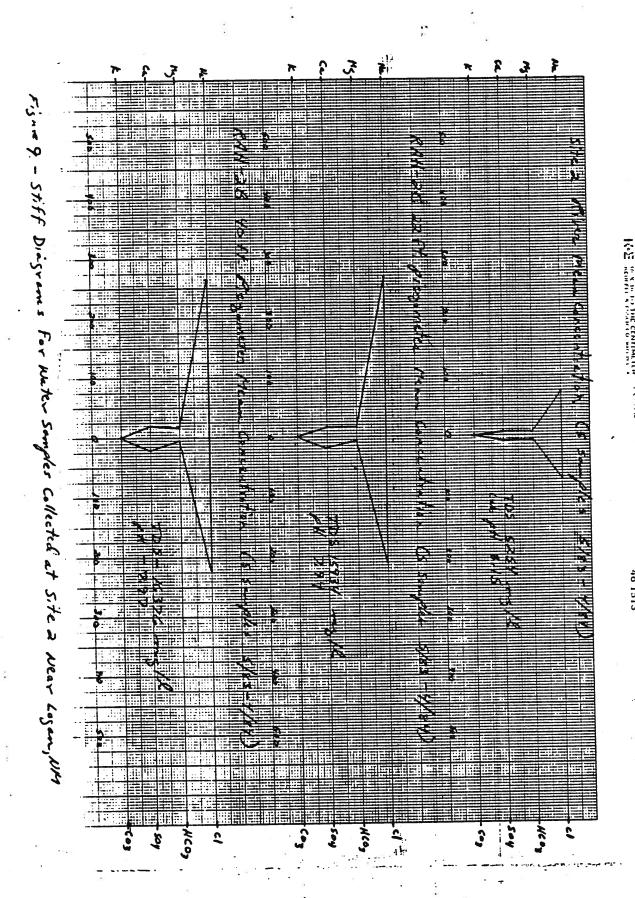
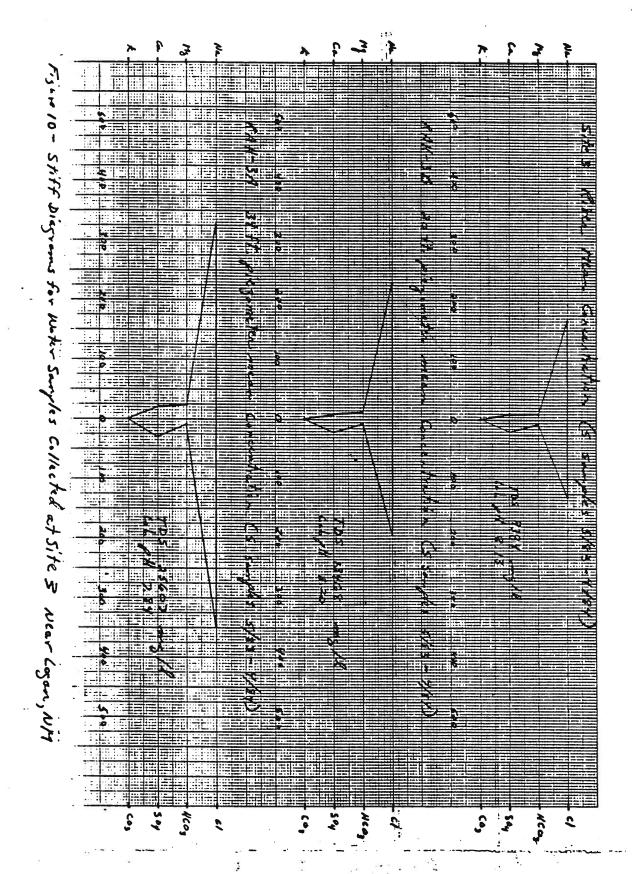


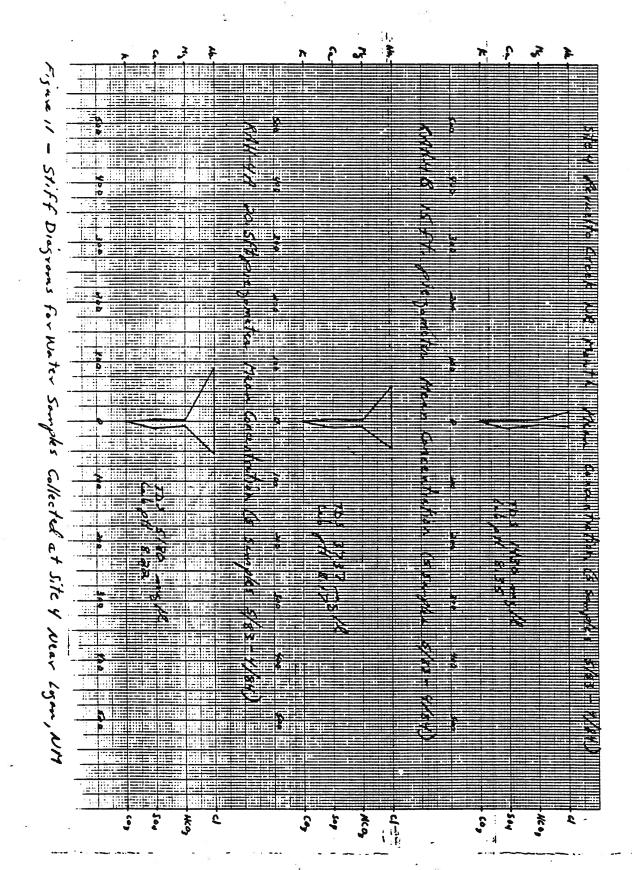
Figure 8 - Stiff Diagrams for Water Sumples Collected at Site 1 Near Logan, NM

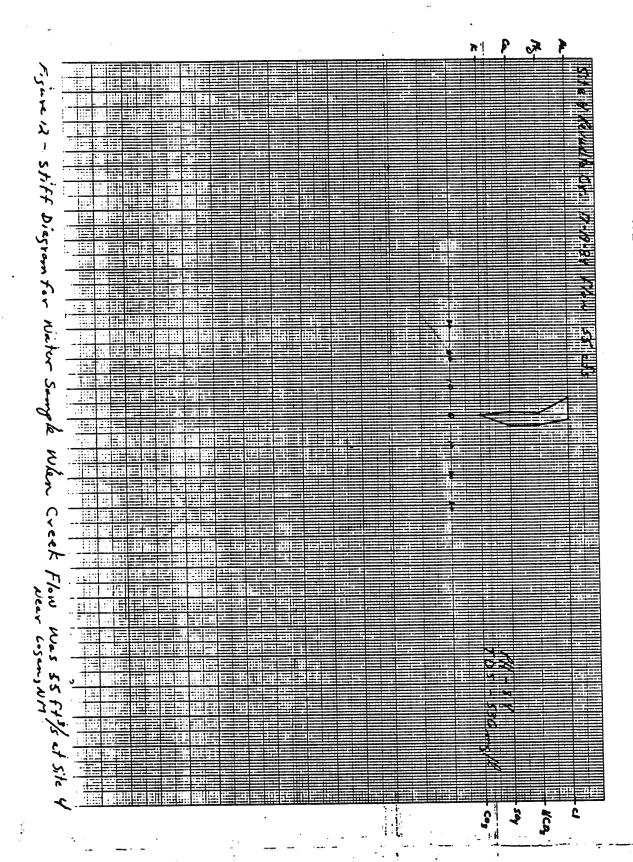


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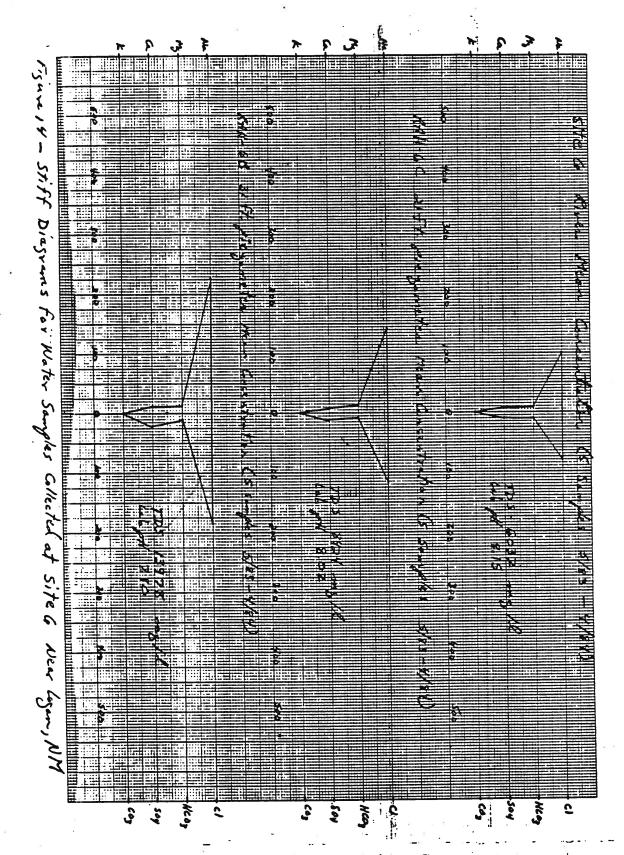
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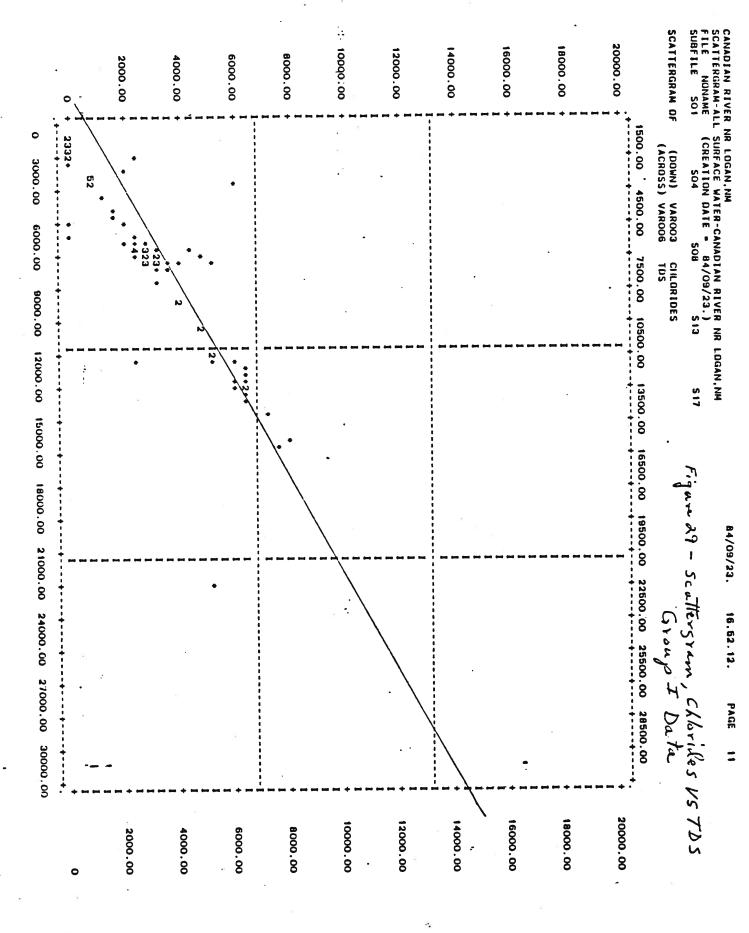


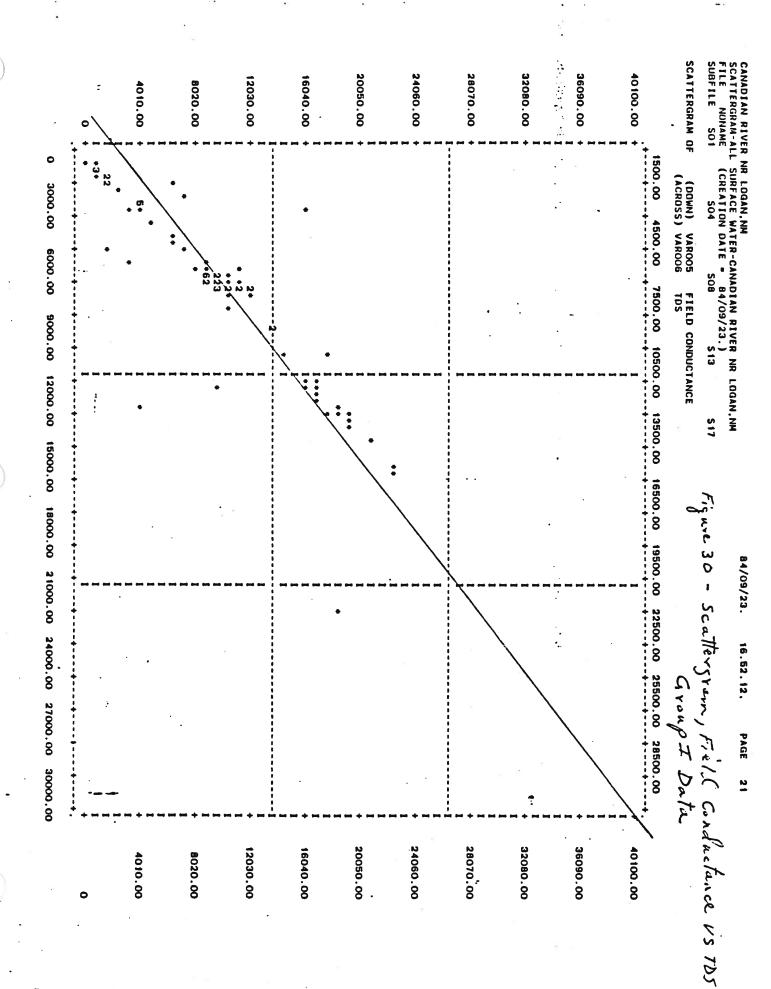
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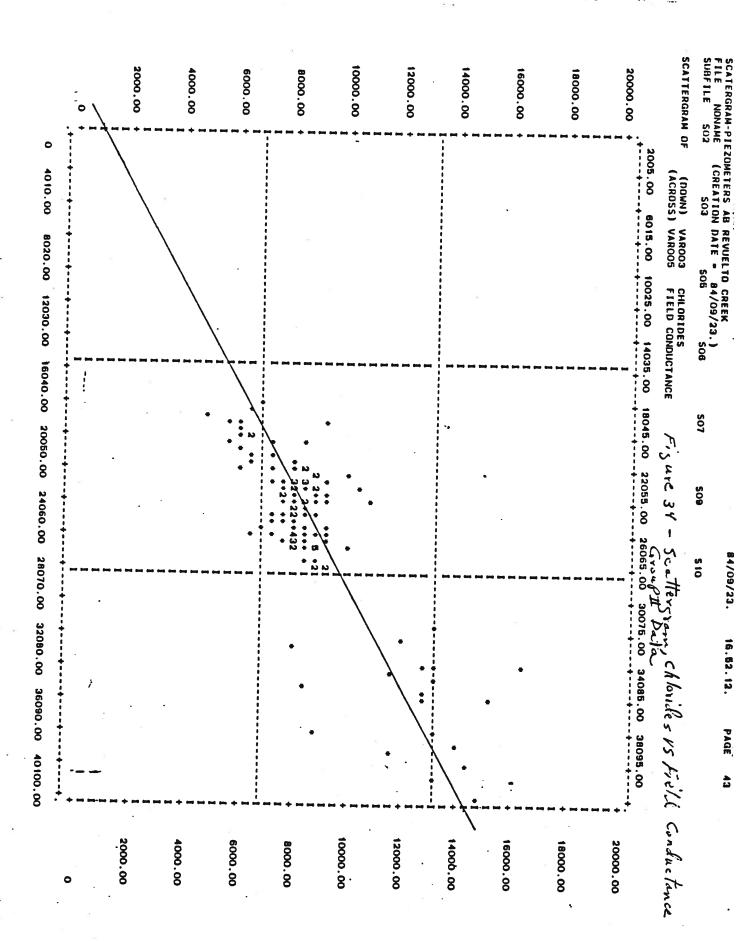


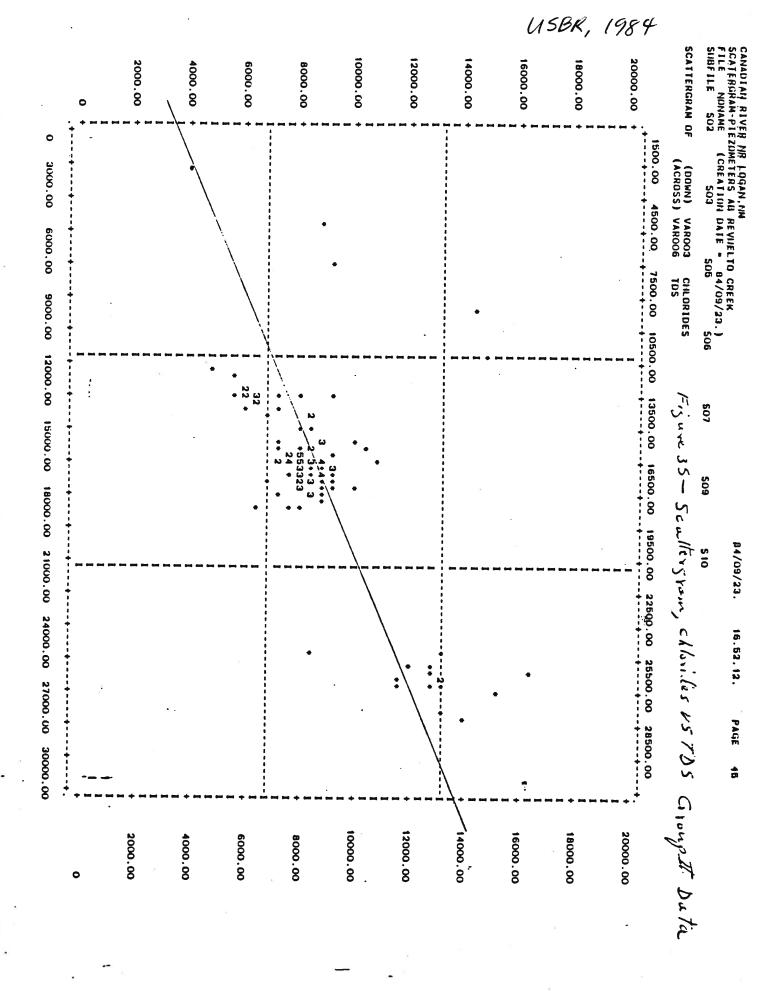
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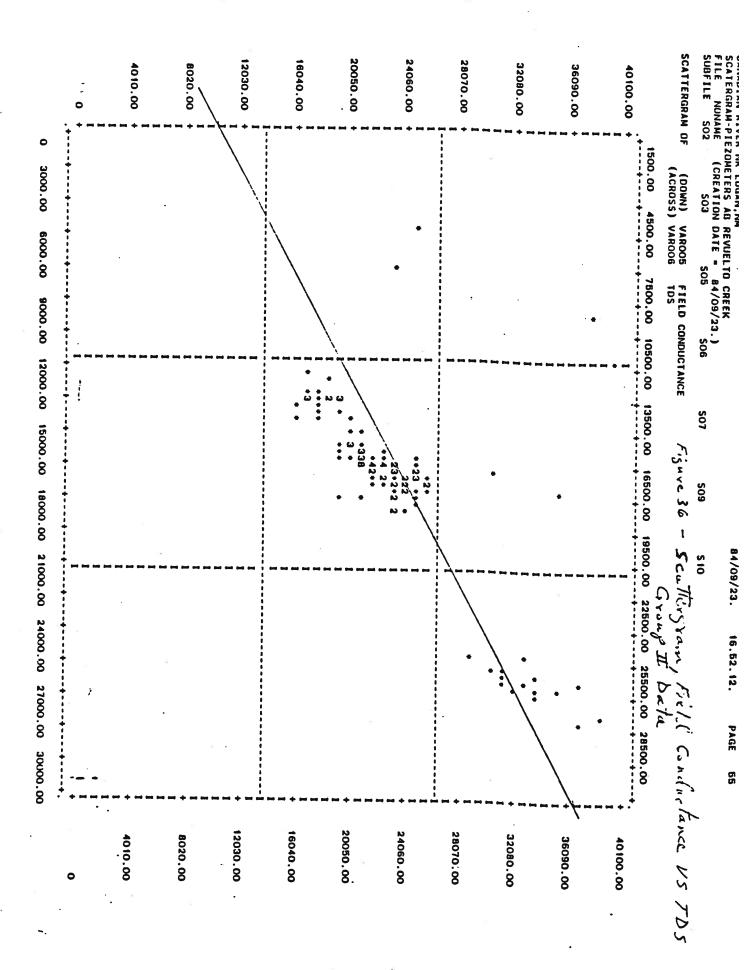
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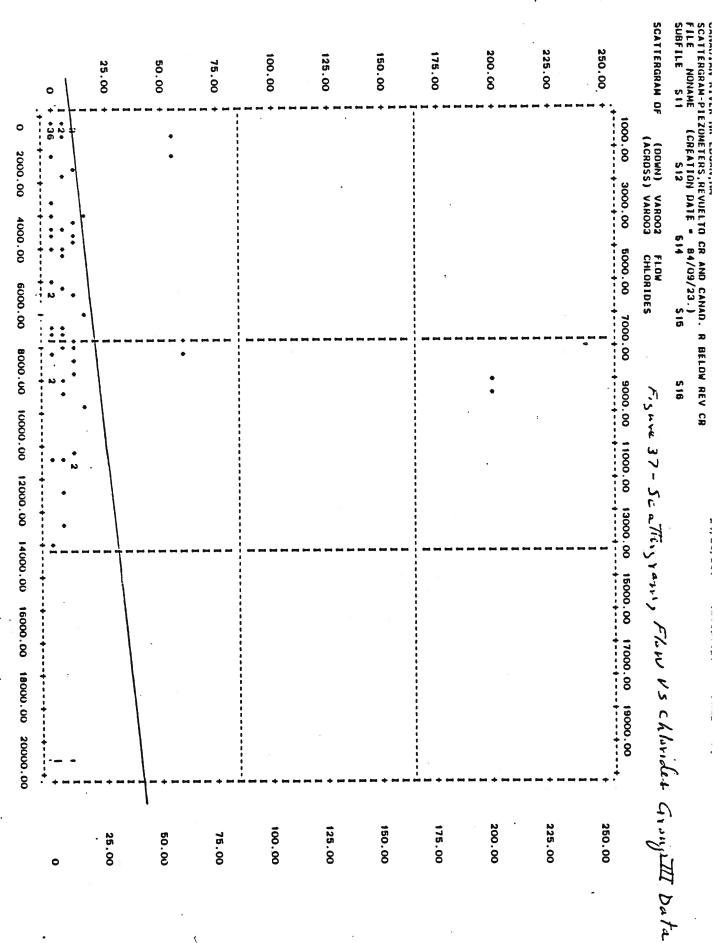
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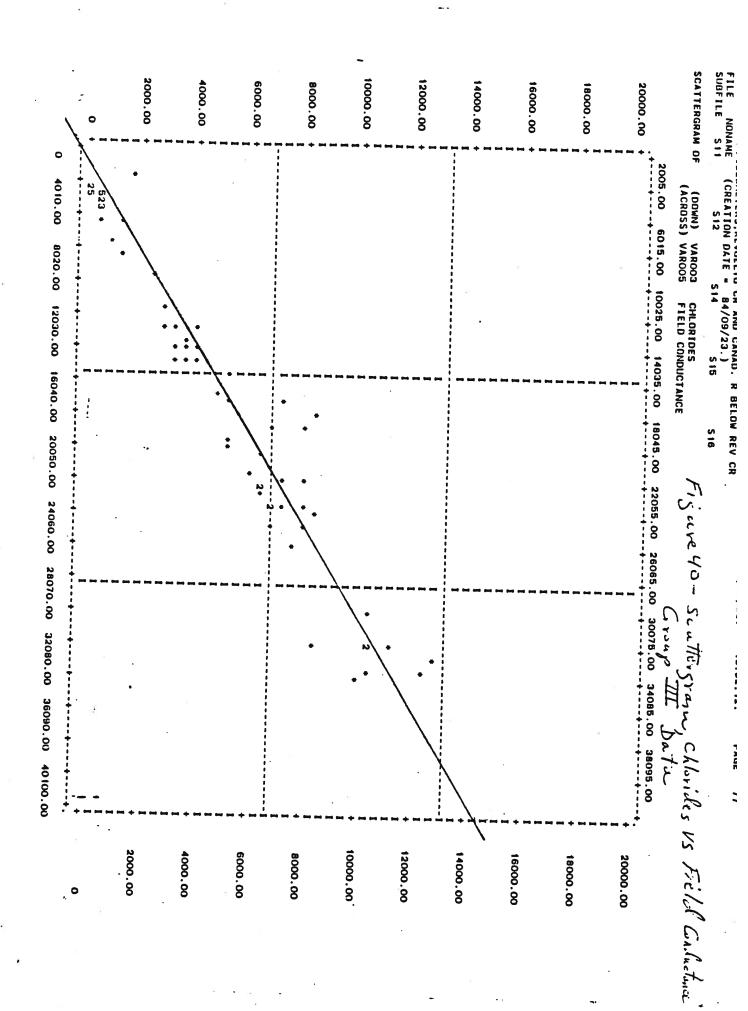


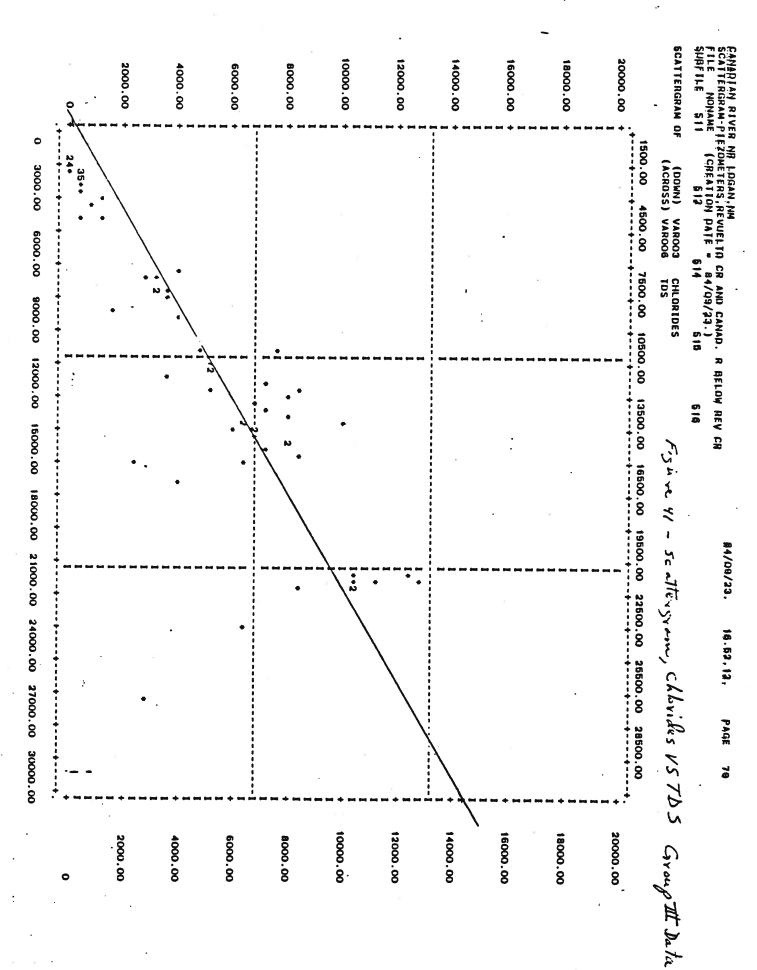


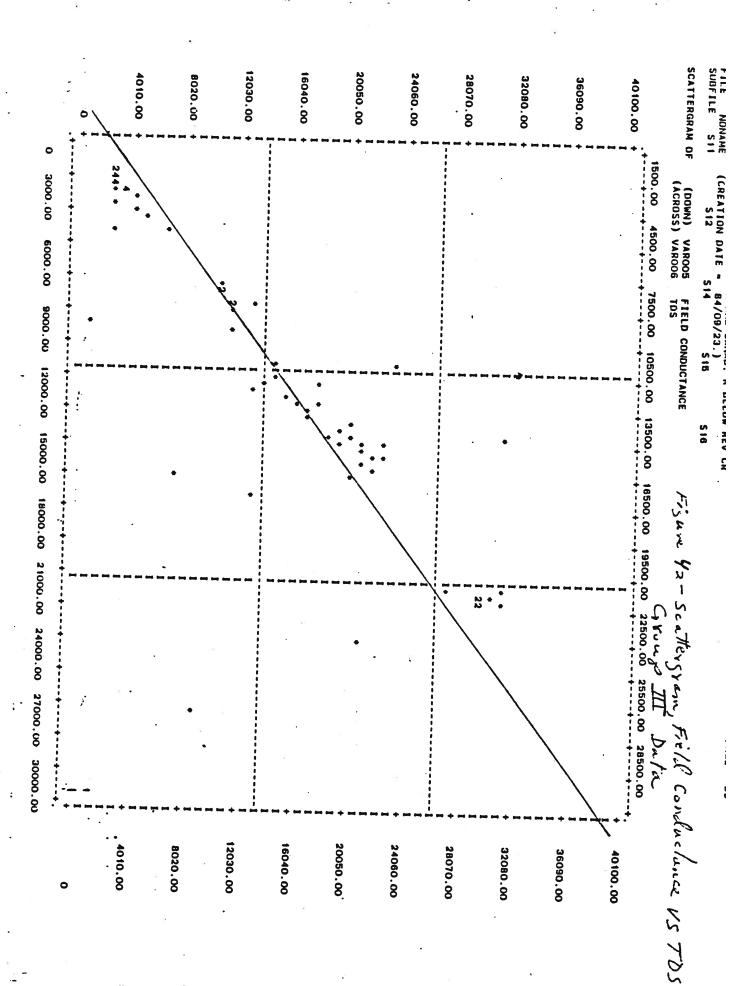
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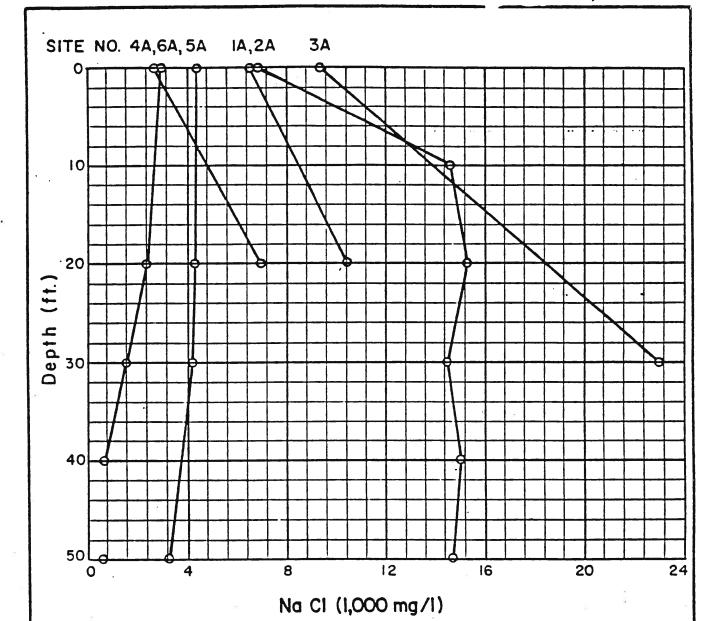
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	0.00 22500.00 25500.00 28500.00	13500.00 16500.00	7500.00 10500.00	(ACRUSS) VAROOG	
Group III Data	Scattergram, Flato 15 TDS	39-	814 (09/23.) 814 515	DATE	SCATTERGRAM-PIE FILE NONAME SUBFILE S11

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Na Cl (mg/l) vs. Depth for water samples from drill . holes in channel alluvium. Sampled 3/6/75

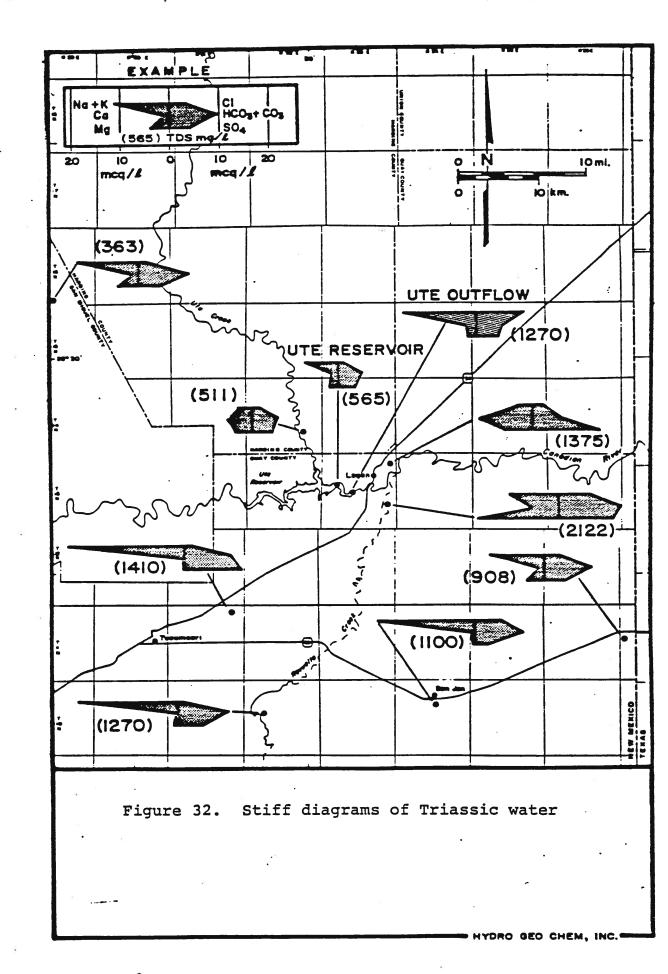
Site IA - I Mi. D/S from Ute Dam Site 2A - 2 Mi. D/S from Ute Dam Site 3A - 5 Mi D/S from Ute Dam Site 4A - 6 Mi D/S from Ute Dam Site 5A - II Mi D/S from Ute Dam Site 6A - 29 Mi D/S from Ute Dam

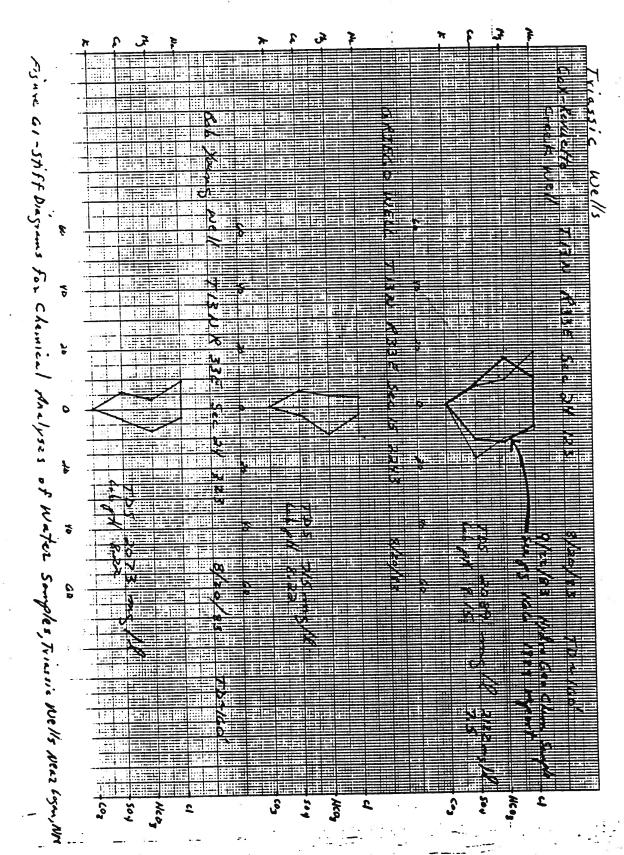
UNITED STATES
DEPARTMENT OF THE INTERIOR
BUREAU OF REGLAMATION
LAKE MEREDITH SALINITY STUDY, TEX.-M.MEX.

Na CI vs. Depth

FIGURE 2

TAB 14 Part B.2 Figures related to Triassic aquifers





TAB 14 Part B.3 Figures related to brine aquifers

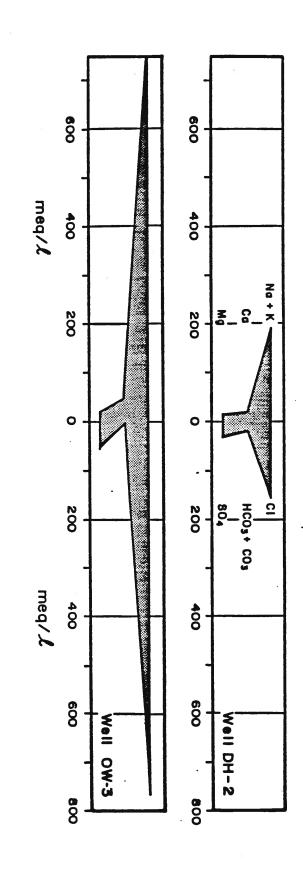
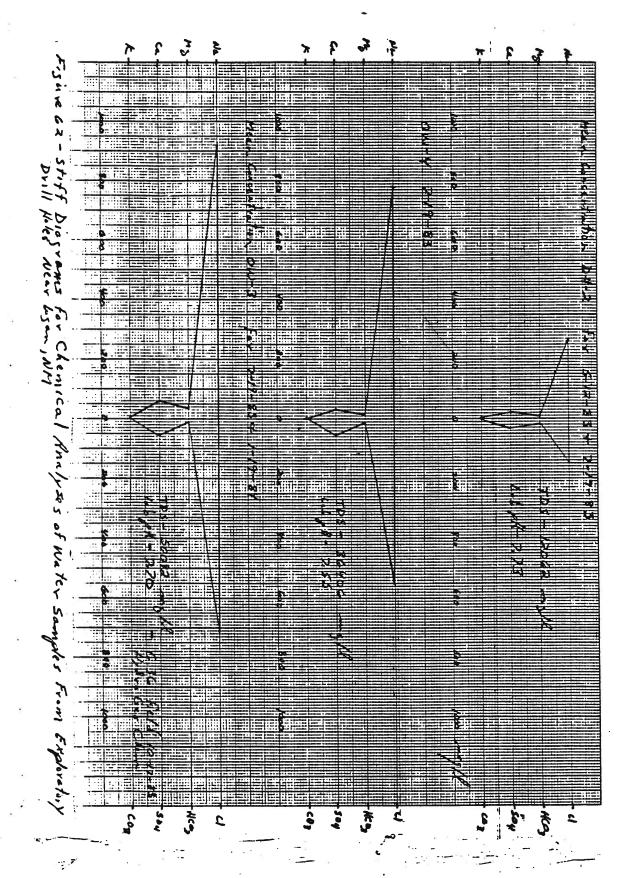


Figure 33. Stiff diagrams of shallow brine aquifer water

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TAB 14 Part B.4 Figures related to Permian aquifers

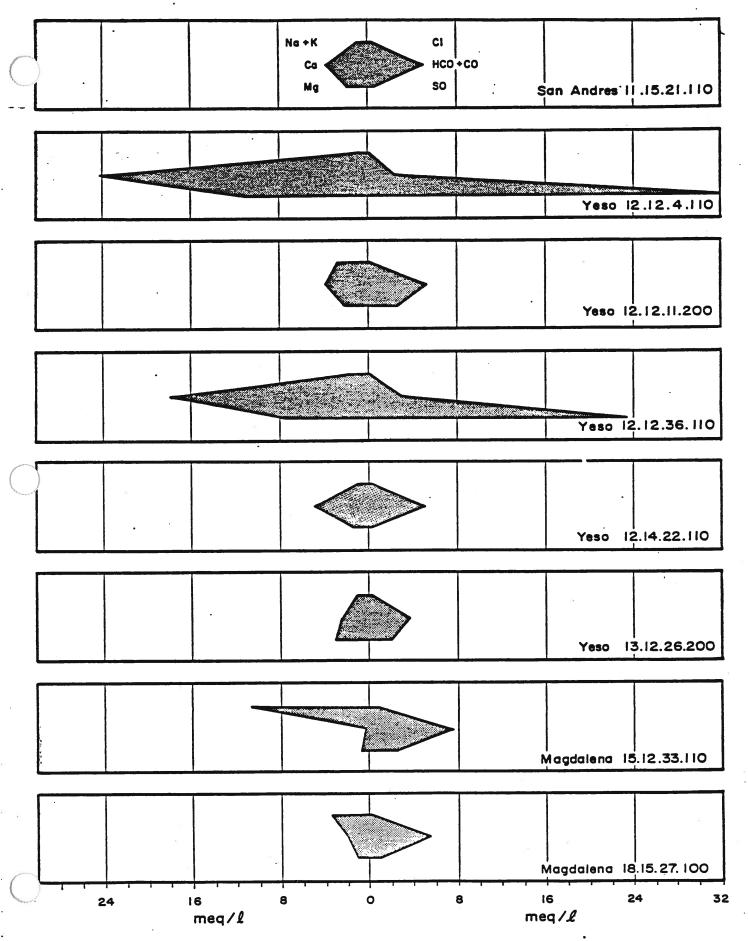


Figure 30. Stiff diagrams of Permian water near recharge area

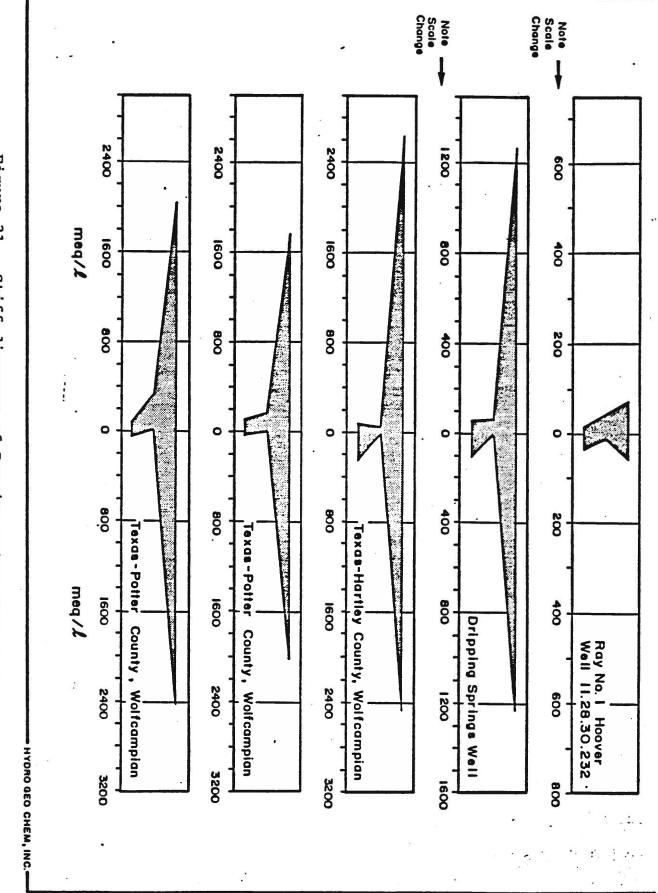


Figure 31. Stiff diagrams of Permian water near study area

TAB 14 Part B.5 Figures related to pre-Leonardian aquifer(s)

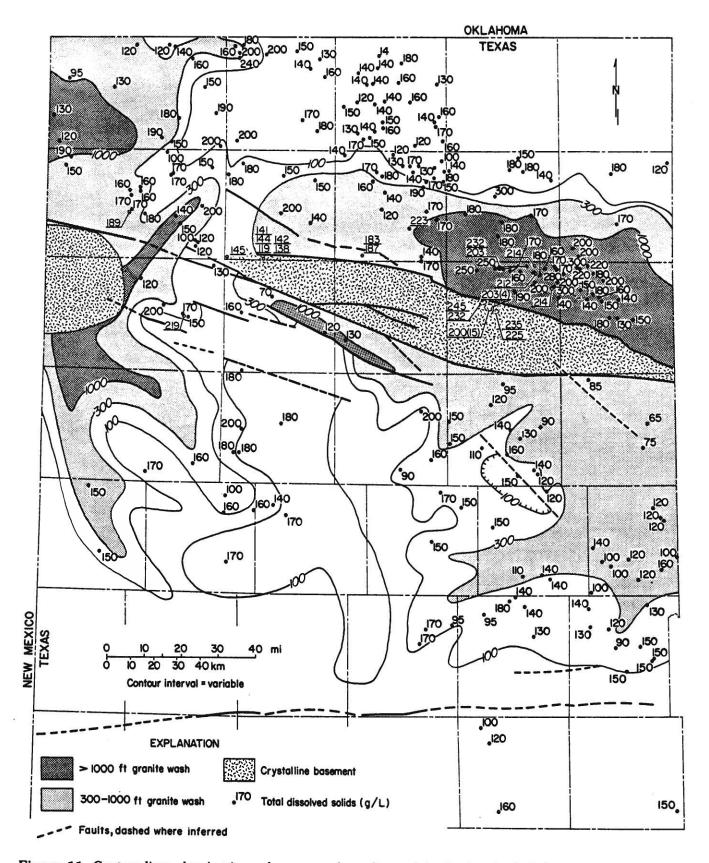


Figure 11. Contour lines showing isopachous map of granite wash in the deep-basin brine aquifer (from Dutton, 1982). Data points indicate total dissolved solids (g/L) from chemical analyses of the brines (see underlined data, this figure, which are from appendix B, table B-1) or computed from spontaneous potential logs (see appendix C, table C-2). Structure data are from A. Goldstein (written communication, 1981).



Table 1. Salt dissolution expressed as rates of horizontal and vertical dissolution. (Solute load data from U.S. Geological Survey, 1968-1977)*

	Mean annual		Annual	rates of		vey, 1906-19	Annual		
-	solute load x 10 ⁵ ft³			dissolution	-		vertical di ean	ssolution Max	Min
Basin	x 10 It	Me ft/yr	an cm/yr	Max ft/yr	Min ft/yr	1	x 10 ⁻³ cm/yr	x 10 ⁻⁵ ft/yr	x 10 ⁻⁵ ft/yr
	/F	10/ 91	CIII/ yi		10/ 91	X 10 10 J	a to cin/yi	x 10 10 y1	
1A Canadian River (Tascosa)	(5 years)** 4.460	0.00189	0.0576	0.00246	0.00132	1.0499	3.2001	1.367	0.735
•		0.00103	0.0570	0.00240	0.00132	1.0455	5.2001	1.001	0.100
1B Canadian River (Amarillo)	(5 years) 6.9542	0.00188	0.0575	0.00239	0.00081	1.0312	3.1431	1.306	0.452
1C	(3 years)	0.00100	0.0070	0.00200	0.00001	1.0012	0.1.01	2.000	0.102
Canadian River (Canadian)	7.9221	0.00186	0.0568	0.00261	0.00118	0.7665	2.3362	1.072	0.484
3 Salt Fork of the	(9 years)								
Red River (Wellington)	2.119	0.00621	0.1893	0.01265	0.00154	0.7405	2.2571	1.509	0.183
4A Prairie Dog Town Fork of the Red	(9 years)								
River (Lakeview)	24.1188	0.00963	0.2935	0.02337	0.00376	5.6674	17.2742	11.926	2.637
4C Little Red River	(9 years)								
(Turkey)	12.851	0.25353	7.7276	0.47850	0.13238	27.1130	82.6404	51.172	14.157
4D Prairie Dog Town Fork of the Red	(9 years)								
River (Childress)	119.5366	0.08485	2.5862	0.01925	0.00564	17.7560	54.1203	29.142	11.816
5A North Pease River	(5 years)								
(Childress)	4.3677	0.01077	0.3283	0.01607	0.00758	1.7911	5.4593	2.672	1.261
5B Middle Pease River (Paducah)	(5 years) 0.5515	0.00100	0.0305	0.00248	0.00018	0.2027	0.6177	0.500	0.037
5C	(8 years)								
Pease River (Childress)	32.5842	0.02408	0.7339	0.03318	0.01737	5.8465	17.8200	8.056	4.216
6-10 Area includes	(5-9 years)								
basins 6-10	115.5136	0.1249	3.8070	0.1735	0.0846	30.8860	94.1405	42.910	20.926
6 North Fork	(8. years)								
Wichita River (Paducah)	19.8165	2.6808	81.7108	3.2283	2.1093	. a		9	
8A South Fork	(6 years)			-					
Wichita River (Guthrie)	13.6156	0.2686	8.1870	0.3115	0.2229				
10B Salt Fork Brazos River	(9 years)								
(Peacock)	25.0487	0.0327	0.9967	0.0672	0.0087				
10C Croton Creek	(9 years)					*Prelin	ninary horizo	ntal rates in	Gustavson
(Jayton)	6.3678	0.0635	1.9355	0.1352	0.0218	and oth	ers (1979c) are edure. All diss	lower due to	a difference
10D Salt Fork	(9 years)					basin v	vere considere the basin.		
Brazos River (Aspermont)	70.1657	0.07216	2.1994	0.1061	0.0447	**Num	per of years of	data.	

